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Ocean acidification increases the impact of typhoons on algal communities

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Ocean acidification increases the impact of typhoons on algal communities

Highlights:

| 2 3 | | Algal community dynamics studied with three-year monthly surveys at a CO ₂ seep |
|------------------|---|--|
| 4 5 | • | Acidification consistently altered community composition across all seasons |
| 6 7 8 9 | | • Structurally complex communities shifted to degraded 'turf' state with rising <i>p</i> CO ₂ • |
| 10 11 | | Acidification-driven community changes were maintained by typhoon disturbance |
| 12 13 | | • Turf-dominated communities displayed low resistance to typhoons |

14 ABSTRACT

15 Long-term environmental change, sudden pulses of extreme perturbation, or a combination of 16 both can trigger regime shifts by changing the processes and feedbacks which determine community assembly, structure, and function, altering the state of ecosystems. Our 17 18 understanding of the mechanisms that stabilise against regime shifts or lock communities into 19 altered states is limited, yet also critical to anticipating future states, preventing regime shifts, 20 and reversing unwanted state change. Ocean acidification contributes to the restructuring and 21 simplification of algal systems, however the mechanisms through which this occurs and 22 whether additional drivers are involved requires further study. Using monthly surveys over three years at a shallow-water volcanic seep we examined how the composition of algal 23 24 communities change both seasonally and following periods of significant physical disturbance by typhoons at three levels of ocean acidification (equivalent to means of contemporary ~350 25 26 and future ~500 and 900 μ atm *p*CO₂). Consistent with most temperate CO₂ seeps around the 27 world, sites exposed to acidification were increasingly monopolised by structurally simple, fast-growing turf algae, and were clearly different to structurally complex macrophyte-28 29 dominated reference sites. The distinct contemporary and acidified community states were 30 stabilised and maintained at their respective sites by different mechanisms following seasonal typhoon disturbance. Contemporary macroalgal-dominated sites were resistant to typhoon 31 32 damage, recovering to the same community composition and pre-disturbance levels of algal 33 cover and structural complexity. In contrast, significant losses of algal biomass represented a 34 near total ecosystem reset by typhoons for the turf-dominated communities in the elevated 35 pCO_2 sites (i.e. negligible resistance). A combination of disturbance and subsequent turf, but 36 not macrophyte, recovery maintained the same structurally devoid state between years (elevated CO₂ levels promote turf growth following algal removal, inhibiting macroalgal 37 38 recruitment). Thus, ocean acidification may promote shifts in algal systems towards degraded

- 39 ecosystem states, and short-term disturbances which reset successional trajectories may 'lock-
- 40 in' these alternative states of low structural and functional diversity.

41 **1 INTRODUCTION**

Long-term exposure to gradual environmental change, short periods of significant perturbation, 42 43 or a combination of both, can alter the structure and functioning of biological communities (Halpern et al. 2008, Petraitis et al. 2009). Ecosystems 'stable' against disturbance display 44 'resistance' (withstanding change in the face of disturbance) and/or 'recovery' (returning to the 45 46 pre-disturbance state after perturbation), maintaining the same composition, functioning, and 47 internal regulating feedbacks (Meredith et al. 2018). When systems fail to resist or recover following disturbance, regime-shifts may occur (Duarte et al. 2009), tipping ecosystems into 48 49 new, reconfigured states comprised of different sets of species and controlled by different processes (Rocha et al. 2015, Wernberg et al. 2016). Regime shifts are concerning as they can 50 51 lead to the dominance of fast-growing, early successional species assemblages, leading to large losses of ecological and economic resources, such as habitat provision, and biodiversity 52 53 (Hastings and Wysham 2010). Many drivers of state change have been identified (for a review, 54 see deYoung et al. 2008), however our understanding of the mechanisms that stabilise alternate 55 ecosystems against regime shifts or lock communities into altered states is limited, yet critical to anticipating future states and the challenge of reversing unwanted state change. 56

Changes in the intensity and duration of disturbances have further implications for their 57 potential to transition communities to a new state and maintain regime shifts, which may 58 59 enhance when multiple stressors interact and operate together (Folke et al. 2004). Both 'press' (long-term sustained perturbation such as ocean acidification), and 'pulse' (short but intense 60 periods of perturbation such as storm damage) disturbances, can push ecosystems beyond 61 62 tipping points, causing regime shifts (e.g. Hughes 1994, Möllmann et al. 2009, Harvey et al. 2021a) with consequences for alternate states that are both difficult to predict and prevent. 63 Crucially, these effects could exacerbate both if 'press' and 'pulse' disturbances overlap, and 64

65 if the frequency, intensity, or duration disturbances increase (Posey et al. 1996, Peterson 1996,
66 Wang et al. 2016).

67 Community stability varies as environmental conditions and ecosystem states change, 68 determined by the positive (self-reinforcing) feedback loops within communities arising from 69 species interactions and physiological traits that can either resist disturbance or promote 70 recovery after perturbation (Folke et al. 2004, Connell and Ghedini 2015, Nimmo et al. 2015). 71 Ocean acidification, a global 'press' disturbance shifting carbonate chemistry conditions as increased atmospheric CO₂ is drawn down by the ocean, is projected to simplify coastal 72 73 ecosystems due to losses in habitat, structural complexity, and biodiversity (Kroeker et al. 74 2011, Vizzini et al. 2017, Agostini et al. 2018). This ecosystem reorganisation arises due to the 75 unequal impact that ocean acidification has on marine organisms (Connell et al. 2018), acting as both a resource for some primary producers (Koch et al. 2013, Cornwall et al. 2017), and a 76 physiological stressor for other organisms, such as marine calcifiers (Harvey et al. 2018, 77 78 Agostini et al. 2021b). By promoting the loss of certain functional groups and restructuring 79 community composition, ocean acidification might reduce the ability of marine communities 80 to absorb disturbances and remain unchanged, reducing their stability (Folke et al. 2004).

81 To capture the complexity of natural systems and investigate the long-term consequences of 82 ocean acidification at the community-level, an increasing number of studies have used marine 83 CO₂ seeps as natural analogues of future conditions (Hall-Spencer et al. 2008, Milazzo et al. 2014, Agostini et al. 2018, Connell et al. 2018, Foo et al. 2018). Some coastal volcanoes cause 84 85 CO₂ to bubble through the seabed, creating localised gradients of acidification. These offer insights into the long-term consequences of ocean acidification on ecosystems by utilising 86 87 communities that are naturally assembled, complex, and shaped by species interactions (Hall-88 Spencer and Harvey 2019). Studies at CO₂ seeps show that ocean acidification increases the

89 probability of regime shifts in many coastal ecosystems with the large-scale loss of complex 90 habitat-forming species (e.g. kelp forests, corals) and more homogenous algal dominance 91 (Enochs et al. 2015, Connell et al. 2018, Harvey et al. 2021a). The degradation of community 92 stability is often associated with shifts to lower baselines of diversity and complexity (Allison 93 2004, Hughes and Stachowicz 2004). Simplified communities may also be more susceptible to 94 additional perturbation, including short-term 'pulse' disturbance events such as storms. As 95 such, acidification-driven changes in species interactions, composition, and community 96 dynamics, by indirectly altering responses to additional stressors, may promote and entrench 97 regime-shifts, maintaining conditions which enable alternate, opportunistic assemblages to 98 dominate.

Typhoons, tropical cyclones with sustained winds that exceed 33 m s⁻¹ (Japan Meteorological 99 100 Agency 2021), are powerful low-pressure weather systems that can be both disruptive and 101 dangerous to marine organisms and human populations (Zhan et al. 2012). Prevalent 102 throughout the Northern and Western Pacific, typhoons typically form in the tropics before 103 tracking poleward (Zhan et al. 2012, Hsu et al. 2013). Typhoons are highly seasonal, peaking 104 in frequency during late summer to early autumn, therefore likely contribute to annual variation within marine communities (which for algal assemblages remains understudied). Typhoons 105 106 represent a major physical 'pulse' disturbance for coastal ecosystems, generating large waves 107 that can severely damage structurally important habitats such as corals (Done 1992, Harmelin-108 Vivien 1994), mangroves (Diele et al. 2013), seagrasses (Wilson et al. 2020), and macroalgae 109 (Vroom et al. 2005, Hall-Spencer and Harvey 2019, Cattano et al. 2020). Typhoons also alter 110 the community structure and functioning of species living in association with these habitats 111 (Gardner et al. 2005, Teixidó et al. 2013), such as fish (Cattano et al. 2020) and invertebrates 112 (Harmelin-Vivien 1994, Diele et al. 2013) amongst seaweeds (Pocklington et al. 2018, Wernberg et al. 2020). The amount of damage caused is determined by the frequency and 113

114 magnitude of typhoons (Lee et al. 2012, Hsu et al. 2013) and the stability of the marine 115 community against physical disturbance. Rising sea surface temperatures in the western North 116 Pacific over the past 30 years have meant that the average latitude at which typhoons reach 117 their peak intensity has increased (Kossin et al. 2016). The impacts of strong typhoons on 118 temperate coastal communities have and will continue to become increasingly common, 119 exacerbated by climate heating (Webster et al. 2005, Mann and Emanuel 2006, Lee et al. 2012, Murakami et al. 2012, Zhan et al. 2012). 120

121 Prior to this study we observed that algal community coverage and composition varied 122 seasonally and was greatly affected following typhoon disturbance, however this response appeared to differ between the ambient and elevated pCO_2 areas of our study site (Figure 1). 123 124 In this study, we test this observation and investigate algal communities along a natural pCO_2 125 gradient in Japan to assess how their composition varies with increasing pCO_2 . We also investigate the seasonality of algal communities to understand patterns of compositional 126 127 change and how this is altered by ocean acidification. We then examine how ocean 128 acidification-driven changes in community structure mediates their stability (resistance and 129 recovery) against seasonal 'pulses' of typhoon disturbance. By conducting the study over three years, we were able to test the responses (resistance and recovery) of algal communities to 130 131 several typhoon seasons. This temporal replication provides new insights into the future impact 132 of typhoons on acidified oceans.

133 2

MATERIALS AND METHODS

134 2.1 Study site and environmental context

135 Algal community assessments were carried out in three locations along a pCO_2 gradient around Shikine Island, Japan (34°19'9" N, 139° 12'18" E) from September 2016 to December 2019. 136 137 These three locations were 1) a reference pCO_2 area outside the influence of the CO₂ seep, 2)

138 a 'near-future' elevated pCO_2 area, and 3) an 'end-of-the-century' elevated pCO_2 area 139 (hereafter referred to as '350 µatm', '500 µatm' and '900 µatm', respectively). The acidified 140 sites offer potential for space-for-time analyses, serving as analogues for future conditions 141 under projections of changing ocean pH (RCP 8.5 scenario, IPCC 2013). The Shikine-jima 142 seep site has been surveyed since 2014, and its carbonate chemistry and biology are well 143 characterised (Agostini et al. 2015, 2018, 2021a, Harvey et al. 2018, 2019, 2021b, 2021a, Witkowski et al. 2019, Kerfahi et al. 2020, Cattano et al. 2020). Located at 34° north, Shikine 144 145 Island is within the temperate-subtropical biogeographic boundary zone (See Figure S1). Here, 146 the coastal communities on rocky reef habitats are dominated by large stands of macroalgae 147 (Harvey et al. 2021b). The pCO_2 conditions are temporally stable over the long-term, while still following natural diurnal and seasonal variation, and are not confounded by differences in 148 149 temperature, salinity, dissolved oxygen, total alkalinity, nutrients or depth relative to reference 150 sites (Agostini et al. 2015, 2018, Harvey et al. 2019, 2021b, Agostini et al. 2021a). The '350 151 μ atm' location had a mean pH_T of 8.137 ± 0.056 (SD), the '500 μ atm' location area had a mean 152 pH_T of 7.983 ± 0.119 (SD), and the '900 µatm' location had a mean pH_T of 7.781 ± 0.105 (SD). 153 The mean carbonate chemistry of the three locations is presented in Table 1. A full description of the locations is provided in the supplementary material (Figure S1). The area of Shikine 154 155 Island, is exposed to frequent and significant disturbance from tropical cyclones (including 156 typhoons), which peak in frequency during the late summer and early autumn (August-September) (Yumoto and Matsuura 2001). For specific information regarding typhoon 157 158 characteristics and dates during the study period, see Table S1.

159 2.2 Experimental design and analysis

160 2.2.1 Data collection

161 To determine how ocean acidification influences the composition and structure of algal
162 communities over time, eight permanent quadrats (50 x 50 cm) were marked using anchor bolts

(8.5 mm width, 70 mm length) at each of the three locations ('350 µatm', '500 µatm' and '900 µatm'). Quadrats within these locations were deployed haphazardly over a *ca*. 400 m² area with at least 3 m between them, fixed to upward-facing substrata. As algae were the focus of this study, random stratified sampling was used to prevent the inclusion of coral within quadrats. Individual quadrats at each location were photographed (Tough TG-5, Olympus, Japan) monthly (with four exceptions when poor conditions prevented access) a total of 36 times over a period of three years (September 2016 – December 2019) for community assessment.

170 2.2.2 Community Analysis

171 Following methods used by Harvey et al. (2021b), algal community composition was assessed using the ImageJ Fiji processing package. An 8 x 8 grid of points (n = 64) was superimposed 172 173 onto each of the photoquadrats, before the abundance of the algal functional group under each 174 point was identified and recorded. Functional groups were assigned following descriptions provided by Steneck and Dethier (1994), sorting algal groups based on their morphology, 175 176 thallus size and complexity: filamentous algae, foliose algae, corticated foliose algae, 177 corticated macrophytes, turf algae, microalgae, leathery macrophytes, articulated calcareous 178 algae, and crustose coralline algae (CCA). For a list of the dominant species and associated functional groups at each location, refer to the supplementary material (Table S2). The 179 180 complexity of the algal community within each quadrat was also determined following Steneck 181 and Dethier (1994), whereby ranks (0-5) assigned to the biogenic habitat complexity provided 182 by each functional group were combined with their relative abundance within each of the communities that they occurred. The total structural complexity was calculated using the 183 184 following ranking categories: other = 0 (ie. macroinvertebrates), bare rock = 0, microalgae = 1, turf = 1, filamentous algae = 2, foliose algae = 3, CCA = 3, corticated foliose algae = 3.5, 185 186 articulated calcareous algae = 4, corticated macrophytes = 4, and leathery macrophytes = 5. To calculate overall community complexity, the complexity value associated with the group 187

represented under each superimposed point were summed (for example, for a community consisting entirely of filamentous algae: $64 \ge 128$). Complexity values were then standardised between 0 and 1 (dividing total quadrat score by the maximum potential score of 320).

192 2.2.3 Algal community stability

Changes in community percentage cover and complexity were assessed between different 193 194 stages of the typhoon season hereby referred to as 'Before', 'After', and 'Recovery' 195 respectively. For all *p*CO₂ levels, communities possessed their highest level of algal coverage 196 and structural complexity during the 'Before' period (April-June) prior to the peak typhoon 197 season, facilitated by summer growth and low-levels of disturbance. The 'After' period 198 (October-December) refers to the months following the peak typhoon season when typhoons 199 had reduced communities to low levels of algal cover and structural complexity. A 'Recovery' 200 period (January-March) coincided with the start of the spring algal bloom, during which the 201 algae began to grow following disturbance.

202 2.3 Statistical Analysis

Statistical analyses were conducted using R (version 4.04; R Core Team, 2021), with the 'vegan' (Oksanen et al., 2019) and 'lme4' (Bates et al. 2015) packages. The 'ggplot2' (Wickham, 2016) and 'ggpubr' (Kassambara, 2019) packages were used for figure production. For each of the analyses performed, the package and specific function used in R are listed below as 'package::function'.

Differences in community composition (based on the relative percentage cover of different functional groups) between locations (three levels: ('350 µatm', '500 µatm' and '900 µatm') were visualised using principal component analysis (PCA; ggord::biplot). The significance of these differences were then determined using one-way and post-hoc pairwise permutational

analysis of variance (PERMANOVA) comparisons based on Bray–Curtis dissimilarity
(vegan::vegdist and vegan::adonis).

214 To test for differences in the responses of both the percentage cover and structural complexity 215 of algal communities exposed to different levels of typhoon disturbance and pCO_2 , general 216 linear mixed effect models (GLMM) were made, with 'Timing' (three levels: 'Before', 'After', 217 and 'Recovery') and 'Location' (three levels: ('350 µatm', '500 µatm' and '900 µatm') as fixed 218 factors, and 'Quadrat', 'Month', and 'Year' as nested random effects (lme4::lmer and emmeans::pairwise). The typhoon period itself was excluded from these analyses as changes 219 220 in community composition is confounded by the exact timing, strength, and number of storms 221 in relation to sampling, reducing the accuracy of interannual comparisons. The same models were also used to test how differences in algal functional group abundance at different times 222 223 (relative to the peak typhoon season) explained these patterns of community percentage cover 224 and complexity change. By assessing how communities respond following typhoon 225 disturbance, we could assess whether mechanisms of resistance (low immediate impact of 226 disturbance), resilience (rapid return to pre-disturbance states), or both were controlling 227 community dynamics.

228

229 **3 RESULTS**

230 3.1 Community differences between pCO₂ conditions and seasons

For all three locations ('350 µatm', '500 µatm' and '900 µatm'), the percentage cover and community composition of the algal communities showed clear patterns within and between years (Figure 2). High algal coverage was observed during the 'pre-typhoon' spring and summer months (around April to June), whereas the occurrence of bare rock increased in the

autumn to winter months (around September to December) following the typhoon season whenphysical disturbance was greatest (Figure 2).

237 Overall community composition was clearly separated by 'Location', associated with changes 238 in pCO_2 (PERMANOVA: Location, $F_{2,858} = 63.61$, p < 0.001; Figure 3; Table 2). Community 239 composition at the 350 µatm location showed clear and consistent differences to both the 500 µatm and 900 µatm locations (Figure 3; Table S3 PERMANOVA *post-hoc*: all p < 0.01). 240 241 Communities in the 350 µatm location had higher structural complexity, with greater coverage of corticated macrophytes, foliose algae, and crustose coralline algae, compared to the turf and 242 243 bare rock dominated acidified sites (Figure 2). Communities within each pCO_2 conditions varied and responded similarly between seasons, however differences were consistently 244 maintained between pCO_2 conditions (Figure 2). At the 350 µatm location, the spring and 245 246 summer months had extensive growth of species such as Asparagopsis taxiformis and Gelidium elegans (Figure 2; Table S2). Although typhoons removed significant algal biomass (Figure 1) 247 macroalgal persistence through the typhoon season, emerging as 'new growth' in winter, 248 249 formed the basis of the summer community composition, maintaining a more consistent 250 structure between seasons (Figure 2; Table S2).

251 At both the 500 µatm and 900 µatm locations, community composition was more variable 252 (Figures 4 (g-i) and S2). The highly productive spring and summer months saw the rapid 253 growth of thick turf algal mats, compared to mix of corticated foliose alga (Zonaria sp.) and 254 bare rock present during the autumn and winter months due to the impacts of typhoon 255 disturbance – ie. the physical removal of turf (Figures 1 and S2). Despite significant differences between the 500 µatm and 900 µatm locations (Table S3), their community composition had a 256 257 higher degree of overlap due to the similarities in their functional groups (structurally simpler 258 turf algae and corticated foliose algae). For 9 of the 12 months, the composition of communities

at the 500 µatm and 900 µatm locations did not significantly differ, suggesting (given
composition did not remain constant) that they respond to environmental change in similar
ways (Figure 3 PERMANOVA results).

262 3.2 Effects of ocean acidification on community stability

Algal communities under different pCO_2 conditions responded to typhoon disturbance 263 differently (Figures 1 and 2). The percentage cover (%) of algal communities was significantly 264 265 impacted by 'Timing', but not pCO₂ condition ('Location'), with a significant interactive effect demonstrating changes in the response of the community over time (Type III ANOVA: 266 Timing*Location, $X^2 = 46.58$, p = < 0.001; Table 3 and Figure 4a-c). Algal percentage cover 267 $(\% \pm SD)$ 'Before' (93.48 ± 8.21, 90.33 ± 8.58, and 91.27 ± 9.85) and 'After' (72.75 ± 16.50, 268 52.93 ± 19.45 , and 49.15 ± 19.39) typhoons differed significantly at each location ('350 µatm', 269 '500 µatm' and '900 µatm') (Figure 4a-c). The algal coverage of 'Before' and 'Recovery' 270 $(89.92 \pm 12.41, 82.01 \pm 13.76, \text{ and } 78.00 \pm 18.81)$ communities (growing in the months) 271 following typhoons) also significantly differed for acidified communities, but not under 272 273 reference pCO_2 conditions (Table S4). The coverage of reference communities remained more 274 stable over time due to resistance to physical disturbance and algal recovery. Although typhoons had a lasting effect on the coverage of acidified communities, these faced far more 275 276 significant reductions in algal coverage due to typhoon disturbance yet still rapidly recovered, 277 returning to near pre-disturbance conditions (Figure 4a-c; Table S4).

The structural complexity of algal communities was also significantly impacted by both 'Timing' (ANOVA: $X^2 = 33.30$, p = < 0.001) and 'Location' (ANOVA: $X^2 = 77.88$, p = < 0.001; Table 3 and Figure 4d-f). Structural complexity was significantly reduced in the 'After' communities when compared to the 'Before' and 'Recovery' communities either side of the peak typhoon season (Figure 4d-f). The structural complexity of these communities did not

283 significantly differ, indicating a return to pre-disturbance levels within the 6 months following 284 the typhoon season. Unlike percentage cover, which started at similar levels for each of the 285 communities prior to disturbance, the mean structural complexity (\pm SD) of the algal 286 communities in the pre-typhoon season ('Before') was significantly higher for the Reference 287 pCO_2 communities ('350 µatm' = 0.670 ± 0.095; '500 µatm' = 0.552 ± 0.182; '900 µatm' = 0.471 ± 0.215 ; PERMANOVA *post-hoc*: all p < 0.001). Despite displaying similar losses in 288 289 structural complexity 'After' typhoon disturbance $(0.552 \pm 0.139, 0.411 \pm 0.093, and$ 0.345 ± 0.076) and a subsequent 'Recovery' of complexity to pre-disturbance levels at each 290 291 location $(0.670 \pm 0.097, 0.584 \pm 0.101, \text{ and } 0.487 \pm 0.119)$, the communities exposed to 292 increased pCO_2 remained less complex than the communities found at the reference location 293 (Figure 4d-f; Table 3, Figure 5; PERMANOVA *post-hoc*: all p < 0.001). As they began at low 294 levels of structural complexity, the increased pCO_2 communities possessed both less structure 295 to lose following disturbance, but also lower baselines to subsequently recover to.

296 Acidification-driven changes in community composition, and differences in the response of 297 individual algal functional groups to typhoon disturbance explained these changes in 298 community percentage cover and structural complexity (Figure 4g-i; Figure S3; Figure S4). 299 Typhoons had a large impact on algal community composition, which differed significantly 300 between all timing pairs, apart from the 'Before' and 'Recovery' communities at '350 µatm' 301 (Figure 4; Figure S4 PERMANOVA results). Calcareous and filamentous algae did not 302 significantly change in abundance between pCO_2 conditions or timings (remaining at low 303 frequency), whereas fleshy macrophytes had reduced abundance following disturbance at 304 every location. Bare rock and turf algae, which increased in abundance following typhoon 305 disturbance, also had a stronger association with typhoon seasonality, underpinning the main 306 differences between the 'After' communities and those of the 'Before' and 'Recovery' timings 307 for each location (Figure S4).

308 4 DISCUSSION

309 Whilst other environmental factors also influence community assembly, composition, and 310 dynamics, global research at CO₂ seeps has consistently demonstrated clear ecosystem shifts towards simplified communities at sites with elevated pCO_2 levels (Johnson et al. 2012, Enochs 311 312 et al. 2015, Sunday et al. 2017, Connell et al. 2018, Foo et al. 2018, Cattano et al. 2020, Harvey 313 et al. 2021b, 2021a, Agostini et al. 2021a). These acidified communities not only have reduced 314 ecological and structural complexity (Figure 1; Figure 4), but also lower biodiversity, raising 315 concerns about a potential loss of ecosystem services (Hall-Spencer and Harvey 2019). Most 316 studies investigating the ecological effects of ocean acidification have not taken seasonality 317 into account, leaving community dynamics between seasons largely unknown (but see Godbold 318 and Solan 2013, Baggini et al. 2014, Cattano et al. 2020, Harvey et al. 2021b). Here, we found that algal community structure consistently differed between locations with different pCO_2 319 320 levels (Figure 3; Figure S2), and this persisted across seasons despite profound physical 321 disturbance from typhoons, a seasonal environmental stress which maintained the differences. In general, increased levels of pCO_2 caused consistent reductions in macrophyte cover (Figure 322 2), leaving communities dominated by turf algae, characterised by low functional diversity and 323 324 structural complexity (Harvey et al. 2021b, 2021a, Agostini et al. 2021a). Conversely, reference communities featured greater algal diversity, with more extensive cover of 325 326 macrophytes and calcareous algae.

Community differences between elevated pCO_2 and reference locations were seen year-round, however they became more pronounced during key periods of seasonal environmental change: (1) the spring algal bloom, and (2) the autumn peak typhoon season. During spring, algal biomass peaked, and the elevated and high pCO_2 communities became dominated by thick mats of turf algae (Harvey et al. 2019) (Figure 2). Reference communities did not become turfdominated, but instead burgeoned with canopy-forming species (e.g. *Asparagopsis taxiformis*,

333 Gelidium elegans, Sarcodia ceylanica. The lack of turf is likely due to several factors, 334 including the year-round persistence of macrophytes (lack of empty space), removal by 335 physical abrasion from fleshy macroalgal thalli (Cheroske et al. 2000), shifts in top-down 336 control (Cattano et al. 2020), and the lack of CO₂ enrichment to boost their competitive ability 337 (Connell et al. 2018). Autumn typhoons removed the turf algal mats, exposing bare rock; a 338 profound reset of the ecosystem state (Figures 1 and 2). Typhoon disturbance caused less 339 change in the composition, coverage, and structural complexity of communities at reference 340 pCO_2 levels (Figures 4 and 5). Algal persistence as low-biomass 'new growth' through the 341 winter following typhoon disturbance meant that reference community structure was more 342 stable between seasons and did not differ between the winter (after typhoon disturbance) and 343 the following spring. It is important to note that our method of calculating structural complexity 344 was limited in that only the functional group and not the size or biomass of the algae was 345 considered. For example, whilst both winter and summer communities were dominated by 346 macrophytes under ambient conditions (receiving similar scores for structural complexity), 347 they differed markedly; summer communities had high biomass and 3-D structure, whereas 348 winter communities had high coverage of the same functional groups, but lower biomass providing less habitat and structure. 349

350 Short periods of intense physical disturbance interacted with the effects of long-term 351 acidification on community composition exposure to alter their dynamics. Acidified 352 assemblages were less resistant (reduction of algal coverage and structural complexity) to 353 disturbance from typhoons during seasonal periods of high storm frequency and strength with 354 more pronounced losses at higher pCO_2 levels (Figures 2, 4, and 5). Community composition and associated life-histories, traits, and physiological susceptibilities underpinned these 355 356 differences in community resistance. Complex macroalgal habitats can buffer the influence of climate change (Krause-Jensen et al. 2018), such as mitigating storm surge intensity due to 357

their capacity to attenuate water flow (Rosman et al. 2007) resisting change until disturbances of a greater magnitude are experienced (Schowalter 2006, Levin and Möllmann 2015). The weaker attachment strength and low complexity of the algal groups associated with acidified conditions, such as turfs (Peterson 1996, Harvey et al. 2021a), leaves acidified communities more susceptible to removal by physical disturbance (Wada et al. 2021), representing low levels of resistance (Figures 4 and 5).

364 Differences were also observed in the recovery of algal communities following disturbance 365 due to acidification-driven shifts in succession trajectories. Despite experiencing a greater 366 initial impact, assemblages found under elevated pCO_2 rapidly recovered to levels of algal 367 coverage near to those seen pre-disturbance. Enrichment of CO₂ promotes the already fast and opportunistic growth of turf algae, leading to rapid turf expansion during springtime 368 369 community succession following the ecosystem reset of acidified sites by the typhoon season (Connell et al. 2018, Ferreira et al. 2021). Whilst turf coverage did not fully recover within 6-370 371 months following typhoons, the non-turf occupied space remained bare and macroalgal 372 communities did not recover in their place (Figure 2). Previous studies have shown that turf 373 removal can promote the recovery of macroalgal canopies (Gorman and Connell 2009), however here turf-dominance was consistently regained at the acidified sites. Due to the 374 375 proximity of our sites and the scale of typhoons, our study lacked a 'disturbance-free' acidified 376 treatment. Despite this, previous studies suggest that even in the absence of disturbance, due 377 to turf-mediated reinforcing feedback loops (sediment trapping, alteration of substrate chemistry, physicochemical environment change, and recruitment inhibition) the development 378 379 of macroalgal assemblages will not eventually replace established turf communities under acidified conditions (Harvey et al. 2021a). As the acidified communities were less complex 380 381 than those at reference locations before typhoon disturbance, their recovery only had a low baseline to reach (Figure 4). This structurally-devoid community represents an alternate stable 382

state maintained by a combination of 'boosted' turf growth , the turf-mediated inhibition of slower-growing, structurally complex macrophytes (Ghedini and Connell 2017b) , and regular 'resetting' of succession by physical disturbance under elevated pCO_2 conditions (Gorman and Connell 2009, Harvey et al. 2021b).

Similar to patterns seen in systems enriched with nutrients (Worm et al. 1999, Gorman et al. 387 388 2009) shifts in dominance towards fast-growing r-selected species have increasingly been 389 documented under acidified conditions (Connell and Russell 2010, Harvey et al. 2019, Agostini et al. 2021a, Harvey et al. 2021a), whereby opportunistic species rapidly monopolise primary 390 391 space, replacing algal canopies (Airoldi 2003, Gorman and Connell 2009). Under reference 392 conditions grazing (top-down control) and competition (bottom-up effects) rapidly exclude turf 393 algae, preventing such a regime shift (Kéfi et al. 2016). Concern is growing that rising pCO_2 394 will threaten the feedbacks that stop turfs from monopolising space, and additional coincidental 395 biotic and abiotic change may further reinforce the competitive advantages turf gain over 396 slower growing, typically dominant groups (such as macroalgae; Hughes 1994, Connell and 397 Russell 2010, Ghedini and Connell 2017a), making state transitions within algal systems more 398 likely (Ghedini et al. 2015, Harvey et al. 2021a).

399 The strength and likelihood of typhoons in temperate latitudes is being enhanced by global 400 warming (Webster et al. 2005, Mann and Emanuel 2006, Lee et al. 2012, Murakami et al. 2012, 401 2012, Zhan et al. 2012, Wang et al. 2022). The competitive edge of turf algae over macroalgal canopy formers are limited to their early life-history stages (O'Brien and Scheibling 2018) and 402 403 is therefore weak under stable conditions. Reductions in recovery time between typhoons and extension of the typhoon season alters the competitive balance of algal communities however, 404 405 particularly those exposed to other stressors such as ocean acidification (Kroeker et al. 2013, 406 Pessarrodona et al. 2021). Although communities with extensive foundation species coverage

display resistance to turf-system regime shifts (Falkenberg et al. 2012), if sufficiently highlevels of disturbance were to remove slower-growing macroalgal canopies, their recovery and
recruitment could be inhibited in favour of *r*-selected turf algae (O'Brien and Scheibling 2018,
Harvey et al. 2021b). Facilitated by rapid recruitment and fast rates of growth, turf algae might
monopolize the space created by stronger typhoons, during the shortened periods occurring
between disturbance events, and following the end of the typhoon season, helping to further
drive regime shifts from macroalgal to turf-dominated communities (Pessarrodona et al. 2021).

414 Our findings reveal that the globally observed tendency for complex ecosystem states to be 415 replaced by simple ecosystem states due to human impacts is likely to become entrenched by 416 ocean acidification. Algal communities differed in their composition with increasing pCO_2 and 417 this was maintained across the year, with the largest changes coinciding with associated 418 environmental change (nutrient rich spring algal blooms and stormy typhoon seasons). 419 Between years, the same ecosystem states were maintained within each respective pCO_2 420 condition. Our findings suggest that the acidification-driven changes in algal community 421 structure (with different life-histories and traits altering the mechanism of response to 422 disturbance by typhoons) have significant implications for the maintenance and stability of the different ecosystem states observed between locations (Figure 5). Over three annual cycles of 423 424 near total ecosystem reset by typhoons following extensive algal removal (i.e. negligible 425 resistance), communities in acidified conditions rapidly returned to the same state (i.e. high 426 recovery). This comprised of dominance by highly productive, opportunistic algal turfs which 427 provide little structural complexity but cover the substrate and inhibit macroalgal recruitment. 428 Whilst the persistence of contemporary, macroalgal-dominated communities relies on both 429 initial resistance to and subsequent recovery from typhoons, the combination of ocean 430 acidification and typhoons increases the probability of ecosystem shifts to simpler states 431 dominated by fast-growing 'weedy' species with low resistance but fast recovery that 'locks-

in' the new ecosystem state (Harvey et al. 2021a). The anticipated increase in typhoon
intensity, frequency, and duration would further reinforce this low complexity state, reducing
the time for algal recovery and keeping the community in an early successional stage.
Crucially, such shifts that reduce ecosystem complexity in the marine realm can also reduce
ecosystem productivity and associated species diversity. Future work should examine how
reductions of community stability against disturbance might change ecosystem capacity to
maintain original functionality following perturbation.

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450 **COMPETING INTERESTS**

451 The authors declare no conflicts of interest.

453 DATA AVAILABILITY

454 [Upon acceptance] Raw data used to create Figs 2-4 and Tables 2-3 will be supplied and stored
455 in the Pangaea repository (http://pangaea.de).

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457 **REFERENCES**

- Agostini, S., B. P. Harvey, M. Milazzo, S. Wada, K. Kon, N. Floc'h, K. Komatsu, M.
 Kuroyama, and J. M. Hall-Spencer. 2021a. Simplification, not "tropicalization", of
 temperate marine ecosystems under ocean warming and acidification. Global Change
 Biology 27:4771–4784.
- 462 Agostini, S., B. P. Harvey, S. Wada, K. Kon, M. Milazzo, K. Inaba, and J. M. Hall-Spencer.
 463 2018. Ocean acidification drives community shifts towards simplified non-calcified
 464 habitats in a subtropical-temperate transition zone. Scientific Reports 8:11354.
- Agostini, S., F. Houlbrèque, T. Biscéré, B. P. Harvey, J. M. Heitzman, R. Takimoto, W.
 Yamazaki, M. Milazzo, and R. Rodolfo-Metalpa. 2021b. Greater mitochondrial energy
 production provides resistance to ocean acidification in "winning" hermatypic corals.
 Frontiers in Marine Science 7:1235.
- Agostini, S., S. Wada, K. Kon, A. Omori, H. Kohtsuka, H. Fujimura, Y. Tsuchiya, T. Sato, H.
 Shinagawa, Y. Yamada, and K. Inaba. 2015. Geochemistry of two shallow CO₂ seeps
 in Shikine Island (Japan) and their potential for ocean acidification research. Regional
 Studies in Marine Science 2, Supplement:45–53.
- 473 Airoldi, L. 2003. The effects of sedimentation on rocky coast assemblages. Oceanography and
 474 Marine Biology: An Annual Review.
- Allison, G. 2004. The Influence of Species Diversity and Stress Intensity on Community
 Resistance and Resilience. Ecological Monographs 74:117–134.

- 477 Baggini, C., M. Salomidi, E. Voutsinas, L. Bray, E. Krasakopoulou, and J. M. Hall-Spencer.
- 478 2014. Seasonality affects macroalgal community response to increases in *p*CO₂. PLoS
 479 One 9:e106520.
- 480 Cattano, C., S. Agostini, B. P. Harvey, S. Wada, F. Quattrocchi, G. Turco, K. Inaba, J. M. Hall481 Spencer, and M. Milazzo. 2020. Changes in fish communities due to benthic habitat
 482 shifts under ocean acidification conditions. Science of The Total Environment
 483 725:138501.
- 484 Connell, S. D., Z. A. Doubleday, N. R. Foster, S. B. Hamlyn, C. D. Harley, B. Helmuth, B. P.
 485 Kelaher, I. Nagelkerken, K. L. Rodgers, G. Sarà, and B. D. Russell. 2018. The duality
 486 of ocean acidification as a resource and a stressor. Ecology 99:1005–1010.
- 487 Connell, S. D., and G. Ghedini. 2015. Resisting regime-shifts: the stabilising effect of
 488 compensatory processes. Trends in Ecology & Evolution 30:513–515.
- 489 Connell, S. D., and B. D. Russell. 2010. The direct effects of increasing CO₂ and temperature
 490 on non-calcifying organisms: increasing the potential for phase shifts in kelp forests.
 491 Proceedings of the Royal Society B: Biological Sciences 277:1409–1415.
- 492 Cornwall, C. E., A. T. Revill, J. M. Hall-Spencer, M. Milazzo, J. A. Raven, and C. L. Hurd.
- 493 2017. Inorganic carbon physiology underpins macroalgal responses to elevated CO₂.
 494 Scientific Reports 7:46297.
- deYoung, B., M. Barange, G. Beaugrand, R. Harris, R. I. Perry, M. Scheffer, and F. Werner.
 2008. Regime shifts in marine ecosystems: detection, prediction and management.
 Trends in Ecology & Evolution 23:402–409.
- Diele, K., D. M. Tran Ngoc, S. J. Geist, F. W. Meyer, Q. H. Pham, U. Saint-Paul, T. Tran, and
 U. Berger. 2013. Impact of typhoon disturbance on the diversity of key ecosystem
 engineers in a monoculture mangrove forest plantation, Can Gio Biosphere Reserve,
 Vietnam. Global and Planetary Change 110:236–248.

- 502 Done, T. J. 1992. Effects of tropical cyclone waves on ecological and geomorphological
 503 structures on the Great Barrier Reef. Continental Shelf Research 12:859–872.
- 504 Enochs, I. C., D. P. Manzello, E. M. Donham, G. Kolodziej, R. Okano, L. Johnston, C. Young,
- 505 J. Iguel, C. B. Edwards, M. D. Fox, L. Valentino, S. Johnson, D. Benavente, S. J. Clark,
- R. Carlton, T. Burton, Y. Eynaud, and N. N. Price. 2015. Shift from coral to macroalgae
 dominance on a volcanically acidified reef. Nature Climate Change 5:1083–1088.
- Falkenberg, L. J., B. D. Russell, and S. D. Connell. 2012. Stability of Strong Species
 Interactions Resist the Synergistic Effects of Local and Global Pollution in Kelp
 Forests. PLOS ONE 7:e33841.
- Ferreira, C. M., S. D. Connell, S. U. Goldenberg, and I. Nagelkerken. 2021. Positive species
 interactions strengthen in a high-CO 2 ocean. Proceedings of the Royal Society B:
 Biological Sciences 288:20210475.
- Folke, C., S. Carpenter, B. Walker, M. Scheffer, T. Elmqvist, L. Gunderson, and C. S. Holling.
 2004. Regime Shifts, Resilience, and Biodiversity in Ecosystem Management. Annual
 Review of Ecology, Evolution, and Systematics 35:557–581.
- 517 Foo, S. A., M. Byrne, E. Ricevuto, and M. C. Gambi. 2018. The carbon dioxide vents of Ischia,
- 518 Italy, a natural system to assess impacts of ocean acidification on marine ecosystems:
- An overview of research and comparisons with other vent systems. Oceanography and
 Marine Biology: An Annual Review 56:237–310.
- Gardner, T. A., I. M. Côté, J. A. Gill, A. Grant, and A. R. Watkinson. 2005. Hurricanes and
 Caribbean Coral Reefs: Impacts, Recovery Patterns, and Role in Long-Term Decline.
 Ecology 86:174–184.
- Ghedini, G., and S. D. Connell. 2017a. Moving ocean acidification research beyond a simple
 science: Investigating ecological change and their stabilizers. Food Webs 13:53–59.

- 526 Ghedini, G., and S. D. Connell. 2017b. Moving ocean acidification research beyond a simple 527 science: Investigating ecological change and their stabilizers. Food Webs 13:53–59.
- 528 Ghedini, G., B. D. Russell, and S. D. Connell. 2015. Trophic compensation reinforces
 529 resistance: herbivory absorbs the increasing effects of multiple disturbances. Ecology
 530 Letters 18:182–187.
- Godbold, J. A., and M. Solan. 2013. Long-term effects of warming and ocean acidification are
 modified by seasonal variation in species responses and environmental conditions.
 Philosophical Transactions of the Royal Society B: Biological Sciences 368:20130186.
- Gorman, D., and S. D. Connell. 2009. Recovering subtidal forests in human-dominated
 landscapes. Journal of Applied Ecology 46:1258–1265.
- Gorman, D., B. D. Russell, and S. D. Connell. 2009. Land-to-sea connectivity: linking humanderived terrestrial subsidies to subtidal habitat change on open rocky coasts. Ecological
 Applications: A Publication of the Ecological Society of America 19:1114–1126.
- Hall-Spencer, J. M., and B. P. Harvey. 2019. Ocean acidification impacts on coastal ecosystem
 services due to habitat degradation. Emerging Topics in Life Sciences 3:197–206.
- 541 Hall-Spencer, J. M., R. Rodolfo-Metalpa, S. Martin, E. Ransome, M. Fine, S. M. Turner, S. J.
- 542Rowley, D. Tedesco, and M.-C. Buia. 2008. Volcanic carbon dioxide vents show543ecosystem effects of ocean acidification. Nature 454:96–99.
- 544 Halpern, B. S., S. Walbridge, K. A. Selkoe, C. V. Kappel, F. Micheli, C. D'Agrosa, J. F. Bruno,
- 545 K. S. Casey, C. Ebert, H. E. Fox, R. Fujita, D. Heinemann, H. S. Lenihan, E. M. P.
- Madin, M. T. Perry, E. R. Selig, M. Spalding, R. Steneck, and R. Watson. 2008. A
 Global Map of Human Impact on Marine Ecosystems. Science 319:948–952.
- 548 Harmelin-Vivien, M. L. 1994. The Effects of Storms and Cyclones on Coral Reefs: A Review.
 549 Journal of Coastal Research:211–231.

- Harvey, B. P., S. Agostini, K. Kon, S. Wada, and J. M. Hall-Spencer. 2019. Diatoms dominate
 and alter marine food-webs when CO₂ rises. Diversity 11:242.
- Harvey, B. P., S. Agostini, S. Wada, K. Inaba, and J. M. Hall-Spencer. 2018. Dissolution: the
 Achilles' heel of the triton shell in an acidifying ocean. Frontiers in Marine Science
 554 5:371.
- Harvey, B. P., R. Allen, S. Agostini, L. J. Hoffmann, K. Kon, T. C. Summerfield, S. Wada,
 and J. M. Hall-Spencer. 2021a. Feedback mechanisms stabilise degraded turf algal
 systems at a CO₂ seep site. Communications Biology 4:219.
- Harvey, B. P., K. Kon, S. Agostini, S. Wada, and J. M. Hall-Spencer. 2021b. Ocean
 acidification locks algal communities in a species-poor early successional stage. Global
 Change Biology 27:2174–2187.
- Hastings, A., and D. B. Wysham. 2010. Regime shifts in ecological systems can occur with no
 warning. Ecology Letters 13:464–472.
- Hsu, P.-C., C.-R. Ho, S.-J. Liang, and N.-J. Kuo. 2013. Impacts of Two Types of El Niño and
 La Niña Events on Typhoon Activity. Advances in Meteorology 2013:e632470.
- Hughes, A. R., and J. J. Stachowicz. 2004. Genetic diversity enhances the resistance of a
 seagrass ecosystem to disturbance. Proceedings of the National Academy of Sciences
 101:8998–9002.
- Hughes, T. P. 1994. Catastrophes, phase shifts, and large-scale degradation of a Caribbean
 coral reef. Science 265:1547–1551.
- 570 Japan Meteorological Agency. 2021. Japan Meteorological Agency | Tropical Cyclone
 571 Information.
- 572 https://www.data.jma.go.jp/multi/cyclone/cyclone_caplink.html?lang=en.

- 573 Johnson, V. R., B. D. Russell, K. E. Fabricius, C. Brownlee, and J. M. Hall-Spencer. 2012.
- 574 Temperate and tropical brown macroalgae thrive, despite decalcification, along natural
 575 CO₂ gradients. Glob. Chang. Biol. 18:2792–2803.
- 576 Kéfi, S., M. Holmgren, and M. Scheffer. 2016. When can positive interactions cause alternative
 577 stable states in ecosystems? Functional Ecology 30:88–97.
- 578 Kerfahi, D., B. P. Harvey, S. Agostini, K. Kon, R. Huang, J. M. Adams, and J. M. Hall-Spencer.
- 579 2020. Responses of intertidal bacterial biofilm communities to increasing *p*CO₂.
 580 Marine Biotechnology.
- Koch, M., G. Bowes, C. Ross, and X.-H. Zhang. 2013. Climate change and ocean acidification
 effects on seagrasses and marine macroalgae. Global Change Biology 19:103–132.
- Kossin, J. P., K. A. Emanuel, and S. J. Camargo. 2016. Past and Projected Changes in Western
 North Pacific Tropical Cyclone Exposure. Journal of Climate 29:5725–5739.
- 585 Krause-Jensen, D., P. Lavery, O. Serrano, N. Marbà, P. Masque, and C. M. Duarte. 2018.
 586 Sequestration of macroalgal carbon: the elephant in the Blue Carbon room. Biology
 587 Letters 14:20180236.
- 588 Kroeker, K. J., F. Micheli, and M. C. Gambi. 2013. Ocean acidification causes ecosystem shifts
 589 via altered competitive interactions. Nature Climate Change 3:156–159.
- Kroeker, K. J., F. Micheli, M. C. Gambi, and T. R. Martz. 2011. Divergent ecosystem responses
 within a benthic marine community to ocean acidification. Proceedings of the National
 Academy of Sciences of the United States of America 108:14515–14520.
- 593 Lee, T.-C., T. R. Knutson, H. Kamahori, and M. Ying. 2012. Impacts of Climate Change on
- Tropical Cyclones in the Western North Pacific Basin. Part I: Past Observations.
 Tropical Cyclone Research and Review 1:213–235.

- 596 Levin, P. S., and C. Möllmann. 2015. Marine ecosystem regime shifts: challenges and
- 597 opportunities for ecosystem-based management. Philosophical Transactions of the
 598 Royal Society B: Biological Sciences 370:20130275.
- Mann, M. E., and K. A. Emanuel. 2006. Atlantic hurricane trends linked to climate change.
 Eos, Transactions American Geophysical Union 87:233–241.
- 601 Milazzo, M., R. Rodolfo-Metalpa, V. B. S. Chan, M. Fine, C. Alessi, V. Thiyagarajan, J. M.
- Hall-Spencer, and R. Chemello. 2014. Ocean acidification impairs vermetid reef
 recruitment. Scientific Reports 4:4189.
- Möllmann, C., R. Diekmann, B. Müller-Karulis, G. Kornilovs, M. Plikshs, and P. Axe. 2009.
- Reorganization of a large marine ecosystem due to atmospheric and anthropogenic
 pressure: a discontinuous regime shift in the Central Baltic Sea. Global Change Biology
 15:1377–1393.
- Murakami, H., Y. Wang, H. Yoshimura, R. Mizuta, M. Sugi, E. Shindo, Y. Adachi, S.
 Yukimoto, M. Hosaka, S. Kusunoki, T. Ose, and A. Kitoh. 2012. Future Changes in
 Tropical Cyclone Activity Projected by the New High-Resolution MRI-AGCM.
 Journal of Climate 25:3237–3260.
- Nimmo, D. G., R. Mac Nally, S. C. Cunningham, A. Haslem, and A. F. Bennett. 2015. Vive la
 résistance: reviving resistance for 21st century conservation. Trends in Ecology &
 Evolution 30:516–523.
- O'Brien, J. M., and R. E. Scheibling. 2018. Turf wars: competition between foundation and
 turf-forming species on temperate and tropical reefs and its role in regime shifts. Marine
 Ecology Progress Series 590:1–17.
- 618 Pessarrodona, A., K. Filbee-Dexter, T. Alcoverro, J. Boada, C. J. Feehan, S. Fredriksen, S. P.
- 619 Grace, Y. Nakamura, C. A. Narvaez, K. M. Norderhaug, and T. Wernberg. 2021.

- Homogenization and miniaturization of habitat structure in temperate marine forests.
 Global Change Biology 27:5262–5275.
- Peterson, C. G. 1996. 13 Response of Benthic Algal Communities to Natural Physical
 Disturbance. Pages 375–402 *in* R. J. Stevenson, M. L. Bothwell, and R. L. Lowe,
 editors. Algal Ecology. Academic Press, San Diego.
- 625 Petraitis, P. S., E. T. Methratta, E. C. Rhile, N. A. Vidargas, and S. R. Dudgeon. 2009.
- 626 Experimental confirmation of multiple community states in a marine ecosystem.627 Oecologia 161:139–148.
- 628 Pocklington, J. B., S. R. Jenkins, A. Bellgrove, M. J. Keough, T. D. O'Hara, P. E. Masterson-
- Algar, and S. J. Hawkins. 2018. Disturbance alters ecosystem engineering by a canopy forming alga. Journal of the Marine Biological Association of the United Kingdom
- 63198:687–698.
- Posey, M., W. Lindberg, T. Alphin, and F. Vose. 1996. Influence of Storm Disturbance on an
 Offshore Benthic Community. Bulletin of Marine Science 59:7.
- Rocha, J., J. Yletyinen, R. Biggs, T. Blenckner, and G. Peterson. 2015. Marine regime shifts:
- drivers and impacts on ecosystems services. Philosophical Transactions of the Royal
 Society B: Biological Sciences 370.
- Rosman, J. H., J. R. Koseff, S. G. Monismith, and J. Grover. 2007. A field investigation into
 the effects of a kelp forest (Macrocystis pyrifera) on coastal hydrodynamics and
 transport. Journal of Geophysical Research: Oceans 112.
- 640 Schowalter, T. D. 2006. Insect Ecology: An Ecosystem Approach. 2nd Edition. Academic
 641 Press.
- 642 Steneck, R. S., and M. N. Dethier. 1994. A functional group approach to the structure of algal643 dominated communities. Oikos 69:476–498.

- 644 Sunday, J. M., K. E. Fabricius, K. J. Kroeker, K. M. Anderson, B. N. E, J. P. Barry, S. D.
- 645 Connell, S. Dupont, B. Gaylord, J. M. Hall-Spencer, T. Klinger, M. Milazzo, P. L.
- 646 Munday, B. D. Russell, E. Sanford, V. Thiyagarajan, M. L. H. Vaughan, S.
- 647 Widdicombe, and C. D. G. Harley. 2017. Ocean acidification can mediate biodiversity
- shifts by changing biogenic habitat. Nature Climate Change 7:81–85.
- 649 Teixidó, N., E. Casas, E. Cebrián, C. Linares, and J. Garrabou. 2013. Impacts on Coralligenous
 650 Outcrop Biodiversity of a Dramatic Coastal Storm. PLOS ONE 8:e53742.
- 651 Vizzini, S., B. Martínez-Crego, C. Andolina, A. Massa-Gallucci, S. D. Connell, and M. C.
 652 Gambi. 2017. Ocean acidification as a driver of community simplification *via* the
- 653 collapse of higher-order and rise of lower-order consumers. Scientific Reports 7:4018.
- Wada, S., S. Agostini, B. P. Harvey, Y. Omori, and J. M. Hall-Spencer. 2021. Ocean
 acidification increases phytobenthic carbon fixation and export in a warm-temperate
 system. Estuarine, Coastal and Shelf Science 250:107113.
- Wang, G., L. Wu, W. Mei, and S.-P. Xie. 2022. Ocean currents show global intensification of
 weak tropical cyclones. Nature 611:496–500.
- Wang, X., W. Wang, and C. Tong. 2016. A review on impact of typhoons and hurricanes on
 coastal wetland ecosystems. Acta Ecologica Sinica 36:23–29.
- Webster, P. J., G. J. Holland, J. A. Curry, and H.-R. Chang. 2005. Changes in Tropical Cyclone
 Number, Duration, and Intensity in a Warming Environment. Science 309:1844–1846.
- 663 Wernberg, T., S. Bennett, R. C. Babcock, T. de Bettignies, K. Cure, M. Depczynski, F. Dufois,
- J. Fromont, C. J. Fulton, R. K. Hovey, E. S. Harvey, T. H. Holmes, G. A. Kendrick, B.
- 665 Radford, J. Santana-Garcon, B. J. Saunders, D. A. Smale, M. S. Thomsen, C. A.
- 666 Tuckett, F. Tuya, M. A. Vanderklift, and S. Wilson. 2016. Climate-driven regime shift
- of a temperate marine ecosystem. Science 353:169–172.

- 668 Wernberg, T., M. Couraudon-Réale, F. Tuya, and M. Thomsen. 2020. Disturbance intensity,
- disturbance extent and ocean climate modulate kelp forest understory communities.
 Marine Ecology Progress Series 651:57–69.
- 671 Wilson, S. S., B. T. Furman, M. O. Hall, and J. W. Fourqurean. 2020. Assessment of Hurricane
- 672 Irma Impacts on South Florida Seagrass Communities Using Long-Term Monitoring
 673 Programs. Estuaries and Coasts 43:1119–1132.
- Witkowski, C. R., S. Agostini, B. P. Harvey, M. T. J. van der Meer, J. S. Sinninghe Damsté,
 and S. Schouten. 2019. Validation of carbon isotope fractionation in algal lipids as a *p*CO₂ proxy using a natural CO₂ seep (Shikine Island, Japan). Biogeosciences 16:4451–
 4461.
- Worm, B., H. K. Lotze, C. Boström, R. Engkvist, V. Labanauskas, and U. Sommer. 1999.
 Marine diversity shift linked to interactions among grazers, nutrients and propagule
 banks. Marine Ecology Progress Series 185:309–314.
- Wumoto, M., and T. Matsuura. 2001. Interdecadal Variability of Tropical Cyclone Activity in
 the Western North Pacific. Journal of the Meterological Society of Japan 79:23–25.
- Zhan, R., Y. Wang, and M. Ying. 2012. Seasonal Forecasts of Tropical Cyclone Activity Over
 the Western North Pacific: A Review. Tropical Cyclone Research and Review 1:307–
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699 Table 1. Summary of the carbonate chemistry for the 350 μatm and 900 μatm locations. The pH_T (350 μatm, n = 1964; 500 μatm, n = 1760; 900 μatm,

700 n = 10,818), salinity (350 μatm, n = 1964; 500 μatm, n = 1760; 900 μatm, n = 10,818), and total alkalinity (A_T; 350 μatm, n = 56; 500 μatm, n = 10; 900

701 µatm, n = 47) are measured values. All other values were calculated using the carbonate chemistry system analysis program CO2SYS: Seawater pCO_2 ,

702 dissolved inorganic carbon (DIC), bicarbonate (HCO₃⁻), carbonate (CO₃²⁻), carbon dioxide (CO₂), saturation states for calcite (Ωcalcite), and aragonite

703 (Ωaragonite). Values are presented as mean, with standard deviation below. NOTE: Carbonate chemistry data are sourced from Agostini et al. 2018,

704 Harvey et al. 2019, and Harvey, Kon, et al., 2021.

| Location | $\mathbf{p}\mathbf{H}_{\mathrm{T}}$ | Salinity | A _T | pCO ₂ | DIC | HCO ₃ - | CO3 ²⁻ | Ωcalcite | Ωaragonite |
|------------|-------------------------------------|----------|--------------------------|------------------|--------------------------|--------------------------|--------------------------|----------|------------|
| | | (psu) | (µmol kg ⁻¹) | (µatm) | (µmol kg ⁻¹) | (µmol kg ⁻¹) | (µmol kg ⁻¹) | | |
| '350 μatm' | 8.137 | 34.504 | 2264.29 | 316.057 | 1962.694 | 1740.629 | 211.979 | 5.087 | 3.301 |
| | 0.056 | 0.427 | 15.34 | 47.466 | 34.376 | 55.084 | 22.221 | 0.534 | 0.348 |
| '500 μatm' | 7.990 | 34.17 | 2264.35 | 471.30 | 2031.16 | 1852.31 | 163.71 | 3.94 | 2.56 |
| | 0.086 | 0.44 | 16.62 | 117.01 | 40.16 | 60.59 | 24.88 | 0.59 | 0.39 |
| (000 ustm? | 7.788 | 34.351 | 2268.33 | 841.148 | 2125.785 | 1984.889 | 115.150 | 2.771 | 1.805 |
| '900 μatm' | 0.106 | 0.484 | 19.45 | 291.762 | 39.381 | 52.510 | 21.308 | 0.512 | 0.336 |

| Term | df | Sum Sq. | Mean Sq. | F | р |
|--------------------------|-----|---------|----------|-------|-----------|
| <i>p</i> CO ₂ | 2 | 21.96 | 10.98 | 63.13 | 0.001 *** |
| Residuals | 858 | 148.1 | 0.173 | | |
| Total | 860 | 170.0 | | | |

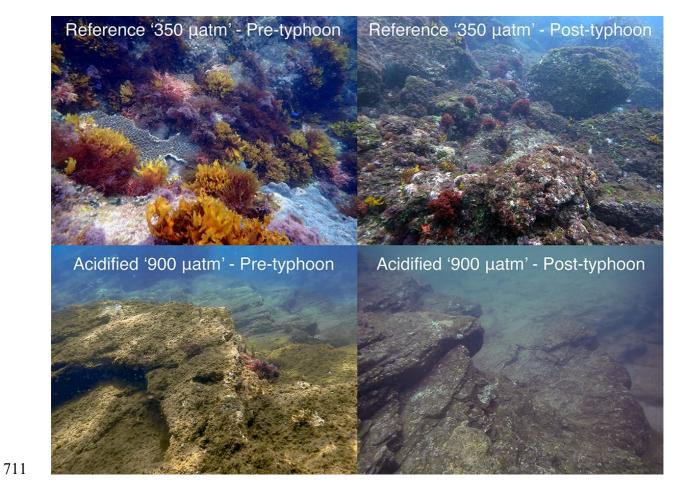
Table 2. PERMANOVA summary for the effects of pCO_2 ('350 µatm' vs. '500 µatm' vs. '900µatm') on algal community composition. For p-values, * p < 0.05, ** p < 0.01, *** p < 0.001.</td>

Table 3. Type III Wald chi-square Analysis of Deviance summary for the effect of timing ('Before' vs. 'After' vs. 'Recovery') and pCO_2 condition ('350 µatm' vs. '500 µatm' vs. '900 µatm') on algal community (a) percentage cover (%) and (b) complexity. For *p*-values, * p < 0.05, ** p < 0.01, *** p < 0.001.

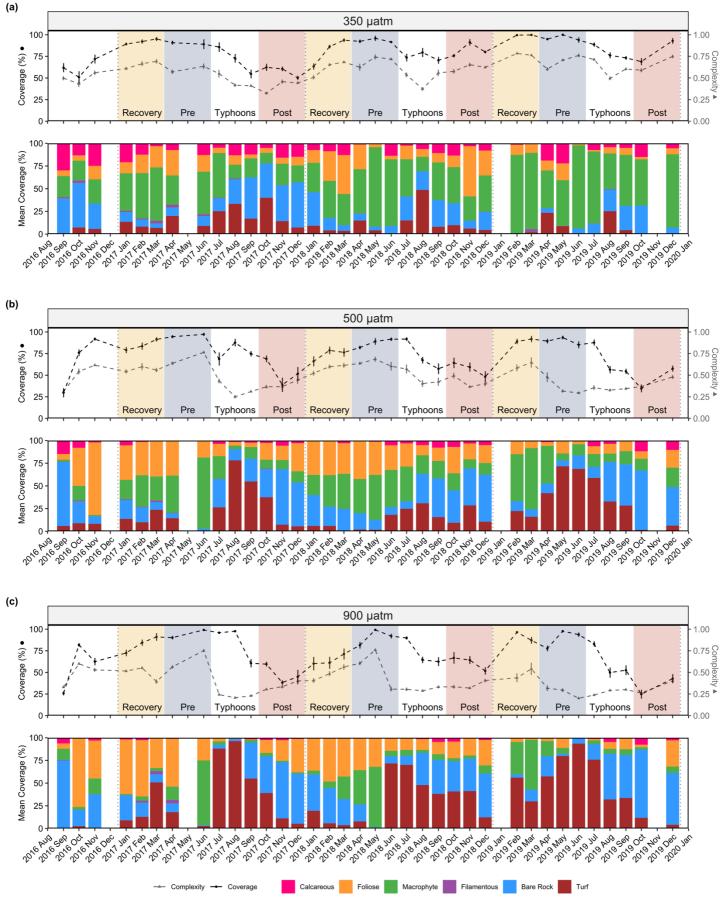
| a) Cover (%) | Term | X^2 | df | р |
|---------------|------------------------------|------------------------|----------------|----------------------|
| | Intercept | 1875 | 1 | < 0.001 *** |
| | Timing | 73.43 | 2 | < 0.001 *** |
| | pCO_2 | 1.208 | 2 | 0.547 |
| | Timing x p CO ₂ | 46.58 | 4 | < 0.001 *** |
| | | | | |
| b) Complexity | Term | X^2 | df | р |
| b) Complexity | Term Intercept | X ² 1539 | <i>df</i> 1 | <i>p</i> < 0.001 *** |
| b) Complexity | | | - | |
| b) Complexity | Intercept | 1539 | 1 | < 0.001 *** |

709

OA reduces resistance, but not resilience of algal communities



- Figure 1. Examples of seascapes found at Reference '350 μ atm' and acidified '900 μ atm' pCO_2
- 713 locations before typhoon disturbance (pre-typhoon) and after (post-typhoon).



OA reduces resistance, but not resilience of algal communities

Figure 2. Average percentage cover (Black); complexity (Grey); and mean coverage (by functional group) of algal communities across each sampling month over a three-year sampling period at (a) 350 µatm CO₂; (b) 500 µatm CO₂; and (c) 900 µatm CO₂. Pre-typhoon periods ("Pre"; April to June; blue), post-typhoon periods ("Post"; October to December; red), and recovery periods ("Recovery"; January to March; yellow) are indicated on the figure. Typhoon periods (July to September) are marked as "Typhoons". Dotted vertical lines denote the start of each new year (2017-2020).



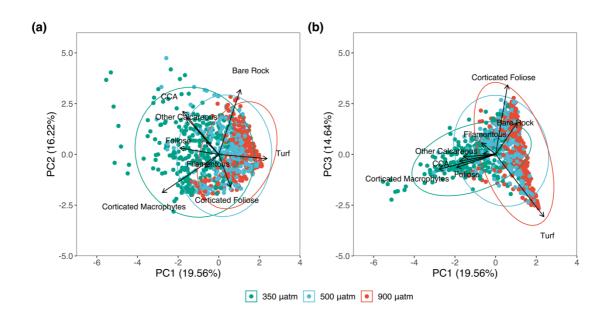


Figure 3. Community composition of algal functional groups between the three locations, 350
µatm CO₂ (Green), 500 µatm CO₂ (Blue) and 900 µatm CO₂ (Red), as assessed by principal
component analysis for (a) PC1 vs. PC2 and (b) PC1 vs. PC3. See Figure S2 for comparisons
of the locations at each month.

728

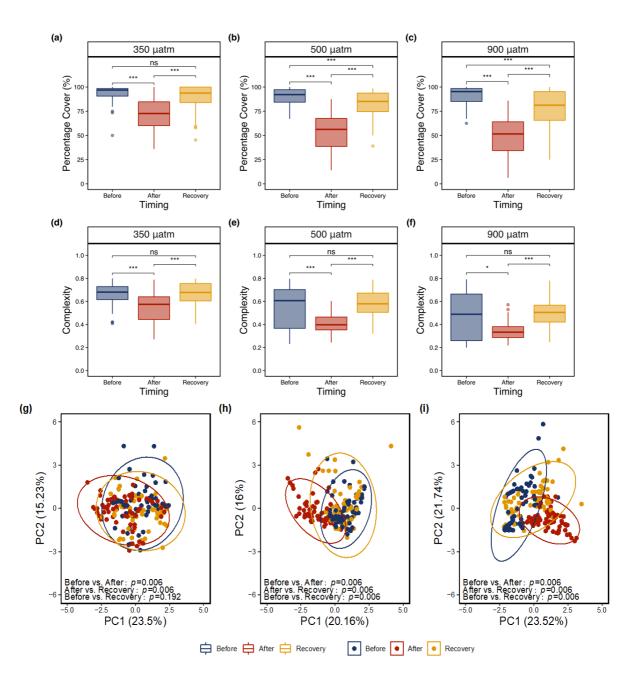
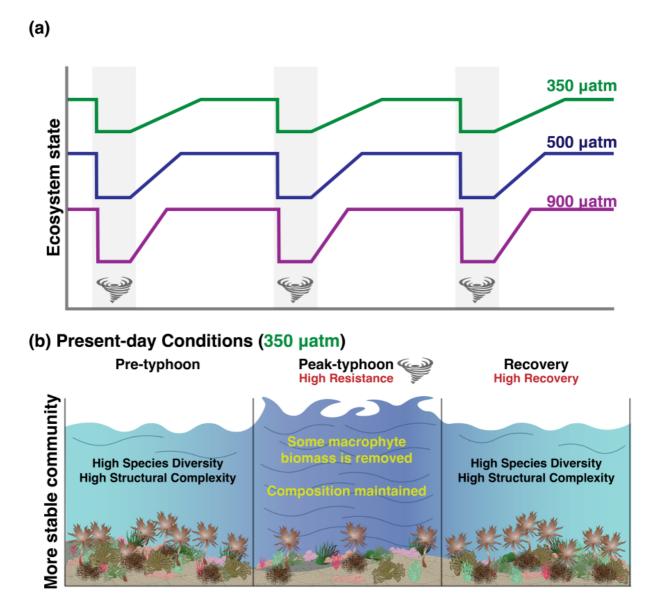
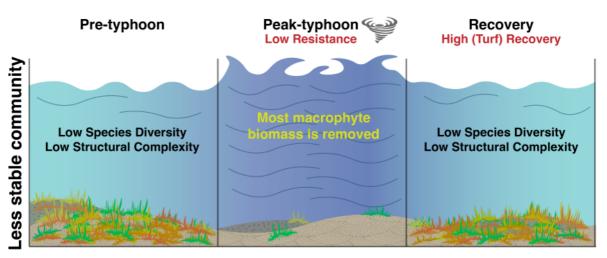


Figure 4. Percentage cover (a-c), complexity (d-f) and composition (g-i) of communities 'Before' (Blue) and 'After' the typhoon season peak (Red), and the subsequent 'Recovery' period (Orange). The 350 µatm location is shown on the left (a,d,g), the 500 µatm location in the center (b,e,h), and the 900 µatm location on the right (c,f,i). NOTES: Pairwise comparisons within Panels a-f are carried out by Type III ANOVA following GLMM (ns, p > 0.05; * p < 0.05; ** p < 0.01;*** p < 0.001) with the individuals points in a-f indicating outliers. To see the Panels a-f separated by year, see Figure S3, and for more detailed statistics, see Table S4 and S5.



(c) Future Conditions (500 µatm, 900 µatm)



730 Figure 5. Schematic summary of the differences in community stability between Present-day 731 (350 μ atm = Green) and future (900 μ atm = Red) CO₂ conditions in response to physical 732 disturbance from typhoons. The ecosystem state of present-day communities is higher than those under future pCO_2 conditions due to persistent press disturbance from ocean acidification 733 734 (leading to community compositional changes, lower functional diversity, and structural 735 complexity). Through time, ecosystem state fluctuates with seasonal and environmental change 736 (temperature, nutrients, disturbance), with larger fluctuations for acidified communities. 737 Periods of intense 'pulse' disturbance from seasonal factors such as typhoons have significant effects on ecosystem state. Differences in the functional traits and life-history strategies of the 738 739 species forming present-day (macroalgal-dominated) and future (turf-dominated) communities 740 lead to reductions in ecosystem stability. Macroalgal assemblages are resistant to disturbance 741 from typhoons, maintaining a more consistent composition between seasons and recovering to pre-disturbance baselines following physical disturbance. Turf-dominated communities 742 743 display low resistance to typhoons, which almost entirely removes the algal community, reducing the benthos to bare substrate. Turf resilience, promoted by boosted primary 744 production and growth under elevated pCO_2 conditions, meant that the empty space created 745 746 was later recolonised by turf algae. In the absence of physical disturbance feedback mechanisms maintain turf-dominance and inhibit macroalgal recruitment (see Harvey et al. 747 748 2021). Regular disturbance stabilised this pattern, locking the community into an early 749 successional stage.