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EVALUATION OF SALT MARSH RESTORATION BY MEANS OF SELF-REGULATING TIDAL GATE – AVON ESTUARY, SOUTH DEVON, UK

Gerd Masselink, Mick E. Hanley, Anissa C. Halwyn, Will Blake, Ken Kingston, Thomas Newton, Mike Williams

Ecological Engineering

ABSTRACT

Salt marshes provide important regulating ecosystem services, including natural flood defence and carbon sequestration, which adds value to restoration and biodiversity offsetting schemes. This study evaluates the success of salt marsh restoration using a Regulated Tidal Exchange (RTE) system in SW England, i.e. a self-regulating tidal gate (SRT), in controlling the partial saline inundation of a 14-ha area of former salt marsh reclaimed for agriculture in 1760. A combination of (a) direct hydrodynamic monitoring of water and sediment flux and (b) repeat surveys to evaluate morphological and ecological (plants and foraminifera) changes over a 5-year period, was implemented immediately following SRT commissioning. Morphological changes were limited to the proximity of the SRT system due to limited sediment influx yielding sedimentation rates that were an order of magnitude below a nearby natural marsh. Ecological change to an ephemeral salt marsh community was only detected after 5 years of inundation cycles, with the delayed response attributed to (a) an initial limited tidal inundation due to conservative SRT settings, followed by (b) excessive inundation due to excessive rainfall and recurring SRT failure in an open position, and (c) a lack of sediment and propagule supply caused by (a) & (b) and the relatively narrow inlet pipe used in the SRT system. While the ecological response under optimum SRT settings was encouraging, the lack of perennial plants and limited foraminifera abundance demonstrated that the marsh was far from reaching natural status. We surmise that this is primarily due to inundation being more rapid than drainage leading to excessive submergence during a tidal cycle. Our study shows that the design of tidal inundation schemes requires a synergistic understanding of core ecological and geomorphological approaches to assess viability and success. We conclude that SRT can be a useful technique for intertidal habitat creation where there are significant site constraints (especially flood risk), but we need to be realistic in our expectations of what it can achieve in terms of delivering a perennial salt marsh community.

32 1. INTRODUCTION

33

34 Sea-level rise and an associated increased frequency and severity of storm surge events ([Martin et al.,](#)
35 [2011](#); [Zappa et al., 2013](#)) are together challenging the long-held view that so-called ‘hard’ engineering
36 alone can protect our coasts from flooding. Instead, an integrated strategy involving natural
37 ecosystems is increasingly held to offer the most cost-effective, sustainable, and effective form of
38 coastal defence ([Temmerman et al., 2013](#); [Bouma et al., 2014](#); [Hanley et al., 2014](#)). In north-west
39 Europe and North America, managed realignment (MR) or ‘de-embankment’ schemes (the deliberate
40 flooding of land situated behind coastal defences) are commonly implemented to create new areas of
41 salt marsh, both as compensation for habitat losses elsewhere, and to enhance flood defence and
42 create accommodation space ([French, 2006](#); [Spencer and Harvey, 2012](#); [Morris, 2012](#); [Foster et al.,](#)
43 [2013](#); [Chang et al., 2016](#)). This is viewed as a desirable outcome, not only to help redress the c. 50%
44 global loss of this habitat ([Adam, 2002](#)), but also because salt marshes have a remarkable capacity to
45 attenuate and dissipate wave energy, store flood waters, and so defend in-land areas from the worst
46 excesses of coastal flooding ([Gedan et al., 2011](#); [Moller et al., 2014](#)).

47

48 Effective intertidal habitat restoration, especially that of salt marsh, has, however, proven difficult to
49 achieve. In their study of 18 MR restorations in the UK, [Mossman et al. \(2012a\)](#) report that managed
50 realignment sites are typified by substantial recruitment failure (bare ground) and, where
51 revegetation has occurred, dominance only by early successional salt marsh communities. Although
52 to some extent these ‘failures’ represent the relative size and age of managed realignment sites
53 compared with adjacent natural marshes ([Wolters et al., 2005](#)), other environmental factors may be
54 important. Sites selected for restoration, usually for opportunistic reasons, are likely to start with
55 physical and biogeochemical conditions very different to natural counterparts ([Spencer and Harvey,](#)
56 [2012](#)). Many (typically) former agricultural sites are especially difficult to restore; livestock and farm
57 machinery cause soil compaction, reducing drainage and susceptibility to channel development, and
58 increasing waterlogging potential ([Spencer and Harvey, 2012](#); [Chang et al., 2016](#)). Long-term
59 agricultural use also leads to soil shrinkage and consolidation; this reduces surface elevation and
60 increases the amount of time the site spends under water post-breach ([Crooks et al., 2002](#); [Spencer](#)
61 [and Harvey, 2012](#)).

62

63 Elevation within the tidal frame is generally considered to be the pivotal factor determining the
64 success of salt marsh establishment ([Adam, 1990](#); [French, 2006](#); [Davy et al., 2011](#)). The duration and
65 frequency of tidal inundation has marked effects on propagule delivery, sedimentation, salinity and

66 soil redox potential, and, therefore, the regeneration potential of any newly arriving species to the
67 site ([Mossman et al., 2012a, b](#); [Spencer and Harvey, 2012](#)). Consequently, the most significant barrier
68 to effective salt marsh transition in managed realignment schemes is the difference in elevation
69 between reclaimed land and adjacent, natural salt marsh ([Wolters et al., 2005, 2008](#); [Spencer and
70 Harvey, 2012](#)). Despite this recognition, [Spencer and Harvey \(2012\)](#) highlight that, while considerable
71 attention has been devoted to monitoring long-term, post-breach shifts in plant and animal
72 communities, there has been minimal attempt to quantify the concomitant physical-chemical changes
73 that effect ecological transitions. There is also a surprising lack of clear demarcation of the optimum
74 tidal inundation characteristics for salt marsh development in managed realignment schemes; an
75 exception being [Ash and Fenn \(1997\)](#) who concluded, based on the Tollesbury MR scheme, that
76 mudflat, pioneer marsh, and mature marsh is characterised by > 38, 25–38 and < 25 inundations per
77 month, respectively. Additionally, according to [Environment Agency \(2003\)](#), salt marsh habitat
78 develops where there are < 42 inundations per month and where the surface has a small gradient (1–
79 3%).

80

81 One frequently used method in managed realignment schemes is to use Regulated Tidal Exchange
82 (RTE) systems which enable habitat creation behind coastal defences, whilst at the same time
83 managing flood risk ([Environment Agency, 2003](#)). RTEs are usually situated in a breach in existing
84 seawalls or embankments and utilise structures such as tide-gates and sluices, to tightly control the
85 amount of water entering the restored area ([Wolters et al., 2005](#); [Cox et al., 2006](#)). As a result, the
86 relative tidal height can be adjusted to match that experienced by local natural marshes, while
87 allowing plant propagules and sediment to enter; both of which are vital for salt marsh establishment
88 ([French, 2006](#); [Mossman et al., 2012b](#)). Buoyancy-controlled systems, or self-regulating tide-gates
89 (SRT), where a float system is adjusted to open the gate until a specified water level is reached, have
90 the further advantage that they can replicate the spring-neap system that drives natural zonation
91 patterns in salt marsh vegetation ([Adam, 1990](#); [Ridgway and Williams, 2011](#)). In theory, these designs
92 also ensure maximum control over water levels in the restored area to minimise flood risk to nearby
93 areas ([Adnitt et al., 2007](#)). In practice however, the elevation of the tidal frame is not chosen as SRTs
94 are often fitted to existing outfalls and the system may also be prone to mechanical problems (flotsam
95 clogging the mechanism). In summary, the efficacy of SRT in facilitating the development of emergent
96 salt marsh vegetation is largely unknown (cf. [Beauchard et al., 2011](#)).

97

98 Along with changes in soil structure and site elevation, SRT-imposed modification of tidal regime is
99 one of three key disturbances to natural physical parameters imposed upon managed realignment

100 schemes ([Spencer and Harvey, 2012](#)). As a consequence, the potential for development of a
101 functioning salt marsh ecosystem may be reduced and with it the ability of the restoration to deliver
102 the very ecosystem services for which it was implemented ([Mossman et al., 2012a](#); [Spencer and
103 Harvey, 2012](#)). To evaluate the ability of SRT to make a consistent and positive contribution to
104 successful salt marsh restoration, it is vital that post-breach plant community development is
105 understood within the context of the tidal environment in which it occurs ([Spencer and Harvey, 2012](#)).
106 The principal aim of this study is to evaluate the performance of a managed realignment scheme,
107 involving an SRT system, through addressing the following objectives: (1) quantify the tidal flooding
108 regime; (2) describe the tidal currents and sedimentation processes; (3) describe the morphological
109 changes; (4) describe the ecological changes, specifically in vegetation abundance and type, seed
110 deposition and foraminifera population; and (5) investigate how the ecological transitions are
111 associated with variation in the tidal environment imposed by the SRT system. We will argue that, at
112 least for our study location, a SRT system might not be the best means by which to control the tidal
113 inundation regime for salt marsh restoration.

114

115 **2. MATERIALS AND METHODS**

116

117 **2.1 Study area**

118

119 The Avon estuary is located on the south coast of Devon in south-west England (Figure [1](#)). It is a
120 relatively small estuary with a total surface area of 213.5 ha, of which 146.2 ha are intertidal, an
121 estuarine shoreline length of 19.8 km and a 7.8-km long tidal channel ([Davidson, 1991](#)). The estuary
122 has steep-sided margins and is generally considered a ria-type (drowned river) estuary, although it
123 does possess a sand barrier at its mouth ([Masselink et al., 2009](#)). The estuary is relatively pristine and
124 the only human modifications comprise several wooden groynes at its mouth (built in the 1930s and
125 now largely obsolete) and the reclamation of a 15-ha salt marsh in the upper estuary more than 100
126 years ago. It is the restoration of this reclaimed salt marsh, South Efford Marsh (Figure [1](#)), that is the
127 subject of this study.

128

129 Figure [1](#) here

130

131 Ocean tides in the Avon estuary are slightly less than at Devonport (Plymouth; closest Primary Port for
132 tidal data) and are characterised by a mean spring and neap tide range of 4.3 and 2.0 m, respectively
133 ([Uncles et al., 2007](#)). The elevation of the high water levels during spring and neap tides at the mouth

134 of the estuary are estimated at 2.5 and 1.2 m ODN (Ordnance Datum Newlyn, which is the mean sea
135 level datum in the UK), respectively. The 1:50 year storm surge height along the south coast of Devon
136 is c. 0.5 m (Lowe and Gregory, 2005).

137

138 Freshwater discharge into the Avon estuary is mainly through the Avon River, supplemented with
139 minor contributions from small streams draining the valley slopes. River discharge is measured by an
140 automatic gauging station at Loddiswell located c. 3 km upstream from the tidal limit (Figure 1). A
141 pronounced seasonal cycle in river discharge is apparent with a monthly-averaged winter discharge of
142 just under 7 m³/s and a monthly-averaged summer discharge of just over 1 m³/s (Uncles et al., 2007).
143 Peak flows during winter can exceed 40 m³/s and minimum flows in summer are often less than 0.5
144 m³/s.

145

146 Investigation of the tidal dynamics was conducted by Uncles et al. (2007) who deployed a current
147 meter in the upper estuary in the tidal channel near South Efford marsh over a 1-week period during
148 summer and winter in 2006. During summer, peak flood flows (0.5–0.6 m/s) were stronger than during
149 ebb (0.3–0.4 m/s) and the salinity profile was well-mixed during flood, but stratified during ebb. During
150 high river discharge flow in the winter however, flows in the upper estuary were almost exclusively
151 directed seaward with peak velocities of 0.4–0.7 m/s. Significant damping of the tidal wave occurs into
152 the estuary and for a typical ocean tide range of 3.7 m, the tide range at Bantham Harbour and South
153 Efford was 3.0 and 1.5 m, respectively. In addition, distortion of the tidal wave was also evident with
154 a duration of the flood phase at Bantham Harbour of only 4–5 hrs.

155

156 **2.2 History of South Efford marsh**

157

158 Around 1760, a large intertidal salt marsh area along the northern margin of the upper Avon estuary,
159 known as South Efford marsh, was reclaimed and converted to pasture through the construction of a
160 surrounding embankment. The embankment was breached in 1943 by a stray bomb and the site
161 reverted to intertidal habitats (mostly mud and sand flats) until the breach was repaired in 1956 and
162 the site returned to agricultural use. In 2011, the Environment Agency (EA) installed a self-regulating
163 tidal gate (SRT) at the southern end of the marsh, utilising the existing outfall with improvements to
164 its headwall, allowing the tide to once more flood the area. Flooding of the marsh by means of an SRT
165 was preferred to breaching the embankment and allowing the breach to evolve naturally, because
166 approval for the restoration scheme was contingent upon not increasing the flood risk to several
167 properties at the back of the north-east corner of the marsh. The SRT has been operational since May

168 2011 and the EA expected that within 5 years of commissioning the lowest part of this area would
169 revert back to an intertidal salt marsh environment, whilst the upper part remains supratidal grazing
170 marsh. The aim was to generate 7 hectares of new intertidal habitat, including both salt marsh and
171 mudflat.

172
173 The LIDAR-derived digital elevation model (DEM) of a 1.2 km x 1.2 km area encompassing South Efford
174 marsh is shown in Figure 1, as well as an aerial photograph. Most noteworthy is the observation that
175 the elevation of the natural marsh to the south ($z = 1.5\text{--}1.6$ m ODN) is at least half a meter higher than
176 that of the realigned marsh surface ($z = 0.9\text{--}1.0$ m ODN). This is due to the fact that the natural salt
177 marsh continued to accrete after South Efford marsh was reclaimed in 1760, possibly exacerbated by
178 soil compaction. Ignoring compaction, an average accretion rate on the natural salt marsh of at c. 3
179 mm/year can be deduced, which corresponds to previous estimates made using salt marsh cores from
180 the Avon estuary (Bugler, 2006). The elevation of the embankment around the marsh site is > 3.5 m
181 ODN, which is higher than the highest water level in the estuary at this location. Some indication of
182 the pre-reclamation tidal creek topography is discernible in the DEM (dark blue meandering patterns)
183 and the low area around Easting = 268600 m and Northing = 46800 m is related to a WWII bomb that
184 struck the embankment at this location. The north-east part of South Efford marsh is significantly
185 higher ($z = > 1.5$ m ODN) than the rest of the marsh. The SRT is located at the south-west end of the
186 marsh and leads straight into the straight ditch that runs along the centre of the marsh.

187

188 **2.3 Self-regulating tidal gate (SRT)**

189

190 The tidal dynamics in the realigned marsh, and the ensuing morphological and ecological changes,
191 reported here are controlled by the design and the settings of the SRT and the way through which it
192 controls tidal exchange between the marsh and estuary. Flooding and draining of the realigned marsh
193 occurs by mean of a concrete pipe (diameter = 0.9 m; elevation of base of pipe = 0.39 m ODN) whose
194 connection to the river is controlled by a rotating tide gate to which floats are attached. The timing of
195 opening and closure of the gate depends on the river water level. At low tide, the gate is closed and
196 water leaves the marsh through a side flap if the marsh level exceeds the river level. As the tide rises,
197 the side flap closes under pressure when the river level exceeds the marsh level; the floats cause the
198 SRT to rotate and the aperture begins to align with the pipe, allowing water to enter the marsh. As the
199 tide continues to rise, the SRT rotates until fully open and then gradually closes again until the river
200 reaches a pre-determined level. The SRT is adjusted so that it is fully closed at high tide when the
201 estuary water level. If uncontrolled, would increase flood risk to dwellings. A conservative setting will

202 cause the gate to be closed well before high tide level is reached, while a less conservative setting will
203 result in more delayed gate closure, or no closure at all (e.g., during neap high tide). As the tide falls,
204 the gate will open again and more water may enter the marsh, until the point when the tide drops
205 below the marsh water level and the marsh starts to drain again.

206

207 A very important implication of this system is that the duration over which the SRT is open, and
208 therefore the amount of water that will flow through the pipe to flood the marsh, is controlled by how
209 long it takes for the river to reach the pre-set gate-closure level: especially for a conservative gate
210 setting, the pre-set water level for gate closure is reached relatively early during spring tides as the
211 tide rises, whereas the pre-set water level for gate closure is reached relatively late in the rising tide
212 during neap tides. Depending on the gate settings, this can mean that spring tides in the river or
213 estuary generate lower water levels in the marsh than neap tides; in other words, a reversed neap-to-
214 spring tidal variation can be generated in the realigned marsh. Another characteristic of the SRT
215 system is that, frequently, the SRT was not working properly and the gate was open almost
216 continuously. Such malfunctioning results in a progressive increase in water level and salinity in the
217 realigned marsh, up to the point that the marsh ceases to drain. This can result in extended periods
218 of marsh submergence which can have significant implications for the ecological development.

219

220 The SRT fitted at South Efford differs from others of a similar design in that, to accommodate the need
221 to open and fully close again over quite a small tidal range, the gate and floats are linked by a
222 mechanism that causes the gate to rotate roughly twice as fast as the float. It is possible to make
223 adjustments to the gate to change to points of opening and closure, but such adjustments are made
224 very infrequently, one of the aims of the design being that of requiring minimal intervention. The
225 South Efford SRT also appears to be more prone to malfunctioning than others of the type. This seems
226 to be as a result of two main factors: (1) the location (facing down the estuary and in a slight
227 backwater) means that debris is prone to collecting around the SRT; and (2) the linkage connecting
228 the rotating gate and the floats means that relatively small pieces of debris can affect the smooth
229 operation of the gate.

230

231 **2.4 Survey grid and monitoring programme**

232

233 Pre-breach, a measurement grid was established using laser total station (Figure 2). This measurement
234 grid has as its origin the location of the SRT, and the x- and y-axis represent the length- and width-axis
235 of the marsh, respectively. A total of 10 across-marsh measurement transects were established (y =

236 25, 50, 100, 150, 200, 300, 400, 500, 600, 700 m) and an additional transect crossing the natural salt
237 marsh to the south of the realigned marsh ($y = -60$ m). Benchmarks were established along the margin
238 of the site and these are used for re-sectioning the total station during surveys. Ecological monitoring
239 sites were located within this measurement grid and are also indicated in Figure 2.

240

241 Figure 2 here

242

243 An extensive data set was collected over a 5-year monitoring period (May 2011 – May 2016) and
244 various parameters were recorded, as discussed below. An overview of these parameters and their
245 sampling frequency is provided in Table 1.

246

247 Table 1 here

248

249 **2.5 Tidal flooding regime**

250

251 Water level was recorded at either side of the SRT using pressure sensors and salinity in the realigned
252 marsh was monitored with a conductivity sensor deployed just inside the gate. The sensors were
253 installed and are maintained by the EA, and data collected every 15 min. The water level inside and
254 outside South Efford marsh is referred to as the marsh level and the river level, respectively.

255

256 The time series of the river water level was used to compute the inundation characteristics of the
257 natural tidal flat or mudflat (TF; mean elevation $z = 1.0$ m ODN) and the natural salt marsh (SM; mean
258 elevation $z = 1.5$ m ODN), and the time series of the marsh water level was used to compute the
259 inundation characteristics of the realigned marsh (RM; mean elevation $z = 0.8$ m ODN). The following
260 parameters were computed for every month for which reliable water level data were available (56
261 and 55 months out of 60 months for TF/NM and RM, respectively): (1) the number of over-tides N_{otides} ,
262 i.e., tides that inundate the RM, NM and RM surface; (2) the total number of hours that the surface is
263 submerged $T_{sub,tot}$ (3) the maximum continuous period of tidal submergence $T_{sub,max}$; (4) the maximum
264 continuous period of exposure $T_{exp,max}$; (5) the mean water depth over the surface h ; and (6) the
265 average rate of the falling tide over the surface dh/dt . Note that these parameters are monthly values.

266

267 **2.6 Tidal currents and sedimentation**

268

269 Two self-recording Acoustic Doppler Velocimeters (ADV – Nortek Vector) with external pressure

270 transducer (PT – Druck PTX 1830) and optical backscatterance sensor (OBS – OBS-3 Downing) were
271 deployed from 3 March to 4 May 2012 to record current velocities and suspended sediment
272 concentrations. One of the instruments was deployed in a tidal channel located in the natural salt
273 marsh just outside the SRT linking the Avon estuary with the realigned marsh. The other instrument
274 was mounted above the SRT to record the flows inside the pipe through which water is exchanged
275 between the Avon estuary and the marsh. In the afternoon of 21 March 2012, water and suspended
276 sediment samples were collected from the natural salt marsh through pumping. Suspended sediment
277 concentrations were used to carry out in-situ calibration of the OBS sensors. Combining suspended
278 sediment concentrations with flow velocities enables quantification of suspended sediment fluxes,
279 and from these potential marsh accretion rates could be estimated.

280 Two squares of ‘Astroturf’ matting (21 cm x 21 cm) with 1.5 cm plastic tufts ([Lambert and Walling,](#)
281 [1987](#)) were secured to the ground surface with a 30 cm steel peg adjacent to each other in the
282 southwestern corner of every 4 m² quadrat on the -60, 25m, 50, 100, 150, 200 and 300 m transects.
283 We focussed on the lower marsh area due to its greater dynamicity and sensitivity to hydrodynamic
284 variation ([Coulombier et al., 2012](#)). Mats were collected and new ones deployed at the same stations
285 for periods of 141–203 days from 2011 to 2014. Once collected and transported for analysis in
286 separate, sealed plastic bags, one mat (of known mass) from each pair was used to determine dry
287 weight (the second mat was used for studying seed deposition and germination; cf., Section **2.7**). Mats
288 were dried for 48 hours at 55°C and reweighed once cooled to determine mass change (i.e. total
289 weight of deposited sediment during deployment). To account for variation deployment time,
290 sediment accumulation was standardized by expressing as g/m²/d⁻¹ ([Reed et al., 1997](#)).

291

292 **2.7 Morphological change**

293

294 Using a total station, surveys of the 11 transects across the realigned marsh were conducted annually
295 to record morphological changes in the form of sedimentation and creek development (Figure **2**).
296 These surveys were always conducted concurrent with the vegetation surveys (refer to Section **2.4**).
297 Although the accuracy of the total station is several mm’s at the most, the actual survey accuracy of
298 the marsh surface is considerably less due to the disturbance of the marsh by cattle and the presence
299 of vegetation, and is considered several cm’s at most. To complement, and concurrent with, the
300 annual total station surveys, a Scan Station 2 terrestrial laser scanner was used to conduct a complete
301 survey of a 100-m radius area in the vicinity of the SRT, with specific focus on any evolution of drainage
302 channels on the realigned marsh. The scanner was installed on top of the embankment at the SRT and
303 acquires a full 360° span of data with a prescribed horizontal grid resolution of 20mm x 20mm. At the

304 end of the 5-year monitoring period, an Unmanned Aerial Vehicle (UAV) was used to acquire high-
305 resolution aerial photographs of the realigned marsh. The data were acquired at low tide, and, with
306 the aid of a large number of ground control points and Structure-from-Motion algorithms, individual
307 photographs were combined to obtain a fully georeferenced DEM of the marsh.

308

309 **2.8 Ecological variability**

310

311 Vegetation surveys were undertaken from 2011 to 2016, using four 4m² quadrats randomly positioned
312 along each transect line, ensuring two quadrats lay either side of the central channel (Figure 2). This
313 arrangement ensured that vegetation across a range of elevations and geomorphological settings was
314 monitored. Surveys were conducted in June 2011 (pre-flooding), October 2011 and then every year in
315 June until 2016. During surveys, the percentage cover of bare ground, dead vegetation and all
316 component species was noted. The species composition in each quadrat was classified according to
317 the NVC Community Type scheme (Table 2; Rodwell, 1992, 2000); only four of the NVC types were
318 observed more than once during the survey period (MG10 was only observed in one transect during
319 one survey).

320

321 Table 2 here

322

323 The second 'Astroturf' mat (cf. Section 2.6) was used to determine the number and species of
324 deposited viable seeds (Goodson et al. 2003). The purpose of this element of the monitoring is to gain
325 insights into whether seeds from the natural salt marsh are imported into the restored salt marsh
326 area. Mats were placed in seed trays (size) filled with potting compost in an unheated, naturally lit
327 greenhouse; a thin layer of vermiculite on each mat prevented desiccation. On germination, seedlings
328 were identified and removed for a maximum of 10-weeks after mat recovery (Goodson et al., 2003).

329

330 Foraminifera samples were collected and analysed to complement the vegetation surveys and provide
331 a more comprehensive overview of the ecological change in the realigned marsh. Sediment samples
332 were initially collected at 3-monthly intervals until 2013, but subsequently less frequently. Three
333 samples were collected each time: one from the 25-m transect (closest to the central creek) and two
334 from the 100-m transect (at either side of the central creek). Samples were wet sieved following
335 Gehrels (2002) and Rose Bengal stain was used to discriminate living and dead foraminifera following
336 Walton (1952). From each sample, 5cc was examined under light microscopy with all foraminifera
337 encountered identified to species level and the results are expressed as tests per cc (or per cm³). A

338 total of 14 species were recorded over the 5-year monitoring period and these were typical of those
339 common to natural salt marsh and estuarine/mudflat assemblages elsewhere in south-west England
340 (Gehrels et al. 2001; Massey et al. 2006; Hart et al. 2014, 2015). On this basis, the foraminifera
341 assemblages were divided into three groups indicating the typical sub-environment with which each
342 community is associated: mudflat, low salt marsh, and high salt marsh (Table 3). Living foraminifera
343 assemblages closely matched death assemblages at each monitoring station which suggests that the
344 deceased foraminifera are autochthonous populations. Due to seasonal bias in living foraminifera
345 populations (cf. Horton and Edwards, 2003), the entire foraminifera (live and dead) assemblage is
346 considered at each station.

347

348 Table 3 here

349

350 3 RESULTS

351

352 3.1 Tidal flooding regime

353

354 Figure 3 shows the 5-year time series of the daily discharge of the River Avon, and the water levels
355 recorded at either side of the self-regulating tidal gate, representing the tidal motion in the estuary
356 and in the realigned marsh. The river discharge is highly seasonal, showing maximum daily discharge
357 during the winter months, peaking at almost 60 m³/s during the 2012/13 winter. Peak discharge was
358 not exceptional during the 2013/14 winter, but during this period, which was the wettest winter on
359 record (Matthews et al., 2014), river discharge remained persistently high (> 10 m³/s) for almost the
360 whole winter period. The pressure sensor in the river was installed above MSL (0 m ODN) and only
361 captured the upper half of the tidal curve. Water levels > 2.5 m ODN were experienced during most
362 of the spring tides and water levels > 3 m ODN occurred when spring tides coincided with larger river
363 discharge (e.g., during 2012/13 and especially the 2013/14 winter). As mentioned earlier, water levels
364 in the realigned marsh are much lower than in the estuary and rarely exceeded 1.5 m ODN. The highest
365 marsh water levels coincided with large river discharge during 2013/14.

366

367 Figure 3 here

368

369 The extent of tidal inundation for the different water levels is illustrated in Figure 4, which shows the
370 water depth across the realigned marsh for water levels of 1, 1.25 and 1.5 m ODN. For these water
371 levels, 5, 11 and 14 ha of the marsh is submerged, respectively, with maximum water depths across

372 the marsh surface (not the creeks) of 0.2, 0.45 and 0.7 m. The tidal prism associated with marsh levels
373 of 1, 1.25 and 1.5 m ODN is 6.2×10^3 , 2.8×10^4 , $5.9 \times 10^4 \text{ m}^3$, respectively. For a marsh level of 0.9 m ODN,
374 the inundation area is < 2 ha (not shown); therefore, for a significant area of the marsh to be flooded,
375 the water level must reach at least 1 m. It is evident from Figure 4 that for water levels higher than 1
376 m ODN, tidal flooding extends right up to the back of the marsh and could potentially increase flood
377 risk to the properties at the north-east corner of the site.

378

379 Figure 4 here

380

381 The monthly tidal inundation characteristics for the complete 5-year monitoring period are shown as
382 boxplots in Figure 5. The number of over-tides per month is largest for the tidal flat ($N_{otides} = 56$) and
383 smallest for the realigned marsh ($N_{otides} = 34$), and the water depth over the tidal flat ($h = 0.63$ m) is
384 also larger than over the realigned marsh ($h = 0.16$ m). Despite the smaller number of tides and
385 shallower water depths over the realigned marsh, the amount of hours per month that its surface is
386 submerged ($T_{sub,tot} = 347$ hrs) is larger than for the tidal flat ($T_{sub,tot} = 263$ hrs). Perhaps more
387 significantly, the maximum continuous period of submergence is also longer for the realigned marsh
388 ($T_{sub,max} = 23$ hrs) compared to that for the tidal flat ($T_{sub,max} = 7$ hrs) and for the salt marsh ($T_{sub,max} = 5$
389 hrs). Over the 5-year monitoring period, the restored salt marsh experienced continuous
390 submergence for more than 4 days during 9 consecutive months, and during December 2013 and
391 February 2014 the marsh was under water for two continuous periods of more than 11 days. On
392 average, the realigned marsh experiences longer maximum periods of exposure per month ($T_{exp,max} =$
393 37 hrs) than the tidal flat ($T_{exp,max} = 10$ hrs), but shorter than the salt marsh ($T_{exp,max} = 98$ hrs). However,
394 over the 5-year monitoring period, the realigned marsh was continuously exposed for more than a
395 week during 14 months; the salt marsh was never exposed for that long a period. Finally, the rate of
396 the falling tide over the realigned marsh ($dh/dt = 1$ mm/min) is considerably slower than for the tidal
397 flat and the salt marsh ($dh/dt = 6$ mm/min); this is expected to lead to lower flow rates during the
398 ebbing tide over the restored salt marsh, reducing the potential for tidal creek development.

399

400 Figure 5 here

401

402 The settings of the SRT were modified several times to try and optimise the tidal flooding of the
403 realigned marsh and the SRT also frequently experienced malfunctioning. As a result, the realigned
404 marsh experienced significant temporal variability in the tidal inundation characteristics and several
405 distinct phases can be identified from the water-level time series recorded over the realigned marsh

406 (Figure 6). The two most extreme situations occurred during phase B (very limited flooding) and phase
407 D (very extensive flooding), and the inundation characteristics that resembled most closely that of the
408 natural salt marsh (blue dashed line in Figure 6) occurred during phase G, at least for N_{otides} , $T_{sub,tot}$ and
409 $T_{exp,max}$ (42 vs 45 tides, 215 vs 150 hrs and 130 vs 95 hrs, respectively). For most of the time, the tidal
410 inundation and exposure on the restored salt marsh most resembled that of the tidal flat (red dashed
411 line in Figure 6).

412

413 Figure 6 here

414

415 For each year, the water-level time series recorded on the realigned marsh for the period 1 Jan – 1
416 July was used to compute monthly tidal inundation statistics across the marsh surface. This 6-month
417 period was selected, rather than the full year, because it was felt that this period (end of winter, spring
418 and start of summer) was most relevant for the ecology. Figure 7 shows the spatial distribution in the
419 number of over-tides (tides that flood the realigned marsh surface) for the years 2012, 2015 and 2016
420 (years 2013 and 2014 are similar to 2015).

421

422 Figure 7 here

423

424 The tidal flooding of the realigned marsh can be compared with that on the natural salt marsh, and
425 the grey (red) area in Figure 7 represents a flooding frequency of 50–100% of that occurring on the
426 natural salt marsh. The figure strongly suggest that for most of the marsh, the tidal inundation during
427 2012 was significantly less than that for the natural salt marsh and that only the area adjoining the
428 central tidal creek and the location of where the WWII bomb struck ($x = -50$ m; $y = 400$ m) was
429 characterised by a tidal flooding regime similar to that of the natural salt marsh. This was also
430 highlighted in Figure 6, showing prolonged periods of continued exposure during 2012. In contrast,
431 the tidal flooding during 2015 (and 2013 and 2014) was excessive compared to that of the natural salt
432 marsh with almost the complete marsh up to $y = 700$ m flooded more frequently than the natural salt
433 marsh. This was also evident from Figure 6, showing prolonged periods of continued inundation during
434 the period 2013–2015. Only in the most recent year 2016 does the flooding regime on the realigned
435 marsh resemble that of the natural salt marsh. A larger part of the realigned marsh is flooded regularly
436 (up to $y = 700$ m), and tidal flooding is only excessive for the region around where the WWII bomb fell.

437

438 3.2 Tidal currents and sedimentation

439

440 Hydrodynamic measurements demonstrated that the maximum water depth across the adjacent
441 natural salt marsh during spring tides is 1–1.2 m, and that it does not flood during neap tides. Current
442 flows recorded in a small tidal channel in the salt marsh are generally weak; peaking at 0.2 m/s during
443 flood and less than 0.1 m/s during ebb. Suspended sediment concentrations in the water are low, with
444 maximum concentrations of 0.015–0.030 kg/m³ (or g/l) at the start of the flooding tide, reducing to c.
445 0.010 kg/m³ for the remainder of the inundation period. The flow characteristics for the realigned
446 marsh were monitored just inward from the SRT. Here, typical flooding velocities are 1.5–2 m/s. These
447 very localised strong flows only occur for a brief period of time at the start of the flooding tide and are
448 followed by weaker ebb velocities (< 0.5 m/s). Suspended sediment data collected near the inflow
449 pipe was circumspect due to the large amounts of organic material that often enters the realigned
450 marsh during the flooding tide. It is assumed that the sediment concentrations of the water entering
451 the realigned marsh through the pipe are the same as the water over the natural salt marsh.

452

453 Using measurements of the suspended sediment concentration and estimates of the tidal prism and
454 inundation area, the potential sedimentation rates in the restored salt marsh can be estimated using
455 the following equation:

$$456 \quad \Delta z = \frac{1}{P\rho} \frac{NAC}{S}$$

457 where Δz is sediment accretion rate per year (m/yr), N = number of over-tides per year (/yr), A = tidal
458 prism (m³), C = average suspended sediment concentration (kg/m³), P = sediment porosity (-), ρ =
459 sediment density (kg/m³) and S = inundation area (m²). The right term of the equation represents the
460 amount of sediment deposition in kg/yr, and the left term converts this to m/y. This equation assumes
461 that all sediment that enters the marsh during the flooding tide will be deposited with no sediment
462 exiting during the ebbing tide. Considering an average high tide level in the realigned marsh of 1 m,
463 which leads to $A = 6,000 \text{ m}^3$ and $S = 50,000 \text{ m}^2$ (Section 3.1), an average sediment concentration C
464 during the flooding tide of 0.02 kg m⁻³, a sediment density ρ of 2,650 kg m⁻³, a porosity P of 0.6 and
465 350 tidal cycles per year results in a vertical accretion rate of 0.0005 m yr⁻¹, or 0.5 mm yr⁻¹. This is at
466 least one order of magnitude less than what can be expected on a natural salt marsh (cf. Cundy et al.,
467 2007).

468

469 Some indication of sedimentation rates can be estimated from the sedimentation maps deployed in
470 the realigned marsh within 50 m from the SRT and also on the natural salt marsh. Typical values for
471 sedimentation rates on the natural salt marsh and the realigned marsh are 0.010–0.025 and 0.001–
472 0.01 kg/m²/day, respectively, with typical organic fraction of this material of 0.2 and 0.4 respectively

473 (White, 2014). These values can be used to estimate vertical accretion rates using the following
474 equation:

$$475 \quad \Delta z = \frac{NQ(1 - O)}{P\rho}$$

476 where Δz is sediment accretion rate per year (m/yr), N = number of days per year (/yr), Q is the
477 sedimentation rate (kg/m²/day), O is organic fraction (-), P = porosity (-) and ρ = sediment density
478 (kg/m³). Application of this equation yields values of Δz for the natural salt marsh of 0.0018 – 0.0046
479 m/y, or c. 3 mm/y. For the realigned marsh, Δz = 0.0003 – 0.0014 m/y, or c. 0.5 mm/y. Perhaps
480 fortuitously, this value is identical to that derived from the suspended sediment computations.

481

482 **3.3 Morphological change**

483

484 Any morphological change was limited to the immediate area around the SRT. The influx of sediment
485 was insufficient to induce vertical accretion of the marsh surface and the flow velocities across the
486 marsh surface were too weak to establish new creeks or modify the existing rather rectilinear drainage
487 system. The planform changes near the SRT recorded using the terrestrial laser scanner data are
488 shown in Figure 8. The main change is the development of a bend in the central drainage channel and
489 the deposition of a mid-channel ‘bar’ in the vicinity of the SRT. This seems to have been a steady
490 process over the 5-year monitoring period.

491

492 Figure 8 here

493

494 The development of the bend is further illustrated by the evolution of transect $y = 25$ m (Figure 9),
495 which clearly shows a widening of the drainage channel through 5–6 m erosion of the westward
496 channel bank (from 2011 to 2016), whilst the eastern bank remained stable. This change was very
497 localized, as no significant widening of the channel occurred at any of the other transects (not shown).
498 This erosion of the western bank is considered a direct consequence of the alignment of the inlet /
499 outfall pipe, which directs high velocity flows onto the western bank. Widening of the channel near
500 the SRT is a response to the volume of water flooding and draining the marsh (i.e., the tidal prism) and
501 is accomplished by the high current velocities through the intake/outfall pipe. It appears that the
502 sediment eroded from the bank and the channel is re-deposited on the bank (in the form of a levee)
503 and in the channel (in the form of a mid-channel bar), and is unlikely to contribute significantly to salt
504 marsh accretion.

505

506 Figure 9 here

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3.4 Ecological variability

3.4.1 Vegetation

All full list of plant species recorded within the flooded area of the site (pre- and post-breach) is given in Table 4. Species richness was greatest just before breach in June 2011 (37 species) and the site was dominated by plants typical of (wetland) pasture communities. In addition, six common upper salt marsh species (*Atriplex patula*, *Glaux maritima*, *Juncus gerardii*, *Puccinellia maritima*, *Spergularia maritima* and *Triglochin maritima*) were present and these were associated with saline seepages under the embankment. Although 30 species were recorded a year later, in subsequent years, species richness more than halved compared to pre-breach values and only 13 species remained in 2016; however, these did include two salt marsh species (*Aster tripodium* and *Salicornia europaea*) that were not present in 2011.

Table 4 here

At the community (NVC) level, the vegetation changes in the realigned marsh from before SRT installation in June 2011 to June 2016 can be subdivided into 4 stages (Figure 10):

- **Pre-breach:** Vegetation was dominated by *Agrostis stolonifera* - *Alopecurus geniculatus* (NVC MG13) grassland, with small patches of *Juncetum gerardii* SM16 salt marsh near saline seeps. On the adjacent natural salt marsh (-60-m transect), the vegetation was, and remained throughout the study, dominated by *Spartina anglica* SM6 salt marsh.
- **2011-2013:** Modest changes occurred in the realigned marsh over the 2-year period following the breach (Figure 10 – left panel), with most MG13 quadrats unaltered, but a small number replaced by bare ground. The SM16 community extended to 150 m.
- **2013-2015:** By 2015 most vegetation had senesced to be replaced by bare ground (32 of the 40 quadrats surveyed) (Figure 10 – middle panel). The decline was progressive, with bare ground increasing from 13% in 2012, 48% in 2013, 70% in 2014 to 80% in 2015.
- **2015-2016:** Over the 12-months from June 2015 to June 2016, vegetation on the realigned marsh recovered and, whilst the five remaining MG13 quadrats above the flood line remained unaltered, 21 of the 32 bare ground quadrats transitioned to the ephemeral *Spergularia marina*-*Puccinellia distans* SM23 Community Type (Figure 10 – right panel).

541 Figure 10 here

542

543 3.4.2. Seed deposition

544

545 Of the 20 different species (from 1390 individual seedlings) recorded on mats recovered from the
546 restored site, only 8 were salt marsh species and these accounted for only 18% of all seedlings. Of
547 these, *Triglochin maritima* and *Juncus gerardii* (both 6% of all seedlings) and *Spergularia marina* (4%)
548 were the most frequently recorded species. There was no increase in seed deposition from time of
549 SRT installation and the relative proportion of salt marsh species did not vary through time (White,
550 2014). From the mats recovered from the natural salt marsh, 12 different species germinated (from
551 635 seedlings) and over 97% (9 species) of these were typical salt marsh plant species with *Aster*
552 *tripolium* dominating the seedling pool (86% of the seedlings).

553

554 3.4.3 Foraminifera

555

556 Foraminifera first colonised the realigned marsh in May 2013, principally at the station closest to the
557 SRT ($y = 25$ m). At this location, a typical low salt marsh community, dominated by the agglutinated
558 taxa *Milammina fusca* (43%), *Jadammina macrescens* (21%) and *Trochammina inflata* (11%), became
559 established between May and November 2013, and peaked in the August 2013 samples with a count
560 of 15.5 tests per cc (Figure 11). This is, however, a very low concentration in comparison to fully
561 established salt marsh communities. As well as a declining foraminiferal stock at this location after the
562 peak in August 2013, there also appears to be a shift to an assemblage more typical of a mudflat
563 community. This is reflected by a decline or absence of the high salt marsh taxa (*Jadammina*
564 *macrescens* and *Trochammina inflata*), to a typical low salt marsh community dominated by
565 *Milammina fusca* with a notable presence of calcareous taxa *Elphidium sp.* and *Haynesina germanica*
566 which are typical of mudflat communities. By the final survey in July 2016, when abundance reached
567 its maximum, it appears this low salt marsh to mudflat community was becoming well-established as
568 indicated by the dominance of the low salt marsh species *Milammina fusca* (35%), the mudflat taxa
569 *Elphidium spp.* (35%) and *Haynesina germanica* (15%), and the decline of the high salt marsh taxa
570 *Jadammina macrescens* (5%) and *Trochammina inflata* (4%). However, the maximum foraminifera
571 concentration of 27.5 tests per cc indicates communities at this site remain immature.

572

573 At $y = 100$ m there are two sampling locations at either side of the central creek. West of the central
574 creek there is evidence of a typical low salt marsh foraminifera community becoming established with

575 concentrations dominated by *Miliammina fusca* present in all the sampling periods from 2013 to 2016.
576 The greatest concentrations of foraminifera in the marsh occurs at this location in the November 2015
577 (68.5 tests per cc) and July 2016 (121 tests per cc) samples, indicating that the *Miliammina fusca*
578 dominated low salt marsh community is becoming well established at this location. Additionally, in
579 the November 2015 and July 2016 samples, significant proportions high salt marsh taxa *Jadammina*
580 *macrescens* (up to 8 %) and *Trocammina inflata* (up to 25%) comprised the assemblages, indicating a
581 subtle transition from low to high salt marsh environments. East of the central creek there was little
582 indication of a foraminiferal community becoming established with few or no foraminifera recorded
583 until November 2015, when concentrations reached a maximum of 4.75 tests per cc. By July 2016
584 there is evidence of a middle-to-high salt marsh community becoming established with a moderate
585 concentration of foraminifera (52 tests per cc) dominated by *Miliammina fusca* (50%), *Trochammina*
586 *inflata* (29%) and *Jadammina macrescens* (19%).

587

588 Figure 11 here

589

590 3.5 Ecological transition and variation in the tidal environment

591

592 The spatial distribution in the inundation statistics across the realigned marsh, such as shown in Figure
593 7, was applied to all the ecological monitoring locations over the complete 5-year monitoring period.
594 Subsequently, the inundation statistics were pooled for each of the NVC Community Type, and the
595 bare ground category ('bare'), and the results presented in a boxplot (Figure 12). All values are
596 monthly statistics, but these were computed using 6 months of water-level data over the realigned
597 marsh (January–June).

598

599 Figure 12 here

600

601 The conditions characterising the 'Bare' vegetation type clearly stands out when compared with the
602 other NVC Community Types; it is associated with the largest number of over-tides ($N_{otides} = 52$), the
603 largest number of total hours of inundation ($T_{sub,tot} = 338$ hrs), the longest continuous period of
604 inundation ($T_{sub,max} = 8$ hrs; also the greatest variability) and also the longest period of exposure ($T_{exp,max}$
605 $= 45$ hrs). This vegetation type was particularly ubiquitous over the period 2013–2015. Compared with
606 the other NVC Community Types in the area, the flooding characteristics for 'Bare' are closest to that
607 of the natural *Spartina* (SM6) salt marsh in terms of number of over-tides ($N_{otides} = 48$, but most similar
608 to the *Juncetum gerardii* (SM16) salt marsh in terms of total hours of inundation ($T_{sub,tot} = 328$ hrs).

609 However, the flooding regime of the 'bare' category is more extreme than that of SM6 and SM16, and
610 is overall more akin that of the tidal flat (refer to Section 3.1).

611

612 The MG13 Community Type (*Agrostis stolonifera* - *Alopecurus geniculatus* grassland) is associated with
613 the least amount of tidal inundation: $N_{otides} = 10$, $T_{sub,tot} = 33$ hrs, $T_{sub,max} = 3$ hrs and $T_{exp,max} = 171$ hrs.
614 MG13 is the original vegetation type, and it seems it can tolerate a significant tidal flooding, as several
615 MG13 plots received more than 20 over-tides and in excess of 100 hrs tidal inundation per month. The
616 SM23 ephemeral *Spergularia marina*-*Puccinellia distans* salt marsh was the dominant vegetation in
617 the inundated area in 2016. It has tidal inundation characteristics that significantly overlap with that
618 of MG13; however, focussing on the median values, the tidal inundation for SM23 is more frequent
619 and of longer duration than for MG13: $N_{otides} = 12$ (*versus* 10), $T_{sub,tot} = 49$ hrs (*versus* 39 hrs), $T_{sub,max} =$
620 4 hrs (*versus* 3 hrs).

621

622 In summary, there is a clear gradient in terms of increasing tidal inundation intensity from MG13 →
623 SM23 → SM6 → SM16 → Bare. There is considerable overlap between the different NVC
624 Communities, but it must be borne in mind that vegetation in the realigned marsh is in transition;
625 therefore, the actual inundation statistics do not necessarily reflect optimal conditions.

626

627 4. Discussion

628

629 The status of the realigned marsh in June 2016 is illustrated by the aerial photograph shown in Figure
630 13. It shows that the majority of the marsh that is regularly inundated by the tide ($z < 1.2$ m ODN) is
631 characterised by the SM23 NVC Community Type. There is a distinct boundary between SM23 and
632 MG13, both in terms of colour Figure 13 (brown-green versus green) and elevation, as the $z = 1.2$ m
633 ODN contour line separates these two NVC types very well. Five years post-breach, the realigned
634 marsh at South Efford is now transitioning into salt marsh. However, while the foraminifera
635 assemblage is typical of mid-high-level salt marsh (e.g., Gehrels et al. 2001), the *Spergularia* dominated
636 (NVC SM23) vegetation is at best only an ephemeral salt marsh community.

637

638 Figure 13 here

639

640 A lack of propagule supply, at least for the first three years after breach, may go some way to
641 explaining why there was no transition to a more typical *Spartina*-dominated (i.e. NVC SM6) perennial
642 plant community. Although we cannot be sure that post-2014 propagules were not entering from the

643 nearby natural marsh, it seems likely that a consequence of the SRT system used at South Efford was
644 that immigration of seeds and root fragments was limited. Indeed, aside from the very small number
645 of (wind dispersed) *Aster tripolium* seeds that germinated on the sediment mats, *Triglochin maritima*,
646 *Juncus gerardii*, and *Spergularia marina* were present inside the embanked area prior to breach. Even
647 with plentiful propagule supply however, the tidal regime at the site would likely have restricted plant
648 establishment.

649

650 The limited inundation during the early post-breach phase (2011–2013) was due to the initial setting
651 of the SRT; this would have further restricted the immigration of salt marsh plant propagules and the
652 development of physical-chemical conditions suitable for their establishment. The excessive
653 inundation during 2013–2015 can be attributed to malfunctioning of the SRT when it was blocked with
654 detritus, and extremely wet winters of 2013/14 and 2014/15. During this period the realigned marsh
655 was submerged continuously for several weeks, conditions which probably contributed to the
656 development of the bare surface typical of many MR schemes (Mossman 2012a). Drainage was further
657 hampered by the SRT system which compelled all water to drain through a single circular 0.9-m
658 diameter outflow pipe at an elevation little different to low tide in the adjacent estuary. Only during
659 2016 did the inundation characteristics on the realigned marsh resemble that of the adjacent natural
660 salt marsh in terms of number of over-tides and duration of tidal inundation. At this time, over most
661 of the realigned marsh, the number of over-tides was between 24 and 48, compared to 48 over-tides
662 on the natural salt marsh and less than 38 and 42 over-tides per month as recommended by Ash and
663 Fenn (1997) and Environment Agency (2003), respectively. Only during 2016 did any typical salt marsh
664 vegetation develop widely across the site and, even then, this was restricted to a plant community
665 dominated by the ephemeral annual *Spergularia marina*. Foraminifera populations remained low (<
666 130 individuals per cc) compared with natural marshes along the Avon Estuary which generally exceed
667 1000 individuals per cc (Stubbles, 1999). This indicates that foraminifera communities in the realigned
668 marsh are immature, but their increasing diversity and abundance over the 5-year monitoring period
669 suggests salt marsh species are beginning to thrive.

670

671 Given that elevation within the tidal frame is the most important factor determining the success of
672 propagule and salt marsh establishment (Adam, 1990; French, 2006; Davy et al., 2011), our
673 observations at South Efford underscore the challenge in achieving the flooding and drainage regime
674 experienced by natural salt marsh, especially when South Efford is, like many managed realignment
675 sites, of much lower elevation than the adjacent natural marsh (Wolters et al., 2005, 2008; Spencer
676 and Harvey, 2012). In addition to a direct impact on propagule delivery and establishment, prolonged

677 flooding reduces sediment redox potential, while a small tidal prism (depth over the marsh < 0.2 m)
678 limits further the amount of suspended sediment entering the marsh. Together these factors mitigate
679 against the establishment of salt marsh vegetation on managed realignment sites ([Mossman et al.,](#)
680 [2012a, b](#); [Spencer and Harvey, 2012](#)). Moreover, and unlike natural topographically complex marshes
681 dissected by numerous tidal creeks, South Efford (like many managed realignment sites) is a horizontal
682 surface with limited drainage channels (only one central channel). The only significant morphological
683 change occurred in the vicinity of SRT due to large water flow velocities entering the marsh through
684 the pipe, and there was very limited evidence of an emergent creek network; the flow velocity over
685 the marsh was too low to entrain sediment and create a dendritic creek network that drains to the
686 main drain. More generally, even if all the sediment entering the realigned marsh is deposited on the
687 marsh surface, the vertical accretion rates are estimated to be one order of magnitude less than on
688 the adjacent natural salt marsh. This means that with rising sea level, the realigned marsh will
689 increasingly lag behind the natural salt marsh in terms of its elevation. This will make it increasingly
690 difficult to maintain a 'natural' inundation regime over the restored salt marsh and ultimately the salt
691 marsh restoration effort will fail. Limited sediment influx is probably a characteristic of intertidal
692 habitat schemes involving regulated tidal exchange (including SRT); in fact, [Pontee \(2014\)](#)
693 recommends installation of a RTE to limit siltation rates in intertidal habitat restoration sites, as
694 opposed to breaching.

695

696 **Concluding remarks**

697

698 A regulated tidal exchange (RTE) option was implemented at South Efford to enable intertidal habitat
699 creation without increasing flood risk to neighbouring properties, and a self-regulating tide gate (SRT)
700 selected to control the water levels over the realigned marsh. One of the main advantages of RTE is
701 the ability to regulate the tidal water levels; nonetheless, our results suggest that through installation
702 of a SRT, it has not been possible to consistently facilitate a natural tidal inundation regime at South
703 Efford. Apart from the non-trivial issue of optimising the gate settings, frequent malfunctions due to
704 jamming by detritus let in too much, or too little water. Even when the SRT operated as planned,
705 inundation was always quicker than drainage, causing extended periods of submergence that
706 prevented the establishment of vegetation. Slow drainage further impeded the development of tidal
707 creek systems which would enable sediment recycling and accretion fed by creek expansion. Perhaps
708 most significantly, the SRT allowed only small amounts of suspended sediment into the site and
709 sedimentation was insignificant; consequently, a major pre-requisite for natural salt marsh
710 development – vertical accretion – was missing. We conclude therefore that the SRT system used at

711 South Efford was unable to impose the natural physical parameters required for salt marsh
712 development, and more generally question the ability of SRT to achieve a sustainable and naturally
713 functioning salt marsh at any managed realignment site. It may be possible to achieve a tidal flooding
714 regime conducive to the development of suitable intertidal habitat (mudflat or salt marsh), but
715 facilitating at the same time sustainable vertical accretion rates (similar to the rate of sea-level rise)
716 and the development of a tidal creek network might be over-ambitious. We conclude that SRT can be
717 a useful technique for intertidal habitat creation where there are significant site constraints (especially
718 flood risk), but we need to be realistic in our expectations of what it can achieve in terms of delivering
719 a perennial salt marsh community.

720

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722

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728

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843

844 **Table 1** – Overview of 5-year monitoring programme at South Efford.

Parameter	Sampling location	Sampling frequency	Sampling period
Water level	In restored salt marsh and in the river	Every 15 min	2011–2016
Salinity	In restored salt marsh	Every 15 min	2011–2016
Tidal currents and suspended sediment concentrations	In restored salt marsh and in natural salt marsh	4 Hz	3 March – 4 May 2012
Suspended sediment concentrations	In natural salt marsh	5-10 min	21 March 2012
Sedimentation and seedling mats	Many locations (> 10) in restored salt marsh	Mats were deployed over several months at a time	2011–2014
Total station surveys	10 transects in restored marsh and 1 transect in natural salt marsh	yearly	2011–2016
Laser scanner surveys	Scan from top of SRT of restored marsh and natural salt marsh	yearly	2011–2016
Unmanned Aerial Survey	Survey of restored marsh and natural salt marsh	once	27 June 2016
Vegetation survey	40 quadrats in restored marsh and 4 quadrats in natural salt marsh	Half-yearly to yearly	2011–2016
Foraminifera sampling	3 samples from restored marsh near SRT site	Three-monthly to yearly	2011–2016

845

846 **Table 2** – NVC Community Types observed during the 5-year monitoring period in South Efford
 847 marsh.

Code	NVC Community Type
SM6	<i>Spartina anglica</i> salt marsh - distinctive salt marsh community on the seaward fringes of marshes and on creek sides,
SM16	<i>Juncetum gerardii</i> salt marsh - characteristic of mid-upper coastal marshes
SM23	<i>Spergularia marina-Puccinellia distans</i> salt marsh - characteristic of disturbed situations with soils of variable but generally high salinity (e.g. upper pans) on coastal marshes
MG13	<i>Agrostis stolonifera - Alopecurus geniculatus</i> Grassland - Typical of inundation; usually in river flood plains and on the edges of ponds
MG10	<i>Holcus lanatus - Juncus effusus</i> rush-pasture - Typically associated with poorly drained permanent pastures
Bare	No vegetation

848

849 **Table 3** – Marsh sub-environment classifications based on typical foraminifera assemblages recorded
850 in southwest England (references in text).

	Foraminifera Community
High salt marsh	<i>Jadammina macrescens</i> , <i>Trochammina inflata</i> , <i>Haplophragmoides wilberti</i>
Low salt marsh	<i>Miliammina fusca</i>
Mudflat	<i>Cibicides lobatulus</i> , <i>Elphidium spp.</i> , <i>Haynesina germanica</i> , <i>Quinqueloculina spp.</i>

851

852

853 **Table 4** – Pre- (June 2011) and post-breach changes in plant species composition recorded in fixed
854 4x4m quadrats positioned along a tidal inundation gradient at the South Efford managed
855 realignment site, Devon, SW England. Species presence in at least one quadrat is denoted by 'P'
856 and species typical of British salt marsh vegetation are highlighted in bold type.

857

Species		Year					
		2011	2012	2013	2014	2015	2016
Grasses	<i>Agrostis stolonifera</i>	P	P	P	P	P	P
	<i>Alopecurus geniculatus</i>	P	P	P	P	P	P
	<i>Anthoxanthum odoratum</i>	P					
	<i>Arrenatherum eliatum</i>	P					
	<i>Cyanosurus cristatus</i>	P	P	P	P		
	<i>Elymus repens</i>	P					
	<i>Festuca pratensis</i>	P	P				
	<i>Holcus lanatus</i>	P	P		P		
	<i>Lolium perenne</i>	P	P				
	<i>Poa annua</i>	P					
	<i>Poa pratensis</i>	P					
	<i>Poa trivialis</i>	P					
	<i>Puccinellia maritima</i>	P	P	P	P	P	P
Sedges/Rushes	<i>Bolboschoenus maritimus</i>		P	P	P		
	<i>Carex otrubae</i>	P	P	P			P
	<i>Carex ovalis</i>	P					
	<i>Eleocharis palustris</i>	P	P	P			
	<i>Juncus articulatus</i>	P	P	P	P		P
	<i>Juncus bufonius</i>		P				P
	<i>Juncus effusus</i>	P	P	P	P	P	
		<i>Juncus gerardii</i>	P	P	P	P	P
Forbs	<i>Aster tripolium</i>		P		P	P	P
	<i>Atriplex patula</i>	P	P		P	P	P
	<i>Cardamine pratensis</i>	P	P				
	<i>Cerastium holosteoides</i>	P	P				
	<i>Glaux maritima</i>	P	P	P			
	<i>Leontodon autumnalis</i>	P					
	<i>Leontodon hispidus</i>	P					P
	<i>Plantago major</i>	P	P				
	<i>Plantago media</i>		P				
	<i>Prunella vulgaris</i>		P				
	<i>Ranunculus acris</i>	P					

<i>Ranunculus sceleratus</i>		P		P		P
<i>Ranunculus repens</i>	P	P	P	P		
<i>Rumex acetosa</i>	P					
<i>Rumex conglomeratus</i>	P					
<i>Rumex crispus</i>	P					
<i>Rumex obtusifolius</i>		P	P			
<i>Salicornia europaea</i>		P		P	P	P
<i>Sonchus arvensis</i>				P		
<i>Spergularia marina</i>	P	P	P	P	P	P
<i>Taraxacum officinale</i>	P	P	P	P	P	
<i>Trifolium dubium</i>	P					
<i>Trifolium pratense</i>	P					
<i>Trifolium repens</i>	P	P	P	P		
<i>Triglochin maritima</i>	P	P	P	P		
Total species	37	30	17	19	10	13
Total salt marsh species	6	8	5	7	6	6

858

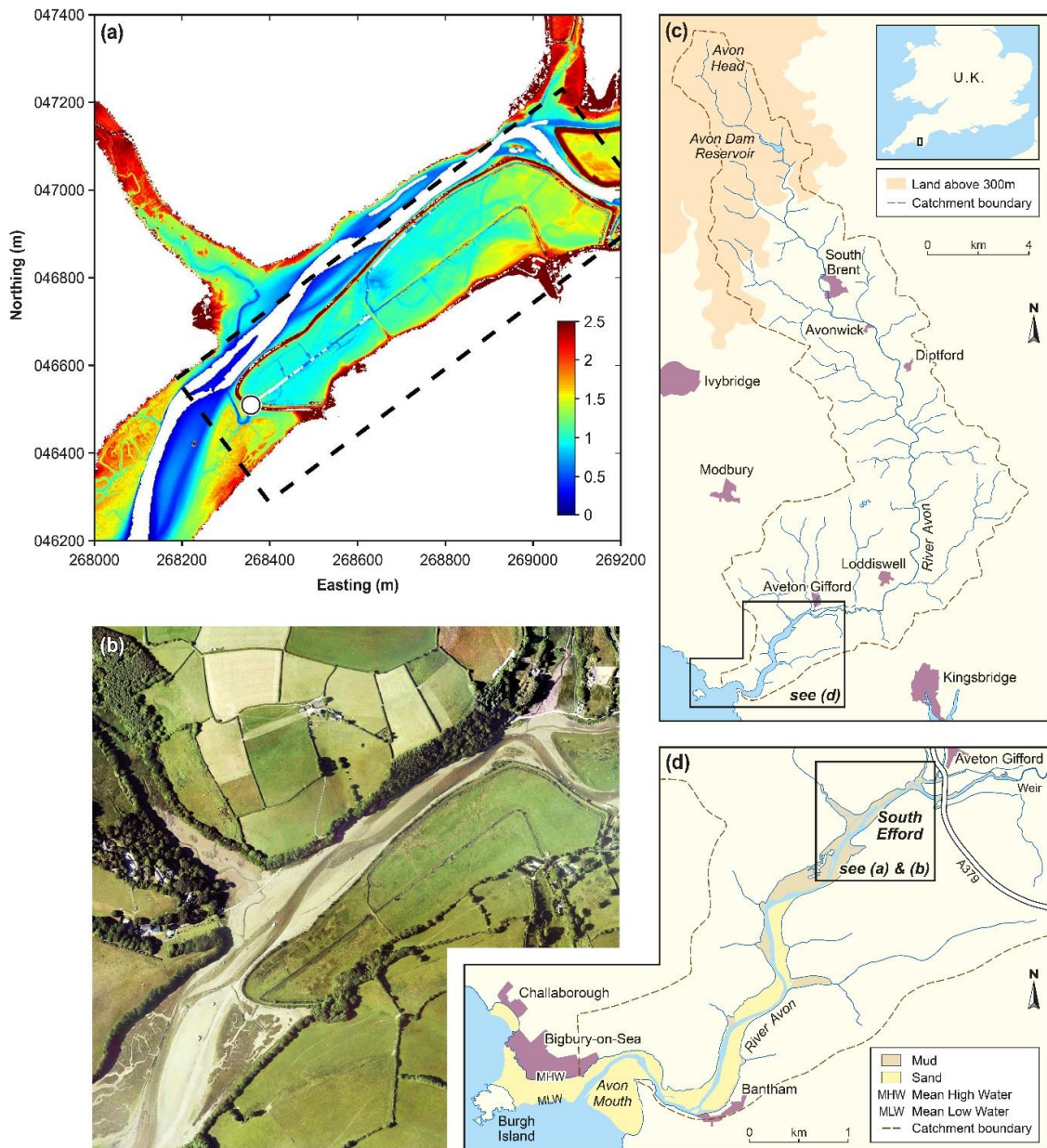


Figure 1 – Location map of the Avon estuary, and 2010 aerial photograph (photo from GoogleEarth) and digital elevation model (DEM) of South Efford marsh. The DEM is based on LIDAR data provided by the Plymouth Coastal Observatory (PCO), and elevations are in m ODN, which is approximately 0.2 m above mean sea level. The black dashed rectangle represents the realigned marsh area plotted in subsequent figures and is 400 m x 1000 m. The red circle represents the location of the self-regulating tidal gate (SRT) and the colour bar refers to the elevation in m ODN. White regions represent standing water (river section) or elevations > 5 m (valleys sides).

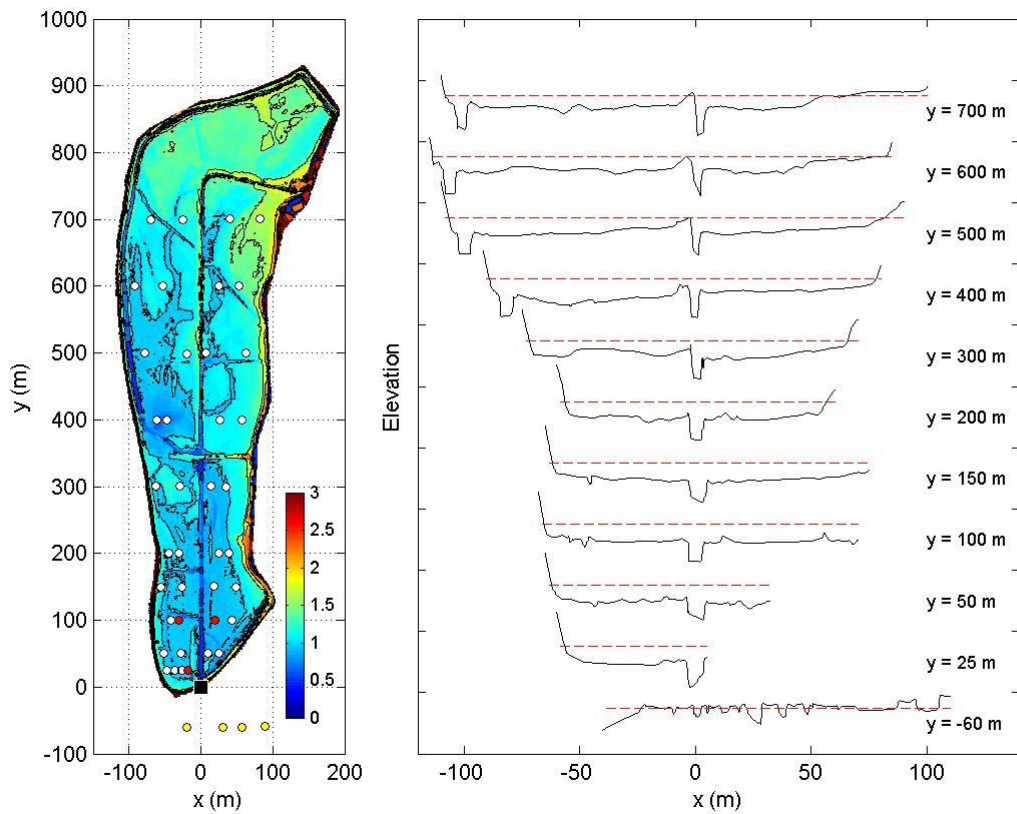


Figure 2 – Left panel shows digital elevation model of South Efford marsh on local coordinate grid with ecological sample locations (circles). The sample sites on the natural salt marsh are denoted by the yellow circles and the red circles represent the foraminifera sample locations. The black square represents the location of the SRT. Right panel shows total station surveys of all across-marsh transects measured during the baseline survey in May 2011. The profiles have been vertically offset by 2 m for ease of comparison and the tick marks on the y-axis represent 1-m intervals. The red dashed horizontal line represents the average level of the natural salt marsh. The contour lines represent 0, 1, 2, 3 and 4 m ODN.

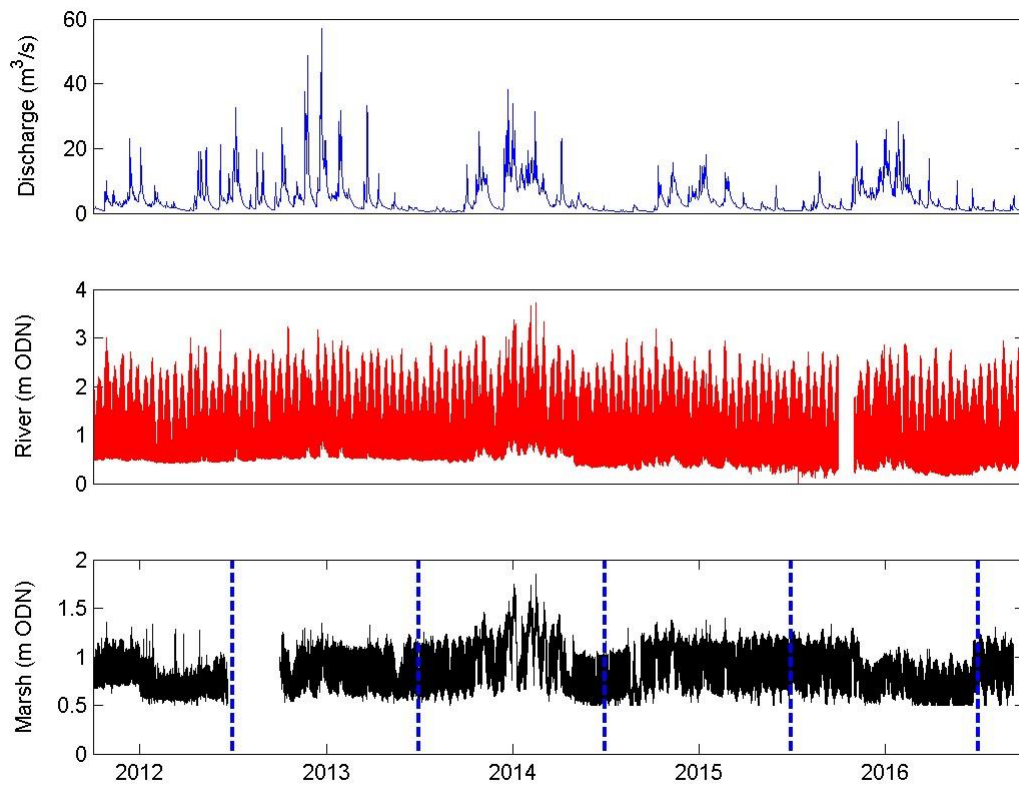


Figure 3 – Time series of river discharge recorded at Loddiswell (upper panel), water level in the river (middle panel) and the realigned marsh (lower panel). The vertical dashed lines in the lower panel represent 1 July, the date by which the annual ecological surveys were finished.

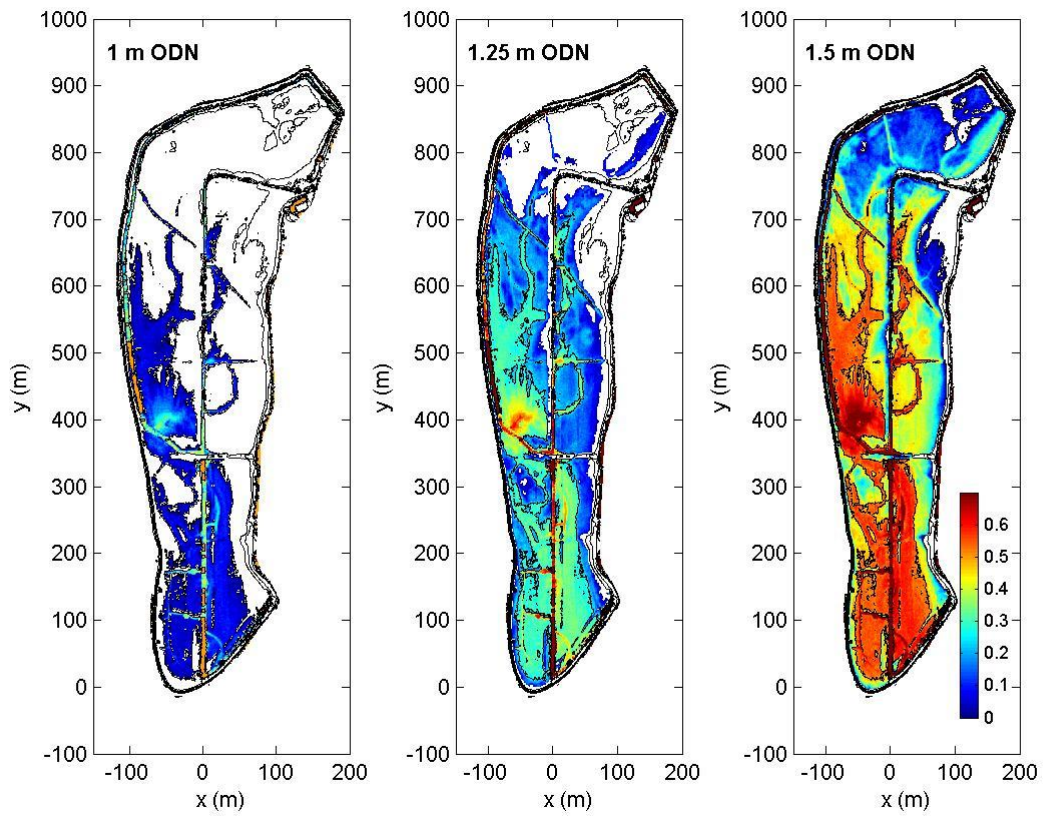


Figure 4 – Water depth on South Efford marsh for marsh water levels of 1, 1.25 and 1.5 m ODN. The colour bar in the right panel applies to all panels.

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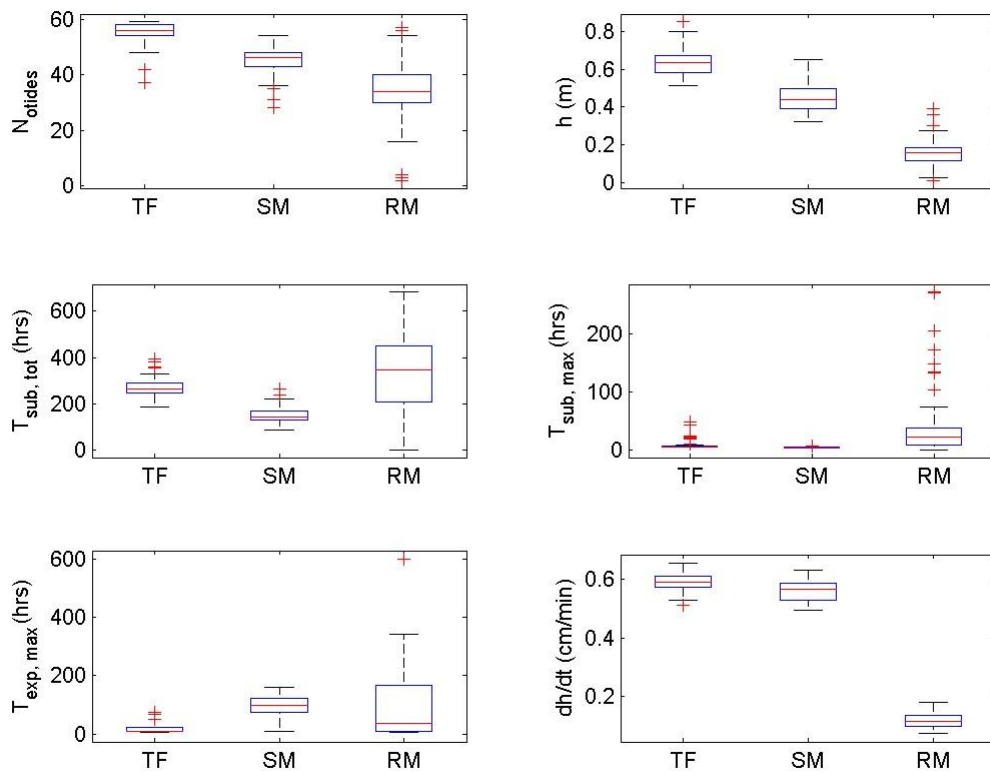


Figure 5 – Box plots of various monthly tidal parameters computed over the complete 5-year survey period for the natural tidal flat (TF; $z = 1$ m ODN), natural salt marsh (SM; $z = 1.5$ m ODN) and the realigned marsh (RM; $z = 0.9$ m ODN): N_{otides} = number of over tides; h = average water depth; $T_{sub,tot}$ = total hours of tidal submergence; $T_{sub,max}$ = maximum continuous period of tidal submergence; $T_{exp,max}$ = maximum continuous period of exposure; dh/dt = average rate of falling tide. On each box, the central mark is the median, the edges of the box are the 25 and 75 percentiles, the whiskers extend to the most extreme data points the algorithm considers to be not outliers (0.7 and 99.3 percentiles), and the outliers are plotted individually.

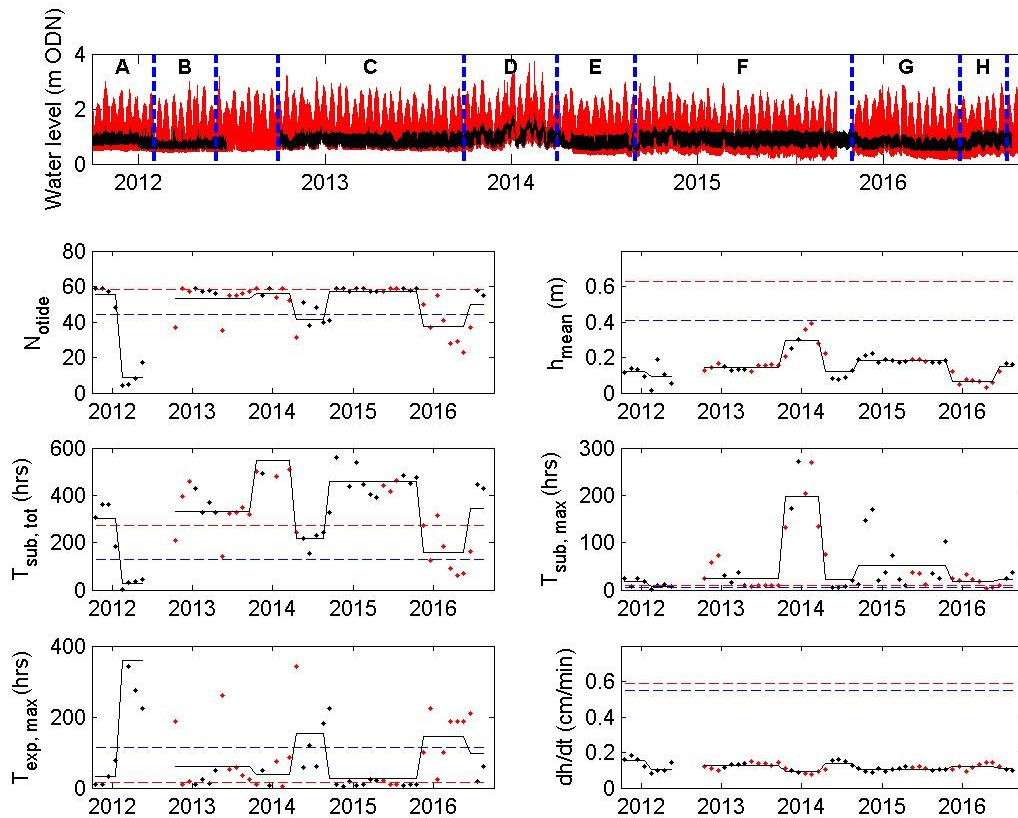


Figure 6 – Top panel shows time series of water level in the river (red line) and the realigned marsh (black line). Smaller panel show monthly time series of tidal parameters for the realigned marsh compared with the mean value over the 5-year period for the tidal flat (red dashed line) and the salt marsh (blue dashed line). The line in the lower panels represents the mean values for the phases indicated in the top panel and the red symbols represent months with significant malfunctioning of the SRT.

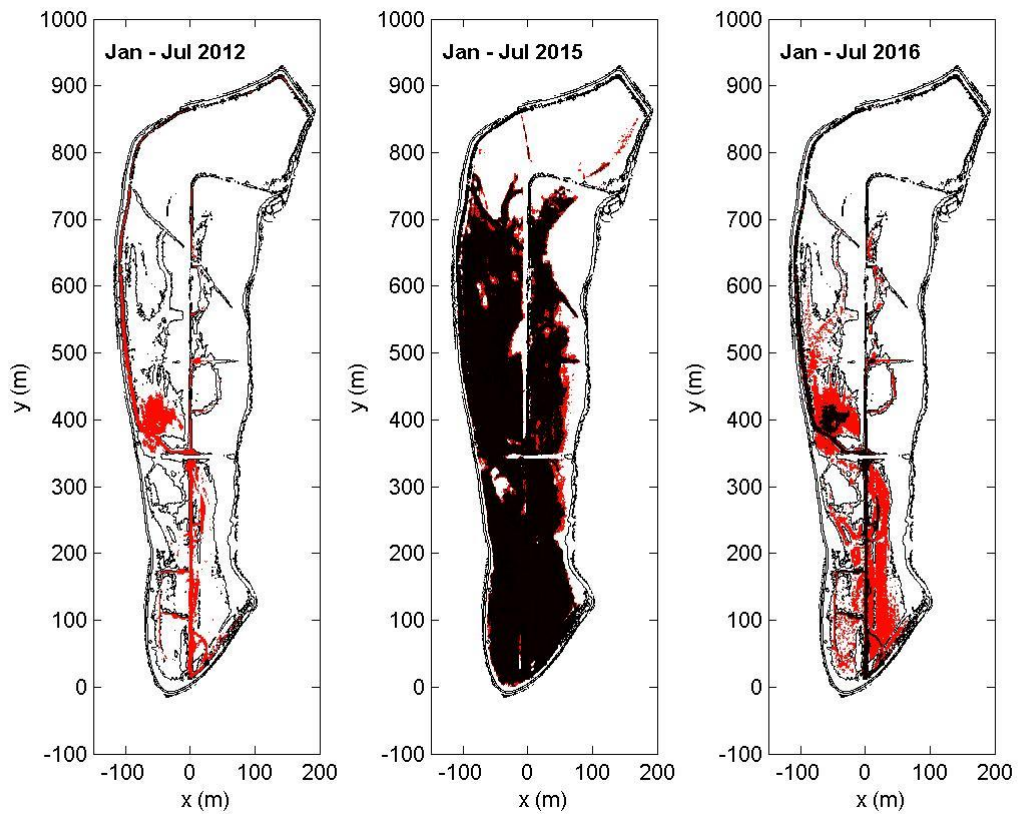


Figure 7 – Spatial distribution across the realigned marsh of the number of over-tides per month (tides that flood the marsh surface) for three years (2012, 2015 and 2016) and calculated for the 6-month period prior to the ecological survey (January to June). Optimal salt marsh conditions are considered to occur in the grey (red) area, representing 24–48 over-marsh tides per month. White and black areas experience less than 24 or more than 48 over-marsh tides per month, respectively. The natural marsh is, on average, flooded 48 times per month (cf. Figure 5). The contour lines represent 0, 1, 2, 3 and 4 m ODN.

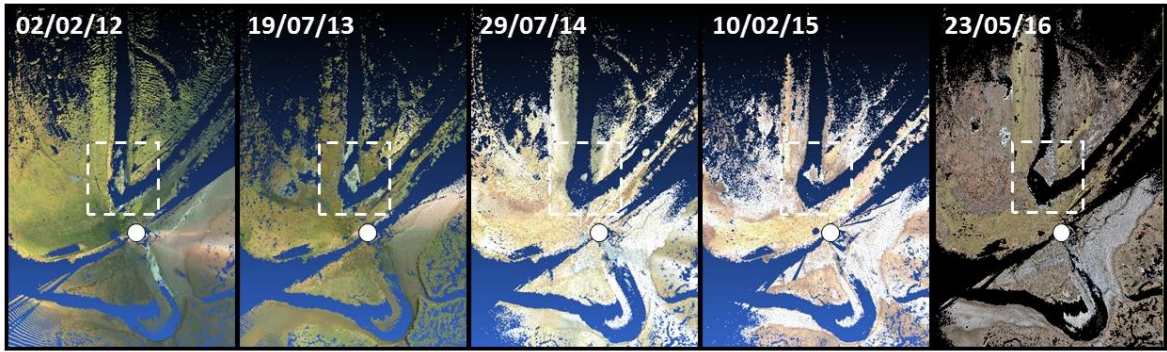


Figure 8 – Terrestrial laser scans obtained from the top of the tidal gate (denoted by white circle), with the realigned marsh at the top of the image and the natural salt marsh at the bottom. A very modest increase in the curvature of the channel in the realigned marsh can be observed and is being achieved through erosion of the left (west) bank of the channel nearest to the tidal gate (see also Figure 9).

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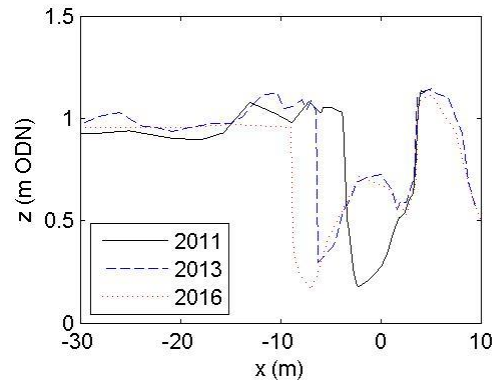


Figure 9 – Photograph taken from the top of the SRT looking towards the realigned marsh (taken in 2014) and morphological evolution of the transect running across the tidal creek near the SRT (at $y = 25$ m) showing progressive erosion of the western (left) bank.

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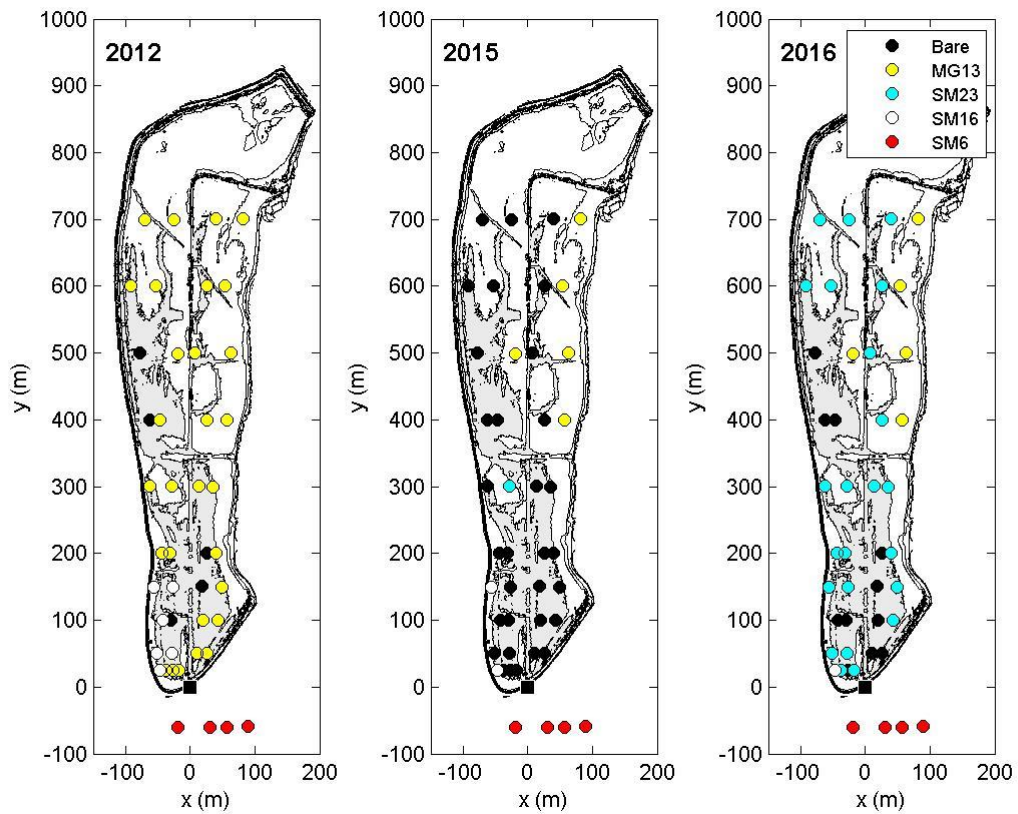


Figure 10 – Spatial distribution of the NVC Community Type at South Efford for the June surveys in 2012 (also representative of 2011), 2015 (also representative of 2013 and 2014) and 2016. The black square represents the SRT and grey shaded area represents below 1 ODN. The contour lines represent 0, 1, 2, 3 and 4 m ODN. The observed NVC Community Types are: MG13 = *Agrostis stolonifera* - *Alopecurus geniculatus* Grassland; SM23 = *Spergularia marina*-*Puccinellia distans* salt marsh; SM16 = *Juncetum gerardii* salt marsh; SM6 = *Spartina anglica* salt marsh.

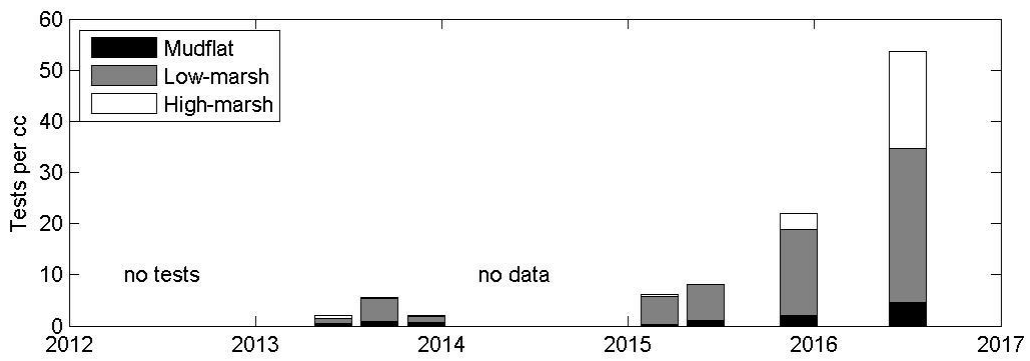


Figure 11 – Average foraminifera concentrations of mudflat, low-marsh and high-marsh biozones over the monitoring period. The concentrations were averaged for the 3 sample locations (1 at y = 25 m; 2 at y = 100 m). Biozone classifications are presented in Table 2.

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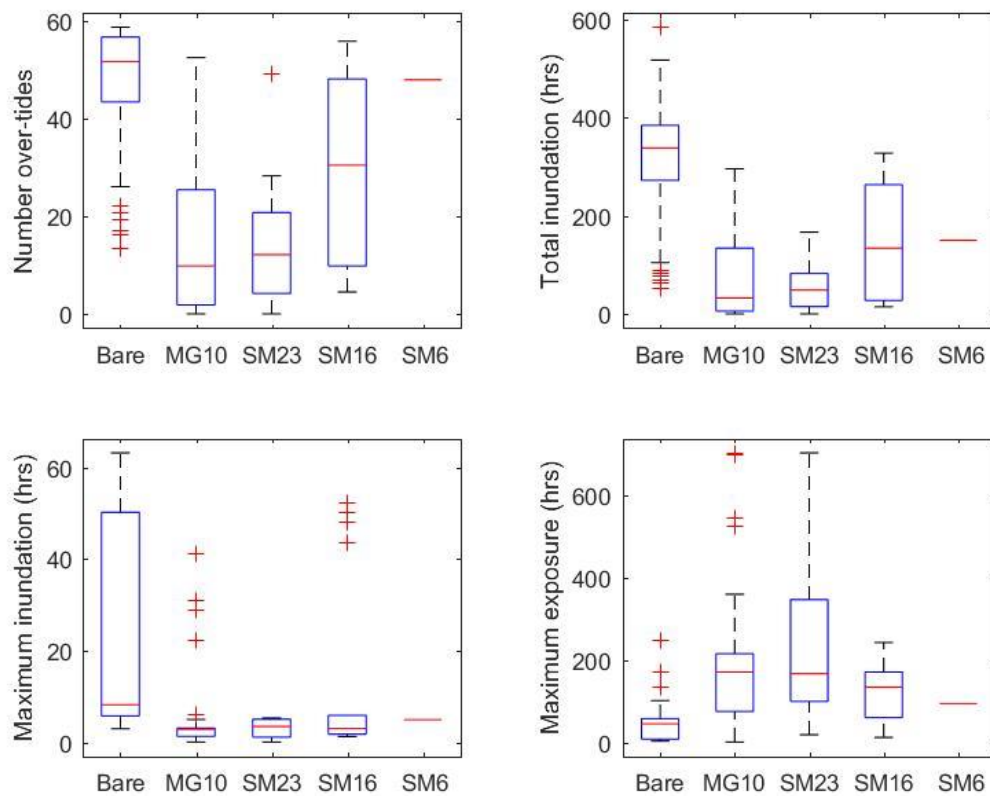


Figure 12 – Tidal inundation statistics for the different NVC Community Types at South Efford using all ecological monitoring locations over the 5-year monitoring period. The observed NVC Community Types are: MG13 = *Agrostis stolonifera* - *Alopecurus geniculatus* Grassland; SM23 = *Spergularia marina*-*Puccinellia distans* salt marsh; SM16 = *Juncetum gerardii* salt marsh; SM6 = *Spartina anglica* salt marsh. On each box, the central mark is the median, the edges of the box are the 25 and 75 percentiles, the whiskers extend to the most extreme data points the algorithm considers to be not outliers (0.7 and 99.3 percentiles), and the outliers are plotted individually.

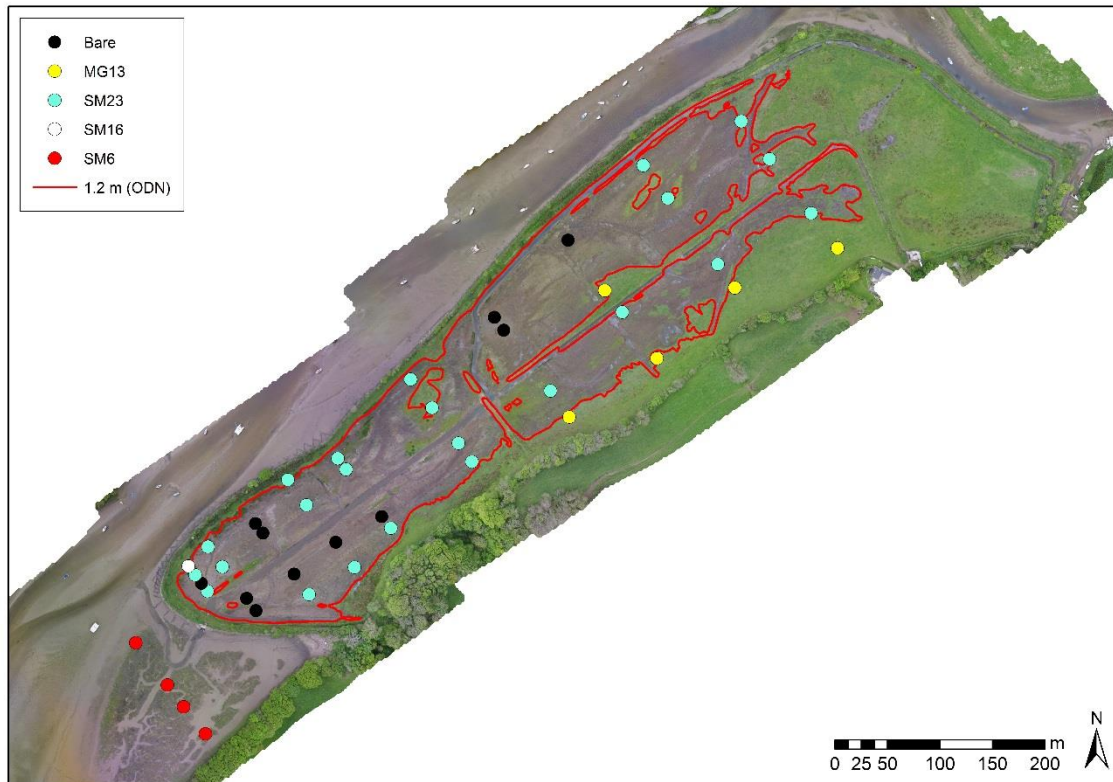


Figure 13 – Aerial photograph of South Efford marsh obtained with UAV flight in June 2016. The ecological sampling locations and the observed NVC Community Types are indicated, as well as the 1.2 m ODN contour line. NVC types: MG13 = *Agrostis stolonifera* - *Alopecurus geniculatus* Grassland; SM23 = *Spergularia marina*-*Puccinellia distans* salt marsh; SM16 = *Juncetum gerardii* salt marsh; SM6 = *Spartina anglica* salt marsh.

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