2022-01-14

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Barnett, M

10.3389/fmars.2021.733799
Frontiers in Marine Science
Frontiers Media

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Total water column analysis shows the importance of a single species in subsurface chlorophyll maximum thin layers in stratified waters

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Submitted to Journal:
Frontiers in Marine Science

Specialty Section:
Marine Biology

Article type:
Original Research Article

Manuscript ID:
733799

Received on:
30 Jun 2021

Revised on:
01 Dec 2021

Journal website link:
www.frontiersin.org

Conflict of interest statement

The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest

Author contribution statement

All authors participated in the field survey, where MB collected all water samples, ANS developed and deployed the digital in-line holographic camera system and pre-processed images, and AESK and DAP aided with sampling procedures. MB processed the CTD and holocam data and analysed all samples. The manuscript was written by MB and AESK with all authors providing comments.

Keywords

Shelf seas, In situ observations, Holography, primary production, Phytoplankton, Dinoflagellate, subsurface chlorophyll maximum, thermocline

Abstract

Word count: 344

Marine phytoplankton form the base of marine food webs and are the driving force of the marine carbon cycle, so understanding the dynamics of their blooms is critical. While near-surface marine productivity (<10 m water depths) is extensively documented, that of the subsurface is less well characterised. Increasing evidence of the importance of subsurface chlorophyll maxima (SCM) and climatically driven increases in stratification of surface waters that promote SCM development call for improved sampling of the subsurface. To address this, we targeted the summer stratified waters of the Western English Channel, part of the NW European shelf seas, where SCM are commonly developed. In situ holography was applied to undertake the highest ever resolution, total water column, quantitative analysis of microplankton distribution, and demonstrated the importance of a SCM, collocated with the thermocline, dominated by a single species, the dinoflagellate Ceratium fusus. This species was dominant in the SCM over a wide area of the NW European shelf in the June/ July 2015 study period and comprised up to 85% of the SCM biomass. Analysis of similarity and multivariate non-metric multidimensional scaling showed the phytoplankton community of the SCM to be statistically distinct from those of the surface and deep waters. Holography also revealed a fine scale layering of taxa at different levels within the SCM, likely reflecting ecological differences. Some taxa followed the peak abundance of C. fusus, while others reached maximum abundances immediately below or above the C. fusus maximum, suggesting the possible operation of exclusion mechanisms. Additionally, the detection of abundant aggregates located only within and beneath the SCM demonstrates the potential importance of this deep production for the export of carbon to the sea floor. Some predictions of phytoplankton productivity propose a shift to smaller cells in the more stratified oceans of the future resulting in declining production and export. Results presented here, however, contribute to a growing body of evidence that suggests, on the contrary, that key species among the larger celled/ colonial, SCM-adapted diatoms and dinoflagellates may instead be selected in stratified conditions, driving increased production and export.

Contribution to the field

Marine phytoplankton undertake around half of global primary production and as they sink from the surface waters, they drive the marine biological carbon pump, drawing down CO2 from the atmosphere leading to carbon sequestration in the deep sea. Understanding phytoplankton ecology and the dynamics of their blooms is therefore critical to predictions of how climate change may affect the global carbon cycle. While the phytoplankton and biological productivity of the top few metres of the ocean have been well observed by satellite sensors and sampled, the subsurface is not well characterised. In this study we use innovative methods to produce the first total water column survey of phytoplankton in a shelf sea setting at the highest ever depth resolution and show the dominance of a single species in subsurface concentrations. We show how the subsurface taxa can be sampled and demonstrate their potential importance for production and carbon export. Climate change is driving increased stratification of the ocean and current models predict that this will lead to reduced phytoplankton productivity and weaken the biological pump. On the contrary, our results suggest that stratification may drive increased productivity with a shift to subsurface phytoplankton concentrations.

Funding statement

This research was supported by a University of Southampton Graduate School of the National Oceanography Centre Southampton (GSNOCS) PhD studentship to MB, with additional support from the Natural Environment Research Council (UK) "CaNDyFloSS: Carbon and Nutrient Dynamics and Fluxes over Shelf Systems" grant NE/K00185X/1 to DAP.

Ethics statements

Studies involving animal subjects
Generated Statement: No animal studies are presented in this manuscript.

Studies involving human subjects
Generated Statement: No human studies are presented in this manuscript.

Inclusion of identifiable human data
Generated Statement: No potentially identifiable human images or data is presented in this study.

Data availability statement

Generated Statement: The datasets presented in this study can be found in online repositories. The names of the repository/repositories and accession number(s) can be found in the article/supplementary material.

Total water column analysis shows the importance of a single species in subsurface chlorophyll maximum thin layers in stratified waters

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Keywords: shelf seas, *in situ* observations, holography, primary production, phytoplankton, dinoflagellate, subsurface chlorophyll maximum, thermocline

Abstract:
Marine phytoplankton form the base of marine food webs and are the driving force of the marine carbon cycle, so understanding the dynamics of their blooms is critical. While near-surface marine productivity (<10 m water depths) is extensively documented, that of the subsurface is less well characterised. Increasing evidence of the importance of subsurface chlorophyll maxima (SCM) and climatically driven increases in stratification of surface waters that promote SCM development call for improved sampling of the subsurface. To address this, we targeted the summer stratified waters of the Western English Channel, part of the NW European shelf seas, where SCM are commonly developed. *In situ* holography was applied to undertake the highest ever resolution, total water column, quantitative analysis of microplankton distribution, and demonstrated the importance of a SCM, collocated with the thermocline, dominated by a single species, the dinoflagellate *Ceratium fusus*. This species was dominant in the SCM over a wide area of the NW European shelf in the June/ July 2015 study period and comprised up to 85% of the SCM biomass. Analysis of similarity and multivariate non-metric multidimensional scaling showed the phytoplankton community of the SCM to be statistically distinct from those of the surface and deep waters. Holography also revealed a fine scale layering of taxa at different levels within the SCM, likely reflecting ecological differences. Some taxa followed the peak abundance of *C. fusus*, while others reached maximum abundances immediately below or above the *C. fusus* maximum, suggesting the possible operation of exclusion mechanisms. Additionally, the detection of abundant aggregates located only within and beneath the SCM demonstrates the potential importance of this deep production for the export of carbon to the sea floor. Some predictions of phytoplankton productivity propose a shift to smaller cells in the more stratified oceans of the future resulting in declining production and export. Results presented here, however, contribute to a growing body of evidence that suggests, on the contrary, that key species among the larger celled/ colonial, SCM-adapted diatoms and dinoflagellates may instead be selected in stratified conditions, driving increased production and export.

1 Introduction

Understanding phytoplankton ecology and the dynamics of their blooms is critical to predictions of how climate change may influence marine ecosystems and affect the global carbon cycle. The surface ocean has been scanned by satellite chlorophyll sensors for 40 years, with global coverage for the last 20 years (Hostetler et al., 2018). For ocean time series and transects, regular sampling has also largely targeted the surface < 10 m (Richardson et al., 2006) or else been restricted to discrete depth intervals spaced at several to 10s of metres, where typically pigments including chlorophylls have been used as a proxy for phytoplankton communities and productivity (Steinberg et al., 2001). Consequently, the phytoplankton that dominate the surface spring blooms or in upwelling regions are relatively well characterised (Behrenfeld and Boss, 2014).

In contrast to the surface waters, the subsurface ocean, beyond the range of satellite sensors (typically > 10 – 20 m), has only begun to be continuously surveyed in the past 15 - 20 years by chlorophyll sensors deployed on towed undulators, autonomous underwater vehicles (AUVs) and gliders. These surveys are now documenting great "patchiness" in the subsurface and are leading to an increased awareness of the significance of subsurface chlorophyll maxima. Such subsurface chlorophyll maxima (SCMs) are now known to be seasonally recurrent and persistent summer features in settings ranging from the mid latitude shelves of the NE Pacific (Perry et al., 2008) or NW Europe (Barnett et al., 2019; Hickman et al., 2012; Weston et al., 2005) to the Arctic (Churnside et al., 2020; Martin et al., 2010), and constitute biomass maxima. In the NW European Shelf, SCM often dominate summer new production with a total annual contribution estimated to be of the same order as (Hickman et al.,2012; Fernand et al., 2013; Williams et al., 2013) or greater than (Richardson et al., 2000) that of the spring bloom. Intensive undulator and sampling surveys such as the pioneering LOCO project in Monterey Bay (Sullivan et al., 2010; Rines et al., 2010) have revealed the widespread presence of SCM thin layers in a wide latitudinal range of coastal and shelf settings (Durham and Stocker, 2012; Greer et al., 2020). Such SCM thin layers are commonly linked to zooplankton distribution (Greer et al., 2013), and in some cases also drive multiple trophic levels via "bottom-up forcing" including top predator distribution (Benoit-Bird and McManus, 2012; Scott et al., 2010). While the dominant species of the surface ocean are well known, the phytoplankton of the subsurface including those of the SCMs are under-sampled. Conventional bottle surveys that sample discrete, multi-meter spaced depth intervals are not geared to target subsurface variability such as the thin layer SCMs that occur on sub-meter scales. Furthermore, comparison between net and bottle samples demonstrate that bottles do not adequately sample the largest phytoplankton (Armand et al., 2008), and it is some of the larger species that may be particularly significant for subsurface production and export (Kemp et al., 2000; Kemp and Villareal, 2013, 2018; Queguiner, 2013). Methods with the appropriate high levels of resolution are therefore required to identify subsurface phytoplankton taxa.

To characterise the phytoplankton of the subsurface it is necessary to deploy systems that sample the entire euphotic zone, not just the first optical depth available to satellites. *In situ* imaging using conventional optical techniques lacks the resolution to capture the nano- and microplankton, while flow cytometry is limited by the input nozzle diameter and may disrupt delicate phytoplankton colonies and aggregates, although developments in imaging flow cytometry can provide insights into temporal changes of phytoplankton community structure (Lombard et al., 2019; Olson et al., 2017). The recent development of commercial *in situ* laser holography systems allows for non-intrusive, 3D reconstructions of plankton over larger volumes and at higher resolutions than other instruments.

While *in situ* laser holography is recognised as having great potential, the scope of studies to date have been limited (Lombard et al., 2019; Nayak et al., 2021). A small number of species have been quantified at a limited number (2 to 3) of depth intervals (Malkiel et al., 1999; Talapatra et al., 2013). Higher resolution depth surveys have investigated total water column variation in phytoplankton versus zooplankton through tidal cycles in the Western English

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Channel (Cross et al., 2015) and the variability of four particle categories (phytoplankton, zooplankton, marine snow and others) in sub-pycnocline depths in the Ross Sea (Bochdansky et al., 2017). Simultaneous studies of holography and microstructure have been used to relate turbulent mixing to the distribution of colonial diatoms (Cross et al., 2014). Holographic studies have also identified subsurface thin layers of key elongate diatom species associated with pycnoclines and zones of low shear (McFarland et al., 2020; Nayak et al., 2018a; Nayak et al., 2018b).

The stratification of subsurface waters has the potential to generate a range of distinct niches for phytoplankton including those associated with the SCM (Cullen, 2015). The objectives of this study were to: 1) to identify the depth zonation of taxa within a stratified summer shelf sea water column; 2) to link depth-related characteristics of the phytoplankton community to water column physical structure; 3) to assess implications for phytoplankton ecology for the main taxa identified with special focus on the SCM. An additional aim was to demonstrate the full potential of *in situ* holography in providing very high resolution imagery of individual phytoplankton cells that make up the small scale layering within intense SCM where chlorophyll layers are <5m thick and referred to here as SCM thin layers - SCMTLs (Durham and Stocker, 2012).

2 Methods

2.1 Study Area and Sampling Procedure

The main study was conducted in the summer stratified waters of the Western English Channel between the 19th of June and 2nd of July 2015 (Fig. 1). A single deployment of a digital in-line holographic camera system, referred to as a holocam, was performed on 19th June from Research Vessel Falcon Spirit. The holocam was mounted on a profiling frame alongside an AML Oceanographic conductivity, temperature, depth (CTD) Plus V2 probe with chlorophyll-fluorescence sensor. A simultaneous and adjacent deployment from RV Callista used a SeaBird SBE19plus V2 CTD probe mounted with a WETlabs ECO FLNTU fluorometer (sensitivity: 0.025µg chl/l; fluorescence excitation/emission wavelengths: 470/695 nm). The configuration of the RV Callista CTD package allowed for slow descent/ascent rates without slowing sensor responses, thus improving dynamic accuracy and allowing small scale structure to be resolved. The CTD system was deployed at a descent/ascent rate of 0.01 - 0.1 m s^{-1} (rate slowed on approach to SCM), with a data acquisition rate of 2 Hz, which provided vertical resolution of 0.5 – 5 cm. Discrete water samples were taken from the surface waters (typically 0-20 m), SCM (at maximum chlorophyll), typically between 20-30 m, near the base of the thermocline, and deep waters extending beneath the SCM from depths of 30-35 m to the sea bed, see Supplementary Figure 2 for the all the complete depth profiles. (The terms deep and bottom are used interchangeably). Samples were collected using a Niskin rosette system (6 x 5 L Niskin bottles) mounted with the CTD package, and analysed for chlorophyll concentration and phytoplankton size and taxonomic community structure. The water sample from the SCM was also analysed for particulate organic carbon (POC). For the remainder of the field survey the RV Callista CTD Niskin rosette system was used to collect vertical water column profiles and discrete water samples (from the SCM, and from surface and deep waters where possible) from a further 39 sites in the Western English Channel showing seasonal summer thermal stratification with total water depths of between 52 and 86 m. (Fig. 1). Fourteen of these sites were in the same general location (repeat station 1) and the remainder of sites were located further afield, often sampled as part of a transect, either inshore-offshore or adjacent to the shore (Fig. 1). Water samples were routinely analysed for chlorophyll for the purpose of calibrating the fluorometer mounted with the CTD, and for phytoplankton community structure in order to provide context for the single holographic profile for the wider Western Channel over the survey period. Further afield, two stratified sites were profiled and sampled in the Celtic Sea on 28th July 2015, approximately 175 km from the study area in the Western English Channel and their locations are indicated by the two stars in the inset to Fig. 1.

2.2 Holocam Deployment and Data Processing

The holocam, updated from that described in Graham and Nimmo-Smith (2010) was mounted on a profiling frame alongside an AML CTD Plus V2 probe with chlorophyll-fluorescence sensor (Fig. 2), which was lowered slowly through the water column, with a sampling frequency of 15 Hz. The holocam was adapted for vertical profiling with a vertical configuration that minimised water column disruption, similar to that described in Graham et al. (2012). It was composed of a laser (658nm, 60 mW) and charge coupled device (CCD) digital camera, separated by low-profile extension tubes and 90° mirrors (82 mm apart) to distance the sample volume (between mirrors) from the CCD camera and laser, as illustrated in Fig. 2. The laser illuminated the sample volume and the CCD camera captured holograms comprising the interference patterns as particles diffracted the laser beam (Graham and Nimmo-Smith, 2010). Each holographic image had a size of 1536 x 1024 pixels, where pixel size was 4.65 µm and so each hologram's sample volume was 7.14 mm x 4.76 mm x 82 mm = 2.78 ml. During the holocam deployment at Site 1 on 29th June, a total of 3323 holograms were digitally recorded on the downcast from the surface to 50 m depth.

On return to the laboratory, the holograms were processed following methodology detailed by Graham and Nimmo-Smith (2010) and Davies et al. (2015). Briefly, the static background was subtracted from the imagery captured by the CCD camera and then each hologram was reconstructed computationally into a "stack" of slices in 3-dimensional space, focussed at 0.5mm intervals through the sample volume. Each stack was then cleaned by removing the lowest 1.1% intensity pixels to reduce background noise, with this threshold determined manually to provide the best balance between reducing noise and retaining the most weakly scattering particles (Davies et al., 2015). Subsequently each stack was analysed manually whilst viewed using a MATLAB graphical user interface that facilitated navigation through the sample volume. All particles, specifically microplankton (including phytoplankton as well as heterotrophic dinoflagellates and ciliates) and aggregates, were counted as they came into focus when stepping through the 3D stack at 0.5 mm intervals, thus > 100 slices were analysed per hologram. Due to the 4.65 µm pixel size of the holograms, any particles below this threshold were not identified. Phytoplankton identification was based on linking the digitally-reconstructed holographic interference patterns to optical microscopy of the bottle samples (Fig. 3 & Supplementary Figure 1), so that all major phytoplankton > 20-30 µm (including chains of smaller cells) were counted. In the case of the rod-shaped diatoms that were thought likely to be *Proboscia alata*, the very thin diameter precluded identification of individual cells in holograms so that counts for this taxon should be regarded as minimum values.

For the whole water column analysis, holograms were analysed at 0.3 m intervals. Where there was less rapid variation in abundance, as evidenced by change in chlorophyll fluorescence, in the deep waters beneath, and the surface waters above the SCM, particle counts were averaged at a 1 m depth resolution. Through the SCM, where the vertical changes in chlorophyll fluorescence were most rapid, a higher vertical resolution of 0.2 m was used without averaging. The choice of resolution was a compromise between identified vertical rates of change in taxa through the water column, as evidenced by rapid change in chlorophyll fluorescence, and how many holograms it was feasible to analyse in the given time.

2.3 Determination of Chlorophyll Concentration

Samples for chlorophyll analysis were collected by filtering 50 ml of water sample through 25 mm Whatman GF/F filters (in triplicate) immediately after collection Parsons et al. (1984). These filters were stored in a -20 °C freezer until analysis, which was conducted as soon as possible on return to the lab to avoid error associated with pigment degradation (Graff and Rynearson, 2011). Chlorophyll was extracted in 90 % acetone via sonication and then chlorophyll concentration was determined using a Turner Designs 10AU fluorometer based on the method of Welschmeyer (1994), whereby the fluorometer excited the extracted sample with blue light (436 nm) and the subsequent red fluorescence emission (680 nm) was recorded. For each SCM a chlorophyll intensity ratio was

calculated, being the ratio of peak chlorophyll concentration to background chlorophyll concentration. Fluorometerically derived values of chlorophyll were used to calibrate the fluorometer mounted with the SeaBird CTD and to derive the carbon: chlorophyll ratio for the SCMs.

2.4 Determination of POC

A water sample of 0.7 L was filtered onto a 25 mm pre-combusted (450 °C, 6hrs) Whatman GF/F filter under low vacuum (< 200 mg Hg) and stored at -20 °C until analysis. Prior to analysis the filter was dried in a 40 °C oven for a minimum of 12 hours, acid-fumed using 35% hydrochloric acid for 24 hours to remove inorganic carbon and then dried again. The sample was then analysed using a carbon, hydrogen, nitrogen, sulphur-oxygen (CHNS-O) elemental analyser (Carlo-Erba Instruments EA1108) (Collos, 2002).

2.5 Microscope Phytoplankton Analysis and Biomass Determination

Samples for phytoplankton community taxonomic structure analysis were collected by decanting 50 ml of water sample into a darkened glass bottle and immediately preserving with Lugol's iodine to a final concentration of 1 %. These Lugol's iodine preserved samples were later counted by settling 10 ml in a sedimentation chamber for 24 hours and examining using a Brunel SP951 inverted trinocular light microscope (Utermöhl, 1958). A single counting unit was an individual cell, whether solitary or part of a chain. Following previous practise (Olenina et al., 2006; Widdicombe et al., 2010), heterotrophic dinoflagellates and ciliates were included in the counts. Numerically dominant taxa (> 50 cells per ml) were counted along a single central traverse of the chamber base at 100x or 250x magnification depending on cell size. Cryptophytes (> 8 µm) and unidentified small naked dinoflagellates (10 – 20 µm and 20 – 25 µm) were also counted along a single traverse of the chamber at 250x magnification. All other cells ≥ 10 µm were counted at 100x magnification upon examination of the entire chamber base plate. Since most nano-phytoplankton < 10 µm and all pico-phytoplankton could not be identified with optical microscopy, CytoSense flow cytometry was used to analyse the contribution of these phytoplankton and provide a complementary measure of biomass in the main different size fractions (meso-, micro-, nano-, pico-). A complete list of identified phytoplankton taxa is presented in Supplementary Table 2.

Diatoms, dinoflagellates and flagellates were identified to species or at least genus wherever possible, and ciliates were grouped according to size and with reference to cell wall structure (loricate or aloricate). Where there was substantial size variation within a diatom or dinoflagellate genus, cells were also classified into size categories. These included *Pleurosigma* (small: ~50 µm length, medium: 80 – 170 µm length, large: 170 - 200 µm length), *Thalassiosira* (xsmall: < 10 µm height, small: 10 - 25 µm height, medium: 25 – 45 µm height, large: > 45 µm height), *Protoperidinium* (small: 10 - 30 µm diameter, medium: 30 – 65 µm diameter, large: 65 - 120 µm diameter) and *Rhizosolenia* (small: ≤ 10 µm diameter, medium: 10 – 20 µm diameter, large: > 20 µm diameter). Any remaining diatoms whose species or genus could not be differentiated accurately with optical microscopy were grouped as pennate or centric according to size (small: 20 – 40 µm length, medium: 40 – 65 µm length, large: 65 – 110 µm length, xlarge: 110 – 175 µm length, and small: 20 – 30 µm diameter, medium: 30 – 50 µm diameter, large: 60 – 150 µm diameter, xlarge > 150 µm diameter respectively). Similarly, any remaining unidentified dinoflagellates were also grouped according to size and with reference to cell wall structure (naked or armoured) where necessary (e.g. 10 – 20 µm naked dinoflagellates, 20 – 25 µm naked dinoflagellates, 10 – 30 µm armoured dinoflagellates).

Cell biovolume was calculated based on the geometric shapes and formulae assigned for each taxon by Olenina et al. (2006). Dimensions of at least 30 cells per taxon (only less in cases of rarely occurring taxa) were measured with the open source software 'ImageJ'. Cell carbon concentrations were estimated using the carbon – biovolume relationships of Menden-Deuer and Lessard (2000).

2.6 CytoSense Flow Cytometric Phytoplankton Analysis

Samples for phytoplankton analysis by flow cytometry were collected by fixing 10 ml of water sample with glutaraldehyde (0.25 % final concentration) and freezing at - 80 °C to preserve chlorophyll fluorescence immediately on return to lab (Fragoso et al., 2019; Marie et al., 2005). Samples were analysed with a CytoBuoy CytoSense flow cytometer and CytoUSB v5.7.5.7 data acquisition software, using two different sets of data acquisition settings; one optimal for larger phytoplankton (meso- and micro- phytoplankton: > 20 µm, and nano-phytoplankton: 2 – 20 µm), and the other optimal for small phytoplankton (pico-phytoplankton: < 2 µm). Meso-, micro- and nano-phytoplankton data were collected using a red (chlorophyll) fluorescence (RFL) trigger (30 mV) at a flow rate of 10 µl s^{-1} for 150 seconds or 10000 cells. Pico-phytoplankton data were acquired using a sideways scatter (SWS) trigger (25 mV) at a flow rate of 0.1 µl s^{-1} for 10000 cells, and pico-particles with a red fluorescence signal < 10 mV were manually removed from the dataset. Cell size derived from forwards scatter (FWS) was calibrated using a set of Thermo Fisher Scientific nonfluorescent polystyrene microspheres with a range of diameters (1, 2, 6, 10, 15 µm).

During data acquisition the CytoSense instrument recorded particle pulse shapes of FWS enabling description of phytoplankton community size structure using CytoClus v4.3.1.1 data processing software. For each sample a cytogram of total FWS (TFWS) and total RFL (TRFL) was generated to identify cell size. Thus clusters of pico-phytoplankton, nano-phytoplankton, and micro- and meso- phytoplankton could be resolved. As TRFL was calculated for each cell, the TRFL of the entire phytoplankton population and of each phytoplankton cluster could be determined. This allowed for the contribution of micro- and meso-phytoplankton (> 20 µm), nanophytoplankton (2 - 20 µm) and pico-phytoplankton (< 2 µm), to total community red fluorescence (TCRFL), a proxy for chlorophyll (and by extension biomass) to be assessed.

At the two Celtic Sea sites (Fig. 1), samples were collected from the SCM and preserved as described above. These preserved samples were run in a CytoSense flow cytometer, and 'Image in flow' was enabled to capture photographs of the most dominant larger (> 30 µm) phytoplankton.

2.7 Statistical Analysis

A number of statistical similarity analyses were undertaken to establish potential similarities and distinctiveness of the phytoplankton community at the different water depths. The analyses were undertaken using the PRIMER v6 software (Clarke and Gorley, 2006). Statistical analysis was conducted on phytoplankton carbon biomass data since biomass is a more biogeochemically relevant property (Paasche, 1960), as it provides a more accurate representation of community structure than abundance when the community consists of taxa of a variety of different sizes. Biomass data were first standardised by dividing biomass values by the total biomass for a given sample, and then normalised by performing a square root transformation to allow each taxon to influence similarity between samples and not just the dominant taxa (e.g. *Ceratium fusus*). Bray-Curtis similarity was calculated within each pair of samples and a cluster analysis subsequently performed to explore similarity of community structure among samples. Samples were grouped by sampling depth, i.e. surface, SCM and deep, and a non-metric multidimensional scale (nMDS) plot was generated to visually display similarity between samples, where a stress level below 0.2 is considered to indicate the ordination to be an accurate representation of the similarity relationship (Zuur et al., 2007). An ANOSIM (Analysis of Similarity) was applied to determine if the three sample clusters (surface, SCM and deep) were statistically distinct from each other in terms of their phytoplankton community structure, and to determine the level of separation between each cluster (given by the global R value, where values close to 0 indicate no separation and values close to 1 indicate high separation). A SIMPER (Similarity Percentage Analysis) was performed to investigate community similarities

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within clusters and dissimilarities between clusters, and to identify contributions of each taxon to overall similarity within each cluster and dissimilarity between clusters.

3 Results

3.1. Site 1 (Holocam Site): Overall Water Column Structure and Chlorophyll Distribution

Site 1 showed a stepped thermocline structure where the lower (main) thermocline was located between 25-28 m and an upper thermocline, between 15-18 m (Figs 4, 5). An intense SCMTL was located at the base of the main thermocline with a maximum chlorophyll concentration of 28.0 µg l^{-1}, and a chlorophyll intensity ratio of 33.3 (ratio of maximum SCM chlorophyll concentration to background chlorophyll concentration). The standing stock of chlorophyll integrated to a depth of 63 m (depth of CTD profile) was 66 mg m^{-2} and the SCM (20-35 m) accounted for 42 mg m^{-2}, being responsible for approximately 64 % of water column chlorophyll. In surface (0 - 20 m) and bottom (30 - 70 m) waters, chlorophyll concentrations did not exceed 1 µg l^{-1}. Phytoplankton counts converted to biomass from the discrete samples from the surface (14 m), SCM (27 m) and bottom layer (42 m) showed distinct differences (Fig. 4): the surface was populated primarily by ciliates (36%), smaller heterotrophic and mixotrophic dinoflagellates (33%) and a minor diatom component (16%); the SCM was dominated by a single species, the dinoflagellate *Ceratium fusus* (85%); the bottom layer contained a more mixed community with larger dinoflagellates (20%) and diatoms (36%) in addition to the smaller dinoflagellates (25%).

3.2. Holocam Site: Whole-Water Column Holographic Phytoplankton Profiling

The most striking feature of the holocam profile was the high concentration and dominance of a single species within the SCM. From holocam counts at individual 0.2 m – spaced depths, the dinoflagellate *Ceratium fusus* comprised up to 81 % of cells identified with concentrations up to 137 cells ml^{-1} (Fig. 6). The discrete water samples taken at the same site for microscopic analysis at the peak SCM chlorophyll were consistent with the holocam results with *C. fusus* comprising 85% biomass in the SCM sample (Fig. 4). The three discrete samples revealed distinct communities within the water column, while the holocam showed the continuity of change, with differences in the vertical distribution of the main taxa and key transitions in community structure at different levels (Figs 5, 6).

Within the SCM, *Ceratium fusus* was consistently the most dominant species, ranging between 41 – 81 % of cells identified from the 0.2 m spaced holograms between 25 – 30 m (Fig. 6). Through the SCM, the *C. fusus* concentrations were > 20 cells ml^{-1} between 25.5 – 29 m, > 60 cells ml^{-1} between 27 – 28 m, with the peak of 137 cells ml^{-1} at 27.4 m (Fig. 6). Holograms analysed through the SCM also revealed that the *C. fusus* concentration (Fig. 6) matched closely the chlorophyll-fluorescence profile (Fig. 4). Both show a continuous downward decrease from the SCM peak concentration at 27.4 m, but above the SCM peak, a sharp decrease is succeeded upwards by a ~ 1.5 m zone of moderately elevated concentration, before *C. fusus* concentrations reduce to < 3 % of cells, by 25 m. In contrast, although the decline was more abrupt below the SCM peak, *C. fusus* contributions of around 22 % were sustained throughout the deeper waters (Fig. 5). The other key contributor to the SCM was the rhizosolenid diatom *Proboscia truncata* that attained peak concentrations around 1 m above and below the *C. fusus* peak (Fig. 6). Similar to *C. fusus*, *P. truncata* had sustained concentrations through the deep layer, below the SCM (> 30 m), comprising typically between 10 – 32 % of cells identified from the individual spaced holograms. The dinoflagellate, *Ceratium lineatum* also made a significant contribution to the community within the lowermost part of the SCM, being 6.3 – 8.1 % of cells identified between 30 – 33 m. Intriguingly, the abundance of the silicoflagellate, *Dictyocha* sp. (likely *Dictyocha fibula* showing the characteristic 4 radial spines – Supplementary Fig. 1) closely resembled that of *C. fusus*, with peak concentrations in the SCM coinciding with those of *C. fusus* between 27-28 m. However, above the

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most intense part of the SCM (25.5 – 29 m) *Dictyocha fibula* replaced *C. fusus* as the dominant taxa, contributing 18.0 – 26.9 % of cells identified between 23 – 25 m at a concentration of 4.3 – 6.5 cells ml^{-1}. *D. fibula* occurred in reduced numbers below 30 m (generally ≤ 1.7 cells ml^{-1}), such that it was generally less than 10 % of cells identified. Significantly, peak concentrations of the heterotrophic dinoflagellates *Ceratoperidinium* and *Gyrodinium* spp. occurred immediately above (27 m) and beneath (28.5) the *C. fusus* SCM peak, but decreased where *C. fusus* was most abundant (Fig. 6).

Above the SCM, in the surface waters, the community was dominated by ciliates (10 – 58 % of cells identified; up to 6.0 cells ml^{-1}), and *Ceratoperidinium* and *Gyrodinium* spp. (generally 10 – 30% of cells; between 1 - 4.0 cells ml^{-1}) (Figs. 4, 5). The larger phytoplankton were particularly sparse above the upper step to the thermocline. The only significant large phytoplankton in the topmost layer were rod shaped diatoms, primarily the rhizosolenid *Proboscia alata* that attained concentrations up to 4.0 cells ml^{-1} and constituted 30 % of cells identified above 15 m (Fig. 5). The distribution of other phytoplankton broadly followed the double thermocline structure with curled chained diatoms (mainly *Chaetoceros* spp.), chained centric diatoms (*Thalassiosira* spp.) and *P. truncata*, the dinoflagellates *C. fusus* and *C. lineatum*, and the silicoflagellate, *D. fibula*, making minor contributions from the top of the SCM up to 15 m, but being rare or absent above this depth (Fig. 5).

The most distinctive feature of the waters below the SCM was the presence of abundant aggregates that were entirely absent from the surface water. These were recorded from the holograms in a number of size ranges covering an order of magnitude variation with the largest > 1 mm (Fig. 5, Supplementary Fig. 1). Most of the material in the aggregates was not identifiable although fragments of diatom chains or large dinoflagellates could sometimes be discerned. Beneath the SCM, *C. fusus* and *P. truncata* as well as *C. lineatum* and *D. fibula* were all persistent (Fig. 4). There was also a significant presence of *P. alata* but with a dominance of shorter (< 450 µm length) specimens. A significant difference was the major presence of chained centric diatoms, likely *Thalassiosira* (10 – 37 % of cells identified; up to 10.4 cells ml^{-1}) in contrast to their near absence above 15 m and a sparse presence 15 – 27 m.

3.3 Holocam site: Discrete Samples: Size Fractionated Chlorophyll and Relation to Phytoplankton Community (Biomass) Structure

Within the SCM at site 1 the phytoplankton community was predominantly composed of larger cells, with micro- and meso-phytoplankton (> 20 µm) contributing 96 % of community chlorophyll, as determined by CytoSense flow cytometry (Table 1). As well as contributing 85 % of the SCM carbon biomass identified using light microscopy (Supplementary Table 2), *C. fusus* also contributed approximately 69 % of total chlorophyll. The chlorophyll concentration was determined by estimating *C. fusus* carbon content using the microscope cell count and biovolume estimate, and converting this to chlorophyll based on a carbon to chlorophyll ratio of 38 calculated using particulate organic carbon (POC) and chlorophyll data collected for the SCM. The other contributors to SCM biomass were the diatoms (mostly *P. truncata*), 3.3 %; other dinoflagellates, 7.1 %; flagellates (mainly *D. fibula*), 2.2 % and ciliates (aloricate and loricate) 2.5 %.

In bottom waters the phytoplankton community was mostly micro- /meso-phytoplankton, and nano-phytoplankton (55 % and 33 % of community chlorophyll respectively; Supplementary Table 2). Diatoms (37.2 %) and dinoflagellates (45.0 %) dominated community biomass identified using microscopy. Ciliates, predominantly aloricate, and flagellates, mainly *Chrysophaerella longispina* and *D. fibula*, also had significant biomass, contributing 8.4 % and 9.4 % respectively. Within the diatoms *P. truncata* (19.6 % of microscope community biomass) and *Chaetoceros* spp. (12.6 % of microscope community biomass) were most dominant, and within the dinoflagellates *C. fusus* (15.4 %), small naked dinoflagellates (17.2 %) and *C. lineatum* (3.6 %) made the largest contributions of biomass (Fig. 4; data in Supplementary Table 2).

3.4 The Wider Shelf Context: the Western Channel and Celtic Sea

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To provide context for the results at site 1 we sampled the SCM at a further 39 stratified sites in a combination of repeat stations and transects (Fig. 1) with representative profiles shown in Fig. 7. Results of flow cytometry are given in Table 1 and the range of biomass contributions of major groups is given in Table 2 with full biomass results in Supplementary Table 2. The sampled sites included 10 where the surface and bottom waters were also sampled. This revealed the widespread dominance of *C. fusus* within the SCM throughout the region. There were also distinct differences in the communities of the surface, SCM and deep layers throughout the survey period that were consistent with the observations at site 1. Size fractionated chlorophyll analysis revealed that the meso- and micro- phytoplankton (> 20 µm) contributed 75% ± 11.1 (n = 40) to the SCM chlorophyll, while the nano- and pico- phytoplankton (< 20 µm) contributed 73.8% ± 15.8 (n = 11) to the chlorophyll of the surface waters (Table 1). The proportion of *C. fusus* in the SCM biomass also varied, with the highest concentrations in the most intense SCM with the highest chlorophyll intensity ratios (Fig. 9). Across all the sites, there was a strong correlation between the abundance of *C. fusus* and chlorophyll concentration (Fig. 10).

A cluster analysis with ANOSIM using carbon biomass data identified the surface, SCM and deep samples to be statistically distinct in terms of their taxonomic community structure ($p = 0.001$), and a global R of 0.82 (R statistic from pairwise tests varied from 0.75 to 0.91) indicated these clusters were well separated. An nMDS analysis provided a 2D spatial representation of the separation between surface, SCM and deep samples based on their biomass values, and a stress level of 0.13 verifies that the 2d plot is a reliable representation of the multidimensional relationships (Fig. 11). Taxa whose cumulative contribution to similarity within a cluster and dissimilarity between clusters was approximately 90 % are given in Table 3.

Within the SCM (Fig. 7c; summarised in Table 2; full data in Supplementary Table 2) dinoflagellates were generally most dominant, and at over 90 % of sites, *C. fusus* was the most dominant dinoflagellate species. The SCM sample cluster had an average similarity of 67.3 %, the top five contributors being *C. fusus* (16.8 %), large aloricate ciliates (8.8 %), 10 – 20 µm naked dinoflagellates (5.5 %), *P. truncata* (5.1 %) and *D. fibula* (5.0 %) (Table 3).

Above the SCM in the surface waters (Fig. 7b; summarised in Table 2; full data in Supplementary Table 2), dinoflagellates, mostly small naked dinoflagellates and *Gyrodinium* + *Ceratoperidinium* spp., and ciliates (mostly aloricate) were quite consistently dominant. The surface sample cluster had an average similarity of 72.0 %, over 40 % of which was contributed by medium aloricate ciliates (11.9 %), large aloricate ciliates (9.7 %), 10 – 20 µm naked dinoflagellates (8.4 %), *Gyrodinium* spp. (5.8 %) and 20 – 25 µm naked dinoflagellates (5. 4 %) (Table 3).

Beneath the SCM in the bottom waters (Fig. 7a; summarised in Table 2; full data in Supplementary Table 2), diatoms, predominantly *P. truncata* and *Chaetoceros* spp., and dinoflagellates, mainly small naked dinoflagellates and *C. fusus*, were most dominant. The deep sample cluster had an average similarity of 67.5 %, and the top five contributors to this similarity (Table 3) were 10 – 20 µm naked dinoflagellates (12.1 %), *P. truncata* (11.7 %), *Chaetoceros* spp. (5.2 %), *Dictyocha* spp. (5.1 %) and large aloricate ciliates (4.8 %).

A broader context is provided by two stratified sites with prominent SCM that were profiled in the Celtic Sea later on 28th July 2015, approximately 175 km from the Western English Channel study area (Fig. 12). Sampling at these two sites indicated a high abundance and biomass dominance of *C. fusus* within the SCM located at the base of the thermocline at depths of 45 – 55 m (Fig. 12). Concentrations of *C. fusus*, with an approximate size range of 350 - 450 µm, were determined at 73 and 249 cells ml^{-1} at site 1 and 2 respectively. The CytoSense flow cytometer 'Image in flow' was used to record images *of C. fusus* (Fig. 12).

The cytograms produced by the cytosense flow cytometer in Fig. 12c show plots of sideways cell scatter (SWS) as an indicator of cell length vs. red (chlorophyll) fluorescence generated for samples collected from the SCM at each of these stations. The cytograms clearly highlight the size dominance of *C. fusus* in the samples collected from the SCM. These findings further point to the

widespread dominance of a single species within the UK continental shelf SCM in the 2015 summer season.

4 Discussion

In line with the objectives of this research we found a broad tripartite make-up of phytoplankton communities that was related to water column structure in the Western Channel in July 2015. Distinct phytoplankton communities occupied the surface waters, the thermocline and associated SCM, and the deep waters. Whereas a few taxa were found throughout the water column, many had more restricted vertical distribution. The highest variability in abundance occurred around the SCM with very rapid changes in the vertical distribution of taxa. This suggests complex interactions amongst taxa and specific adaptations to the SCM/ thermocline environment.

4.1 Dominance of a Single Dinoflagellate Species, *Ceratium fusus* in the SCM

The most remarkable feature was the dominance of the dinoflagellate *C. fusus* in the SCM, with its greatest abundance in the most chlorophyll-rich and sharpest SCM. A further notable aspect was the widespread distribution of *C. fusus* ranging to the SCMTLs of the Celtic Sea in addition to the Western Channel (Fig. 12). The presence of this species in the SCM over such wide distances suggests effective dispersal mechanisms over the UK continental shelf in July 2015. While *C. fusus* is a mixotroph, the size fractionated chlorophyll-fluorescence, and close match of *C. fusus* abundance variation to that of chlorophyll, show that it was actively growing through photosynthesis in autotrophic mode at site 1.

Within the NW European shelf seas high concentrations of *C. fusus* have not previously been reported. Continuous plankton recorder (CPR) surveys, sampling depths of 10 m over the past several decades, reported *C. fusus* as one of the 10 most abundant phytoplankton in the North Sea and as common in the Bay of Biscay/ Western Channel/ Celtic Sea (Beaugrand et al., 2000). More broadly, *C. fusus* is classified as a cosmopolitan species in the North Atlantic, found in a wide temperature range (2 – 29.5 °C) based on surface water surveys (Dodge and Marshall, 1994).

Relevant information on the physiology of *C. fusus* is provided by culture experiments on samples taken from the Sagami Bay area of Japan where subsurface maxima are frequently observed at depths of 5 m associated with pycnoclines (Baek et al., 2007). Laboratory experiments show that at 12 °C (similar to the thermocline temperatures in the Western Channel) specific growth rates decrease with increasing photon irradiance between 53 to 183 µmol m^{-2} s^{-1} (Baek et al., 2007). Light levels at the site 1 thermocline were around 34 µmol m^{-2} s^{-1}. Culture experiments also show that cells can survive at least 15 days darkness (Baek et al., 2008b). Taken together, these insights underscore the adaptations of *C. fusus* to the SCM niche.

The average swimming speed of *C. fusus* is 76 µm s^{-1} (27 cm/ hour) (Baek et al., 2009) (although an earlier study report speeds of up to 278 µm s^{-1} (Hasle, 1954)). Such speeds are sufficient to exceed the effects of the typical vertical eddy diffusivities of the shelf sea thermocline (Sharples et al., 2001) so that *C. fusus* would be capable of vertical movement to respond to and exploit environmental gradients of light and nutrients. Indeed, field and laboratory evidence indicate that *C. fusus* can migrate vertically to avoid strong sunlight (Baek et al., 2009). Thus, in periods of sustained stable stratification, in the absence of wind/ wave induced turbulence, *C. fusus* is well-equipped to optimise growth conditions at the thermocline by, for example, moving to exploit peak nutrient concentrations.

While the highest specific growth rates recorded for *C. fusus* are 0.59 d^{-1}, these are attained at 24 °C, whereas the growth rate at 12 °C are only 0.1 d^{-1} (Baek et al., 2008b), leading these authors to suggest introduction of populations by advection from warmer waters to explain high concentrations in colder waters. These relatively slow reported growth rates, may suggest

that there is a further mechanism responsible for the most intense *C. fusus* thin layers other than merely *in situ* growth. The most intense *C. fusus* SCMTL coincided with strong stratification, and a steep thermocline which would dampen turbulence and enhance vertical shear (Durham and Stocker, 2012) (see also associated higher buoyancy frequency values, Supplementary Table 1). This shear may exert a torque on plankton swimming at relatively low speeds (100 μm s^{-1}) leading to rotation and gyrotactic trapping that may enhance layer formation (Durham and Stocker, 2012). However, in the absence of data on shear in our study area, the role of such a mechanism cannot be assessed. A preference for strongly stratified conditions is also evidenced by experiments showing that *C. fusus* growth rates are unaffected by small scale turbulence but decreased up to 50 % with high turbulence (Sullivan and Swift, 2003).

C. fusus also has a number of other traits that may enhance its ability to thrive in the deeper SCM niche. In common with the rhizosolenid diatoms it has a high aspect ratio, so with long axis horizontal, an orientation promoted by shear, light absorption would be enhanced (Nayak et al., 2018a). The large cell size and spines also act as a defence against grazing and *Ceratium* spp. is avoided by all but the largest calanoid copepods (Nielsen, 1991). The ecology of *C. fusus* is less well known, but the close relative *Ceratium furca* can undertake luxury nutrient uptake and, if necessary, feeds via phagotrophy (Baek et al., 2008a), both strategies that would facilitate survival through nutrient poor conditions in the shelf sea thermocline, where supply of nutrients via tidal-driven turbulence is intermittent (Sharples et al., 2001).

4.2 Water Column Structure and Layering of the Plankton Community

The overall plankton community structure at site 1 followed the water column structure, while there was also smaller scale layering of taxa within the SCM (Figs. 5, 6). The surface waters were divided by a stepped thermocline with an "upper thermocline" above the main lowest thermocline (Fig. 4). Such features may develop when sustained wind mixing is succeeded by a prolonged calm period resulting in a single deep thermocline. This is then followed by a renewed windy period of lesser duration or magnitude that mixes the uppermost waters but not as far as the deep thermocline. This produces a new, shallower mixed layer and associated, upper thermocline (Beer, 1983). Alternatively, stepped thermoclines may also form in response to internal-wave-induced mixing (Navrotsky et al., 2004). The uppermost layer above the upper step of the thermocline was dominated by ciliates and heterotrophic dinoflagellates with the rhizosolenid *Proboscia alata* somewhat anomalous as the only significant large diatom present. Intriguingly, several species of rhizosolenid diatoms engage in vertical migration between deeper nutriclines and the surface through buoyancy regulation (Moore and Villareal, 1996), and although not demonstrated for *P. alata*, there are reports of positive buoyancy in this species (Villareal, written com, 2017). Such behaviour could explain the otherwise incongruous presence of *P. alata* in the surface. The major phytoplankton taxa were more uniformly present in the lower segment of the surface waters and there was an increase of key species downwards to the SCM (Figs. 5, 6).

The holocam revealed a fine-scale layering of taxa within the SCM at site 1 on a vertical scale hitherto undocumented, with some taxa reaching maximum abundances above and below, but declining within the *C. fusus* peak while others followed the *C. fusus* peak (Fig. 6). The silicoflagellate *Dictyocha fibula* had a sustained presence in the lower segment of the surface layer but peak abundances coincided with the *C. fusus* peak (Fig. 6). *D. fibula* is generally only reported in low abundances, for example, in the deeper part of the euphotic zone above a DCM found at 60 - 80 m in NW Mediterranean (Estrada et al., 1993). It has been observed in higher abundances (20 - 80% of cells observed) in subsurface (70 – 80 m) maxima in the Western Mediterranean, between SE Spain and Algeria, in close association with the nitracline (Lohrenz et al., 1988), but has not previously been reported to occur in a concentrated thin layer. The presence of maxima of the heterotrophic dinoflagellates *Gyrodinium* and *Ceratoperidinium* just above and beneath the *C. fusus* peak is in keeping with their requirements for relatively high prey concentrations for growth (Hansen, 1992) but their marked decrease within the peak itself may suggest some mechanism for

exclusion from the *C. fusus* maxima. This also highlights a different feeding strategy from their potential competitors, the ciliates, which did not increase in abundance within the SCM. Maximum abundances of the diatom *P. truncata* in the SCMTL may reflect similar concentration mechanisms to those of *C. fusus*. Peak abundances of *C. lineatum* occurred some 2-3 m below the *C. fusus* max. layer. *C. lineatum* has been observed to grow better than *C. fusus* at lower temperatures (Nordli, 1957), therefore temperatures of ~11.6 °C on the downslope of the SCM, compared to those on the upslope of the SCM down to the depth of maximal fluorescence (12.0 - 13.2 °C) may have been more favourable for *C. lineatum*.

4.3 Aggregates and Export from the SCM

A striking feature of the holocam results is the distribution of abundant aggregates that occurred solely within and beneath the SCM (Fig.5). While occasional diatom or dinoflagellate fragments may be observed in these, most contained indeterminate nano-sized particles and closely resemble the transparent exopolymer particles (TEP) imaged by Passow and others (Passow, 2002). The distribution of aggregates suggests that they were generated by activity within the SCM and represent a sustained flux of organic material from the SCM. The key phytoplankton of the SCM were also present throughout the lower layer, suggesting that they, also, were settling from the SCM. This suggests a key role of the SCM in generating production and export of organic matter and supports wider evidence for the importance of SCM production in shelf seas (Fernand et al., 2013; Hickman et al., 2012; Richardson et al., 2000; Williams et al., 2013).

4.4 Predictions of Future Enhanced Stratification and Implications for SCM Development

Increases in upper ocean temperatures and localised freshening, due to increased precipitation at high latitudes, are resulting in increasing surface water stratification in the global ocean and across shallow shelf seas (Bindoff et al., 2007). Future projections depict warming and freshening of the seas around the UK, with seasonal stratification projected to become more intense and persist longer leading to longer-lived and steeper thermoclines (Holt et al., 2012; Lowe et al., 2009). Previous work suggests that SCMTLs are more likely to form in stronger thermoclines that promote the growth and concentration of the larger diatoms and dinoflagellates, so that SCM may be of increased biogeochemical significance in future oceans (Barnett et al., 2019). Some predictions of phytoplankton productivity propose a shift from larger to smaller cells in the more stratified oceans of the future. This is predicted to result in declining production and export (Finkel et al., 2010). Our results suggest, on the contrary, that the larger celled, SCM-adapted phytoplankton may instead be selected in stratified conditions, driving increased production and biomass. The abundant aggregates that we identify below the SCM might also suggest that export may be significant in such scenarios. It is therefore important to improve our understanding of the ecology of the SCM environment.

4.5 Future Potential of Holocam Studies

This study highlights the spatial scale of analysis that is required to adequately understand phytoplankton ecology in stratified settings. The decimeter-scale variability of different taxa through the SCM may provide the basis for insights into the interactions between autotrophs and heterotrophs. Furthermore, even two closely related diatom species, the rhizosolenid diatoms, *P. alata* and *P. truncata* appear to be responding quite differently, with the former present in the surface layer (possibly through buoyancy regulation) and the latter in the SCM and bottom layer. This suggests that simple traits like cell size (Barton et al., 2013) are insufficient to parameterise biogeochemical models, but rather, the ecology of the dominant species needs to be better understood. The location of the SCMTL locked to the thermocline raises questions on the nature

of potential biological – physical interactions, which could be investigated, for example, by deployment of a free-fall turbulence microstructure profiler mounting a holocam.

With regard to the phytoplankton identification from holocam imagery, this was done entirely by visual identification through comparison with optical microscopy and was time consuming. There is clearly scope for use of image analysis techniques and the potential development of fully unsupervised (automated) methods with ongoing increases in computing capability. Whereas there has been significant progress in aspects of image processing including focusing, segmentation and sizing, and simple particle classification (Davies et al., 2015), the unsupervised classification of more complex particle populations remains a challenge. The present study highlights the importance of a relatively small number of key taxa, so a way of simplifying the task of automated particle recognition would be limiting the number of shapes to be targeted for identification (Davies et al., 2015), which could be done following a brief initial pilot study. Thereafter automated identification could also be improved by the use of neural networks/ machine learning (Deglint et al., 2019; Luo et al., 2018; Guo et al., 2021).

5 Conclusions

Surveys of the summer stratified waters of part of the NW European shelf seas in June/ July 2015 showed a broad tripartite structure with the surface and deep waters separated by a thermocline with a co-located subsurface chlorophyll maximum (SCM).

Analysis of similarity and multivariate non-metric multidimensional scaling showed the phytoplankton communities of the surface waters (above the thermocline), the SCM and the deep waters (beneath the thermocline) to be statistically distinct.

Throughout the survey area the SCM was dominated by a single species, the dinoflagellate *Ceratium fusus*.

The proportion of *C. fusus* in the SCM biomass varied, with the highest concentrations of up to 85% of the SCM biomass and 69% of SCM chlorophyll in the most intense SCM with the highest chlorophyll intensity ratios.

The holocam results revealed a fine scale layering of taxa within the SCM with some taxa following the peak abundance of *C. fusus* but with others reaching maximum abundances immediately above and below it suggesting the possible operation of exclusion mechanisms.

The abundances of certain diatoms including *Proboscia alata* (most abundant in surface waters) and *Proboscia truncata* (most abundant in the SCM and bottom waters) highlights the known adaptation to stratified waters of rhizosolenid diatoms and may indicate buoyancy regulation.

Abundant aggregates present beneath the SCM throughout the bottom waters suggest mass settling and export from the SCM.

6 Conflict of Interest

The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

7 Author Contributions

All authors participated in the field survey, where MB collected all water samples, ANS developed and deployed the digital in-line holographic camera system and pre-processed images, and AESK and DAP aided with sampling procedures. MB processed the CTD and holocam data and analysed all samples. The manuscript was written by MB and AESK with all authors providing comments.

8 Funding

This research was supported by a University of Southampton Graduate School of the National Oceanography Centre Southampton (GSNOCS) PhD studentship to MB, with additional support from the Natural Environment Research Council (UK) "CaNDyFloSS: Carbon and Nutrient Dynamics and Fluxes over Shelf Systems" grant NE/K00185X/1 to DAP.

9 Acknowledgements

We thank the crew of the RV Callista (University of Southampton) and the RV Falcon Spirit (University of Plymouth) for their assistance in data collection. We are grateful to Mark Moore (University of Southampton) for collecting the Celtic Sea samples that were analysed using the CytoSense flow cytometer.

Figure Legends

Fig. 1. Study area in the Western English Channel where sampling occurred between the 19th of June and the 2nd of July 2015. Inset shows location with reference to the NW European continental shelf and location of two stations sampled in the Celtic Sea later in July. Main figure: the encircled blue star indicates the location of Falmouth. Symbols indicate the 40 stratified sites profiled and sampled (symbols labelled with site numbers), and on which date (represented by different symbols). The red/orange encircled cross indicates where the holocam was deployed on the 19th June. Repeat station 1 is indicated, including sites sampled there.

Fig. 2. Photograph of the holocam system components mounted on a profiling frame with the AML CTD Plus V2 probe and fluorometer. The sample volume is indicated by the yellow arrow between low-profile housing extensions containing the laser and CCD camera. Raw data was collected by the on-board data logger.

Fig. 3. Selected comparative images of key taxa from optical microscope (left panel) and holocam (right panel). (A) the dinoflagellate, *Ceratium fusus*, (B) the dinoflagellate, *Ceratium macroceros*, (C) the diatom, *Proboscia truncata*, (D) the chained diatom, *Thalassiosira*, (E) the silicoflagellate, *Dictyocha* sp. These were used as a basis for the holocam taxa counts. A more comprehensive set of comparative images is provided in Supplementary Figure 1.

Fig. 4. Temperature and chlorophyll profile, and phytoplankton (biomass) community structure determined by microscopy at site 1 where the holocam was deployed. The green line represents chlorophyll concentration determined from CTD chlorophyll-fluorescence, the blue dashed line represents temperature and the red Xs (and corresponding red arrows) where water samples were collected for phytoplankton analysis. An optical microscope image of the community within the SCM clearly shows the dinoflagellate *Ceratium fusus* to be dominant.

Fig. 5. Holocam determined abundances of the 10 most numerically dominant large (> 30 μm) phytoplankton for the whole water column. Holograms were analysed at 0.3 m depth intervals, increased to 0.2 m intervals for the SCM (between 25 – 30 m). For the whole water column analysis, particle counts were averaged over each meter, where circle size is proportional to cell abundance standardised to maximum cell abundance, and the centres of the circles correspond to the the depth analysed. Aggregates counted are shown in the right hand panel of the plot, where circle size is proportional to the number of aggregates counted in the sample volume analysed, and the centres of the circles, again, correspond to the depth analysed. The AML CTD Plus V2 probe temperature (dashed blue line) and chlorophyll-fluorescence (green line) depth profile of the holocam profiled site is shown in the left hand panel of the plot, red Xs indicate where discrete water samples were collected.

Fig. 6. Holocam determined abundances of the dominant *C. fusus* and other major plankton through the SCM showing layering of key taxa. Analysis through the main SCM was at a higher resolution (20 cm).

Fig. 7. Representative profiles of temperature and chlorophyll fluorescence from the survey sites shown in Fig. 1. A full reproduction of all the survey profiles is given in Supplementary Fig. 2.

Fig. 8. Phytoplankton community structure within (A) deep waters (11 sites), (B) surface waters (11 sites) and (C) the chlorophyll peak (all 40 sites) at sites profiled and sampled for optical microscope microplankton analysis. Percentage contribution of diatom taxa, dinoflagellate taxa, flagellates, ciliates and non-flagellated chlorophyceae to community carbon biomass as identified by microscopy, where diatom taxa are indicated by blue colouration and dinoflagellate taxa by

red colouration (NB small naked dinoflagellates refer to 10 – 25 µm naked dinoflagellates that were not identified to genus/species). On plot (C) dates of sampling are given, sites sampled at repeat station 1 are labelled R1, and sites sampled as part of a transect are indicated by an arrow (two arrow head ends indicates an across shore transect and a single arrow head indicates an inshore-offshore transect, where the direction of the arrow indicates movement inshore to offshore).

Fig. 9. Biomass of *Ceratium fusus* plotted against chlorophyll intensity ratio (ratio of maximum SCM chlorophyll concentration to background chlorophyll concentration) for the SCMs surveyed (see Fig. 6).

Fig. 10. Relationship between cell carbon of *C. fusus* derived from cell counts and chlorophyll concentrations from the SCM in all the surveyed sites $y=38.4x-96.9$; $r^2 = 0.84$.

Fig. 11. Non-metric multidimensional scaling (nMDS) plot representing the similarity in phytoplankton community taxonomic structure and showing the separation between SCM (blue triangles), surface (red circles) and bottom water (green squares) samples, based on carbon biomass values. The 2D stress value of 0.13 indicates that the plot represents a good representation of the multidimensional relationships.

Fig. 12.
Results from the two sites sampled in the Celtic Sea in July 2015 with locations indicated by the numbered purple circles (1 and 2) in panel (A) , (blue star is Falmouth and red/orange cross is the holocam deployment site). Panel (B) shows the temperature and chlorophyll profiles of these two sites exhibiting an SCM. (Green line – chlorophyll; blue dashed line – temperature). CytoSense 'Image in flow' from the SCM samples reveal abundant *C. fusus*. Panel (C) shows two cytograms generated by the Cytosense flow cytometer comparing sideways cell scatter (SWS) length vs. red (chlorophyll) fluorescence, with *Ceratium fusus* cell clusters indicated in red symbols.

Tables

Table 1. Contribution of micro- and meso-phytoplankton (> 20 µm), nano-phytoplankton (2 – 20 µm) and pico-phytoplankton (< 2 µm) to total community red fluorescence (TCRFL; a proxy for chlorophyll and by extension biomass) as identified using CytoSense flow cytometry.

Table 1

Date	Site no.	Sample location	micro- and meso- (% of TCRFL)	nano- (% of TCRFL)	pico- (% of TCRFL)
19/06/2015	1	Bottom	55	33	12
	1	SCM	95.8	3.7	0.5
	1	Surface	36.1	42.8	21.1
	2	Bottom	27.6	47.1	25.3
	2	SCM	55.2	41.3	3.5
	2	Surface	11.9	70.1	18.1
23/06/2015	3	SCM	42.6	55.2	2.2
	4	Bottom	32	50	18
	4	SCM	66.6	30.9	2.5
	4	Surface	17	61.8	21.2
	5	SCM	73	23.1	3.8
	6	SCM	85.3	12.6	2.2
	7	SCM	92	6.3	1.7
	8	SCM	87.2	10.9	1.9
	9	SCM	89	5.3	5.7
24/06/2015	10	Bottom	14.9	50.4	34.7
	10	SCM	81.7	16.9	1.4
	10	Surface	21.3	59.6	19.2
	11	SCM	88.4	9.2	2.4
	12	SCM	88.2	9.7	2
	13	Bottom	20.7	55.6	23.7
	13	SCM	73.7	21.5	4.8
	13	Surface	10.7	68.2	21.2
25/06/2015	14	SCM	50	41.1	9
	15	SCM	54.9	40.7	4.4
	16	Bottom	22.4	35.7	42
	16	SCM	60.8	31.9	7.3
	16	Surface	32.2	44.1	23.6
	17	SCM	78.5	19	2.6
	18	SCM	89.8	8.4	1.8
26/06/2015	19	Bottom	21	45.2	33.8
	19	SCM	91.7	7.1	1.2
	19	Surface	40.2	44.7	15
	20	SCM	82.6	8.2	9.2
	21	SCM	88	10.7	1.2
	22	SCM	60.9	31.8	7.3
	23	SCM	66.1	28.5	5.4

	24	SCM	80.9	14.8	4.2
	25	SCM	91.9	6.1	2
	26	SCM	94.6	4.1	1.2
	27	SCM	79.6	14.4	6.1
	28	SCM	72.2	22.5	5.2
	29	Bottom	21.5	23.3	55.3
27/06/2015	29	SCM	81	14.6	4.4
	29	Surface	41.9	43.9	14.2
	30	SCM	92.4	7	0.7
	31	SCM	56.5	35.8	7.7
	32	SCM	78.7	18.2	3.1
	33	Bottom	20.1	41.6	38.3
	33	SCM	73	21.6	5.4
29/06/2015	33	Surface	14.6	36.7	48.6
	34	SCM	67.2	23.7	9.2
	35	SCM	42.4	50.9	6.6
	36	SCM	59.3	36.4	4.3
	37	Bottom	22.8	23.1	54.1
30/06/2015	37	SCM	76.1	21.7	2.2
	37	Surface	25.8	66.1	8.1
01/07/2015	38	SCM	67.1	28.2	4.7
	39	SCM	42.7	54.5	2.8
	40	Bottom	17.9	39.5	42.6
02/07/2015	40	SCM	54.1	34.1	11.8
	40	Surface	23	59.9	17.1

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Table 2. Contribution of carbon biomass by diatoms, dinoflagellates, flagellates, ciliates, non-flagellated chlorophyceae, and key taxa within these major groups in the surface layer, SCM and bottom mixed layer of the 39 stratified sites sampled within the Western English Channel excluding the holocam deployment site (site 1). Ranges and mean values given.

		Contribution to biomass (%)	
		Range	Mean
Surface	Diatoms	10.8 – 34.9	18.3
	→ *P.truncata*	→ 0.0 – 9.8	4.6
	→ *Chaetoceros spp.*	→ 0.1 – 6.0	1.9
	→ *Leptocylindrus spp.*	→ 1.8 – 20.6	7.1
	→ *P.alata/Rhizosolenia spp.*	→ 0.6 – 4.2	1.9
	→ *Thalassiosira spp.*	→ 0.1 – 3.0	0.7
	Dinoflagellates	27.4 – 45.0	35.3
	→ Small naked dinoflagellates	→ 9.3 – 25.2	15.5
	→ *Gyrodinium + Ceratoperidinium spp.*	→ 1.6 – 10.3	3.9
	Flagellates	1.9 – 6.6	3.9
	Ciliates (mostly aloricate)	32.4 – 53.9	41.9
	Non-flagellated chlorophyceae	0.0 – 6.0	0.6
SCM	Diatoms	0.8 – 22.7	7.2
	→ P.*truncata*	→ 0.0 – 13.7	3.8
	Dinoflagellates	20.5 – 89.0	64.3
	→ *C.fusus*	→ 1.3 – 80.9	39.6
	→ *C.lineatum*	→ 0.0 – 7.3	2.5
	→ *Gyrodinium + Ceratoperidinium spp.*	→ 1.3 – 7.4	3.1
	→ *Protoperidinium spp.*	→ 0.2 – 4.8	1.9
	→ *Dinophysis spp.*	→ 0.0 – 39.1	5.0
	→ Small naked dinoflagellates	→ 1.5 – 23.9	6.3
	Flagellates	1.1 – 50.1	7.8
	→ *Dictyocha spp.*	→ 0.3 – 49.0	5.9
	Ciliates (mostly aloricate)	4.4 – 54.9	20.7
	Non-flagellated chlorophyceae	0.0 – 3.4	0.1
Deep	Diatoms	14.2 – 52.7	37.0
	→ *P.truncata*	→ 4.0 – 43.3	22.4
	→ *Chaetoceros spp.*	→ 1.2 – 9.7	4.4
	Dinoflagellates	24.7 – 51.5	35.4
	→ *C.fusus*	→ 0.0 - 8.7	4.2
	→ Small naked dinoflagellates	→ 12.8 – 40.0	23.2
	Flagellates	4.5 – 15.1	8.8
	→ *Dictyocha spp.*	→ 2.3 – 6.1	3.4
	Ciliates (aloricate and loricate)	8.4 – 29.4	18.3
	Non-flagellated chlorophyceae	0.0 – 2.1	0.4

Table 3. The five greatest contributors to similarity within each cluster, where cluster 1 (C1) contains deep samples, cluster 2 (C2) contains SCM samples and cluster 3 (C3) contains surface samples. Average similarity within each cluster is also given.

Top five contributors to similarity (with % contributions)	Deep (C1)	SCM (C2)	Surface (C3)
1.	10-20 µm naked dinoflagellates (12.14)	*Ceratium fusus* (16.75)	M aloricate ciliates (11.87)
2.	*Proboscia truncata* (11.65)	L aloricate ciliates (8.82)	L aloricate ciliates (9.72)
3.	*Chaetoceros spp.* (5.22)	10-20 µm naked dinoflagellates (5.46)	10-20 µm naked dinoflagellates (8.39)
4.	*Dictyocha spp.* (5.09)	*Proboscia truncata* (5.14)	*Gyrodinium spp.* (5.81)
5.	L aloricate ciliates (4.82)	*Dictyocha spp.* (4.99)	20-25 µm naked dinoflagellates (5.41)
Cumulative contribution (%)	38.91	41.16	41.21
Average similarity (%)	67.48	67.30	71.96

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Figure 1.JPEG

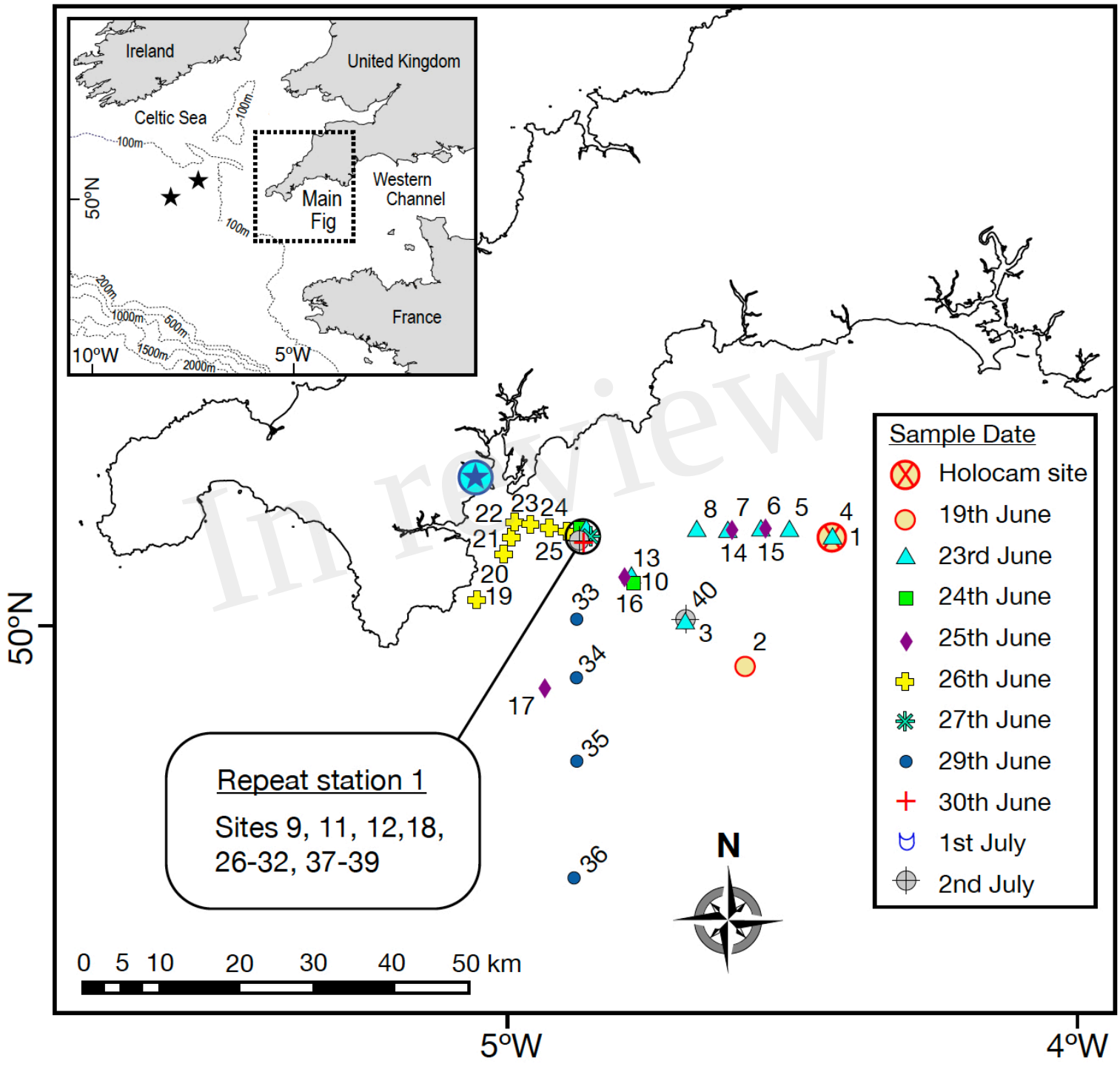


Figure 2.JPEG

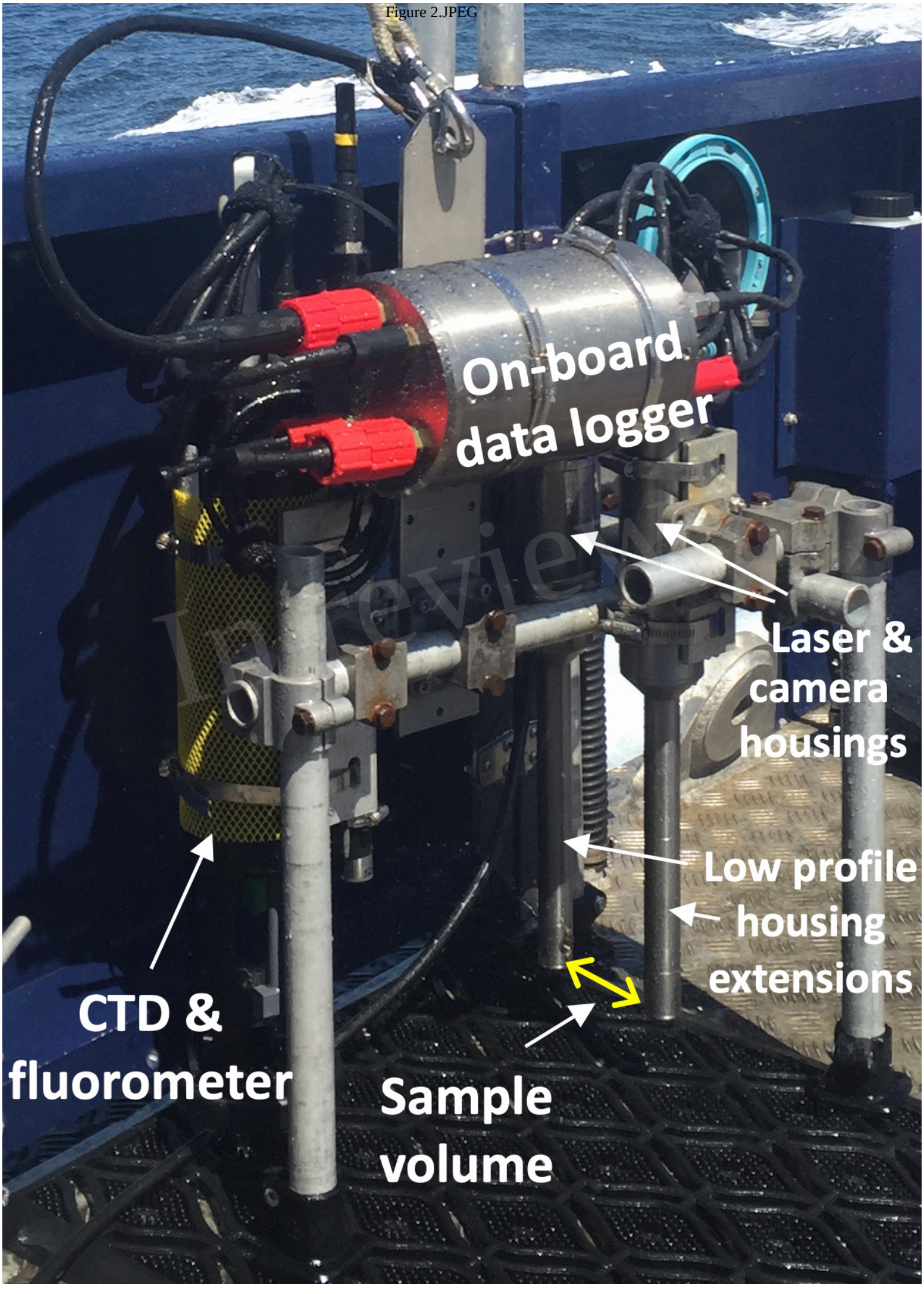


Figure 3.JPEG

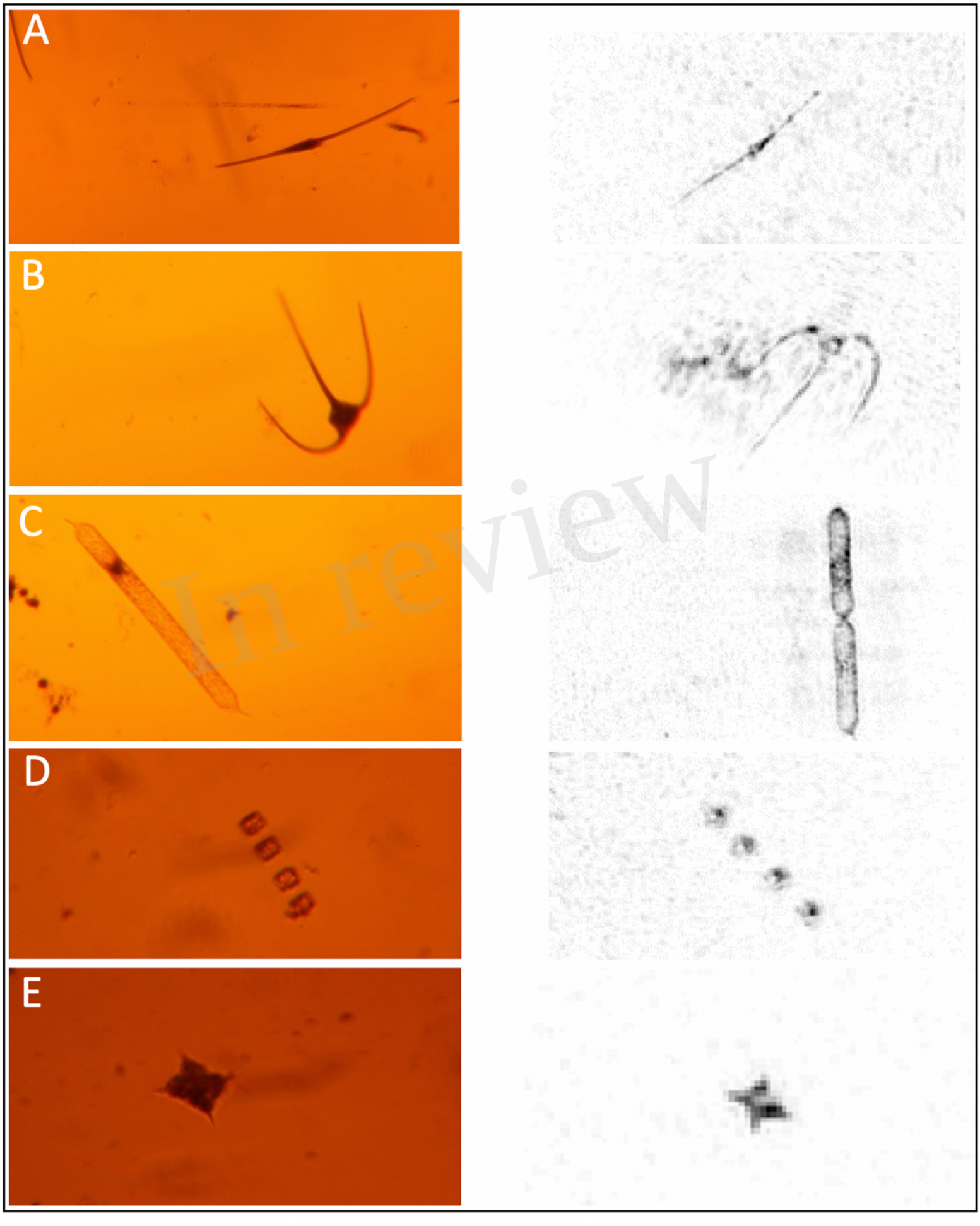


Figure 4.JPEG

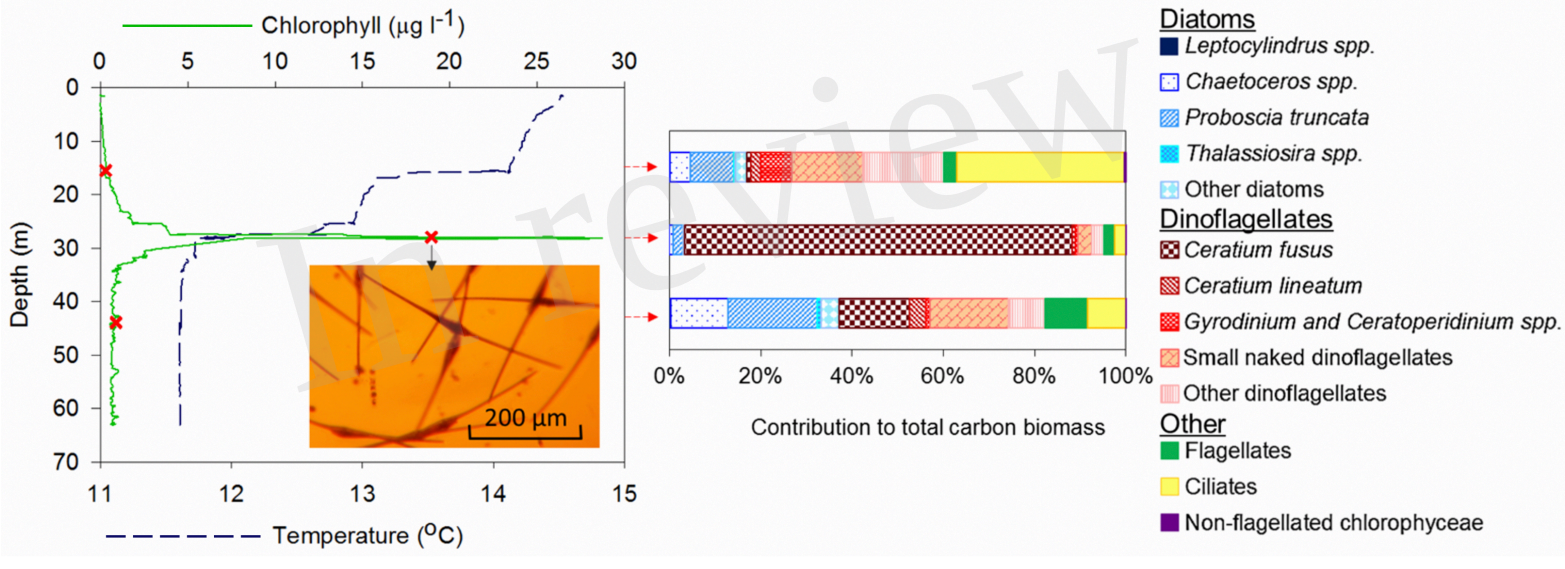


Figure 5.JPEG

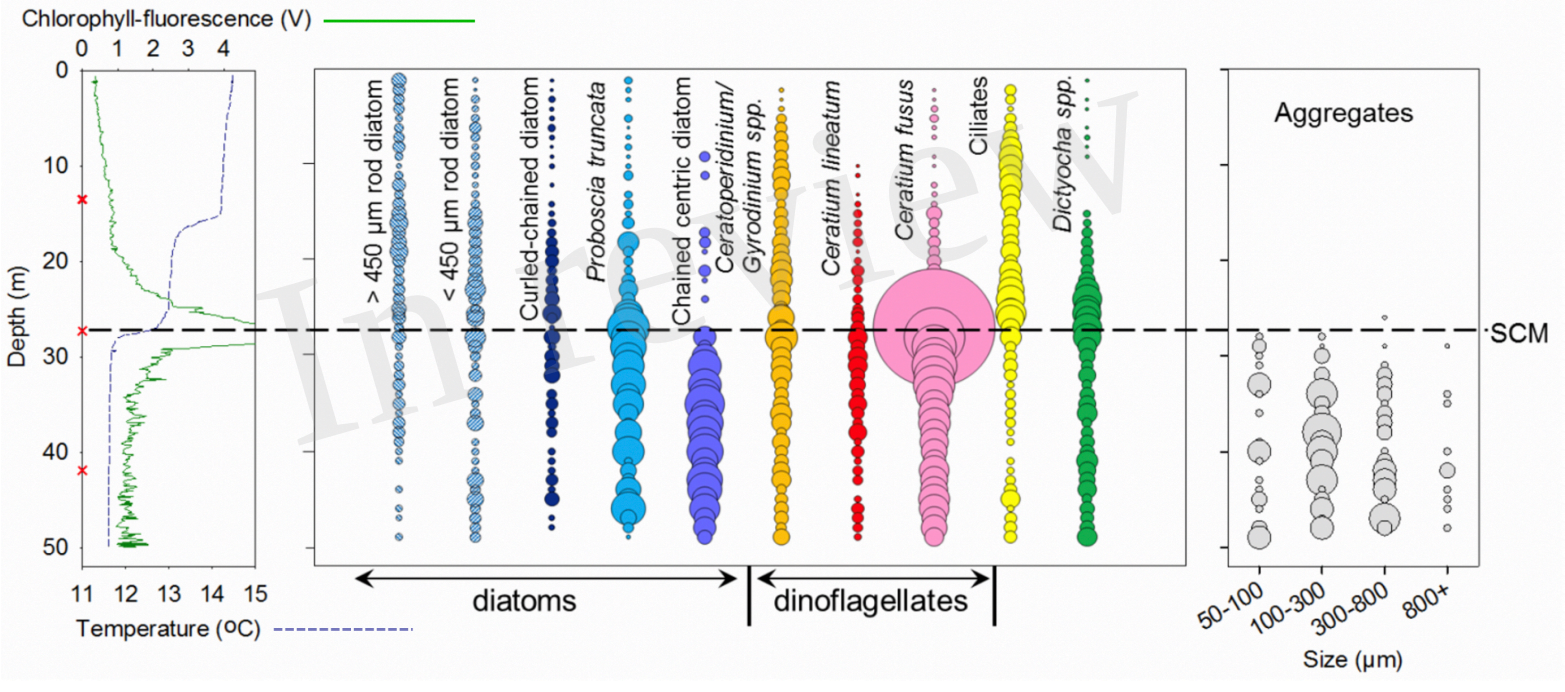


Figure 6.JPEG

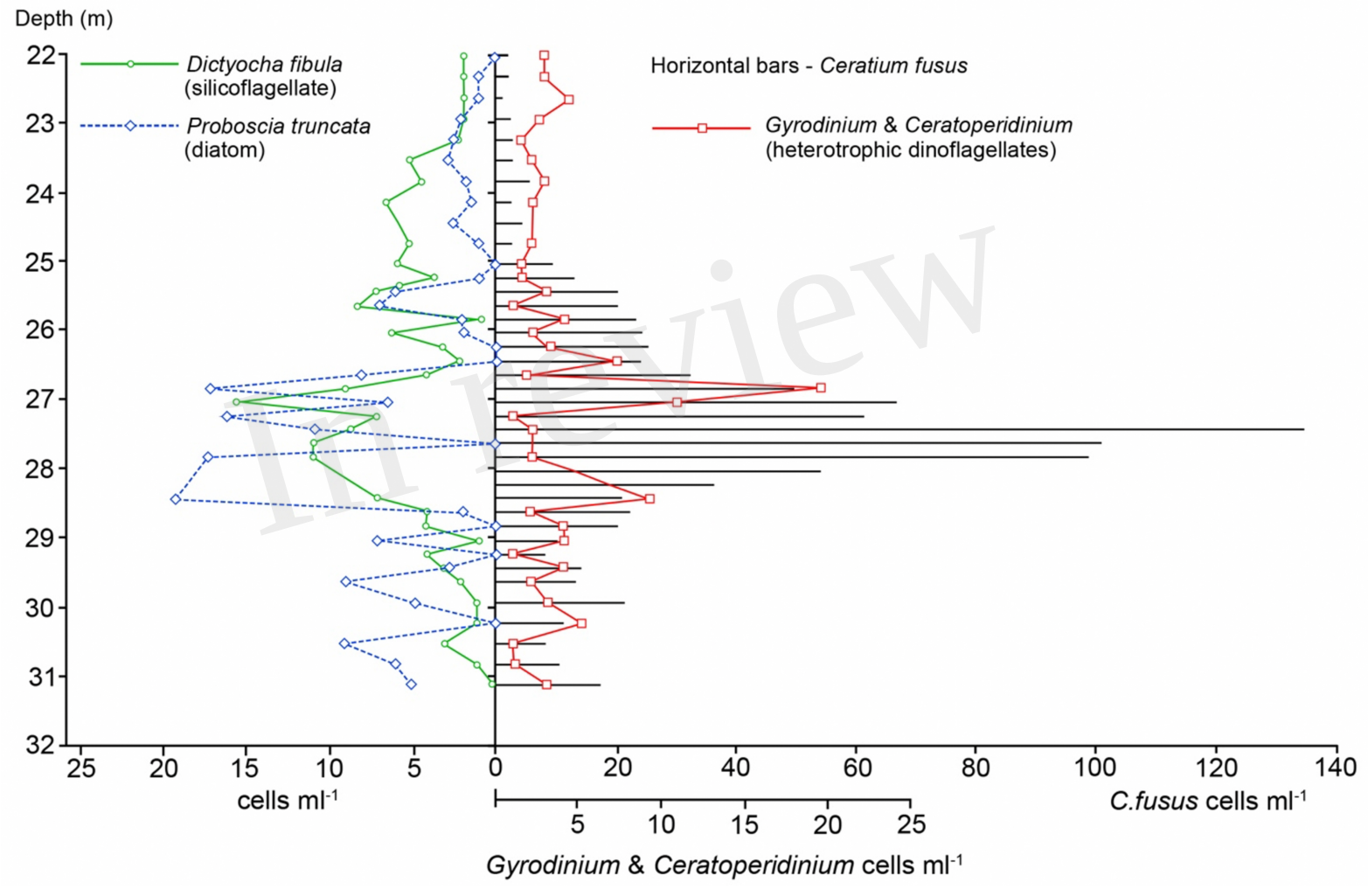


Figure 7.JPEG

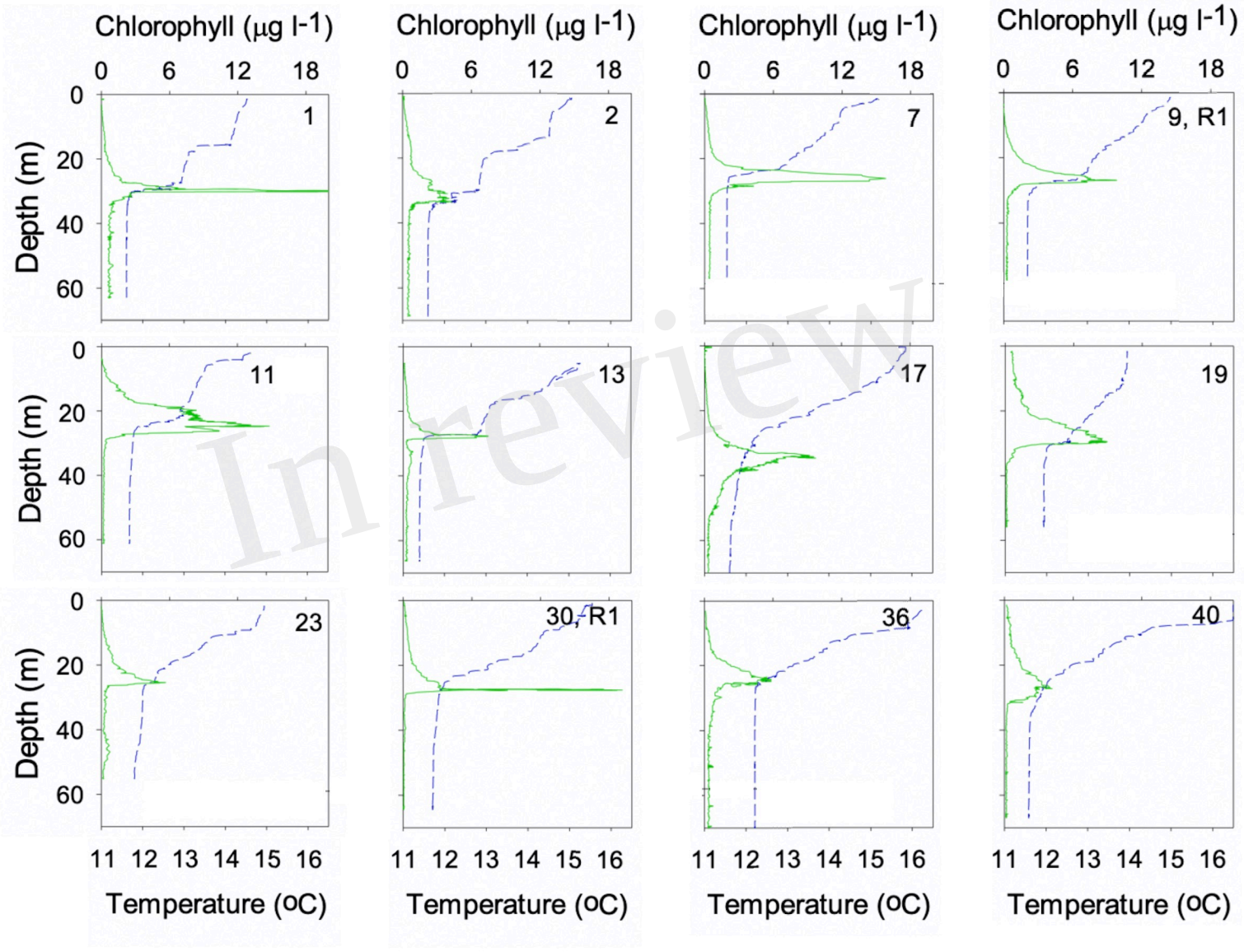


Figure 8.JPEG

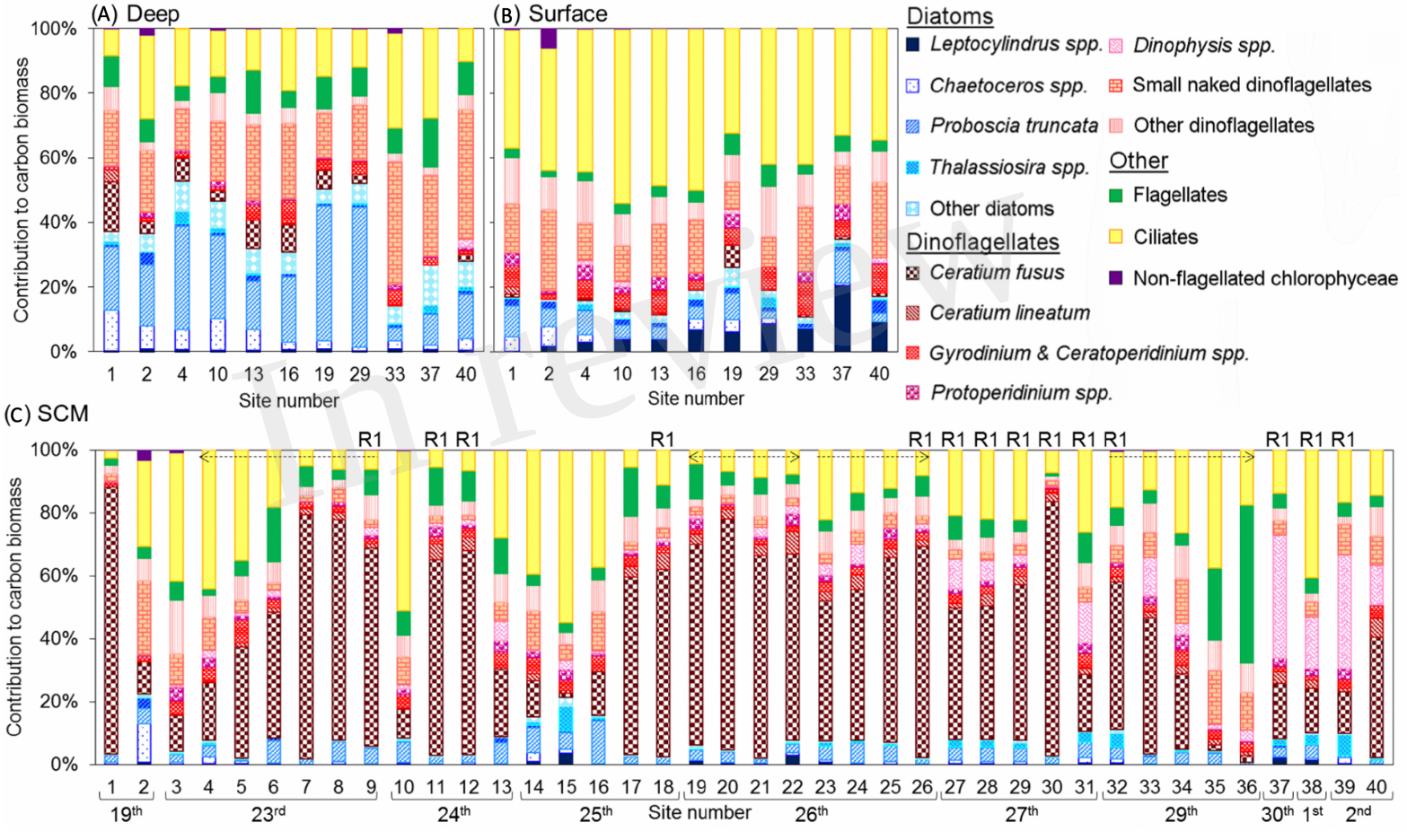


Figure 9.JPEG

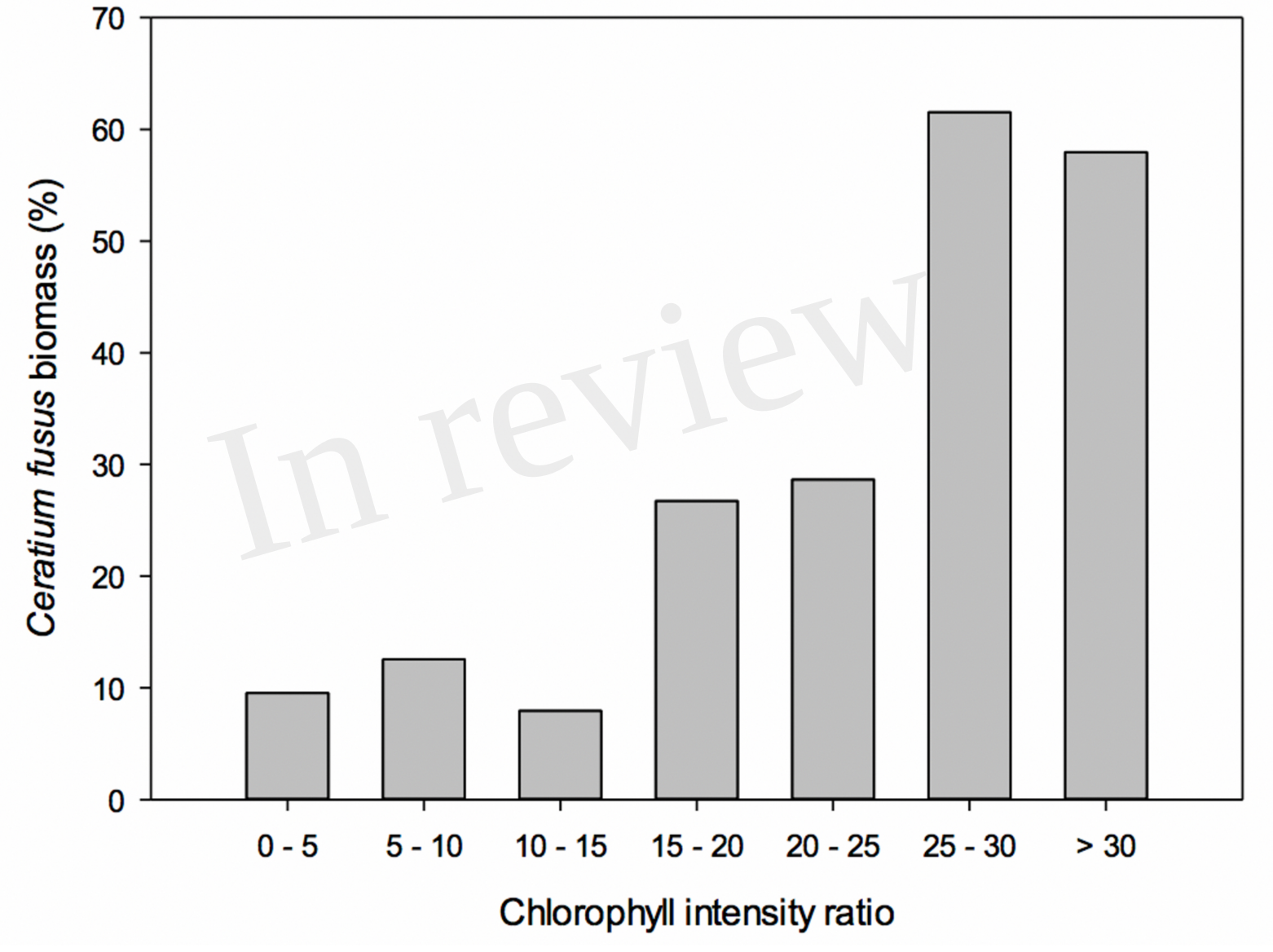


Figure 10.JPEG

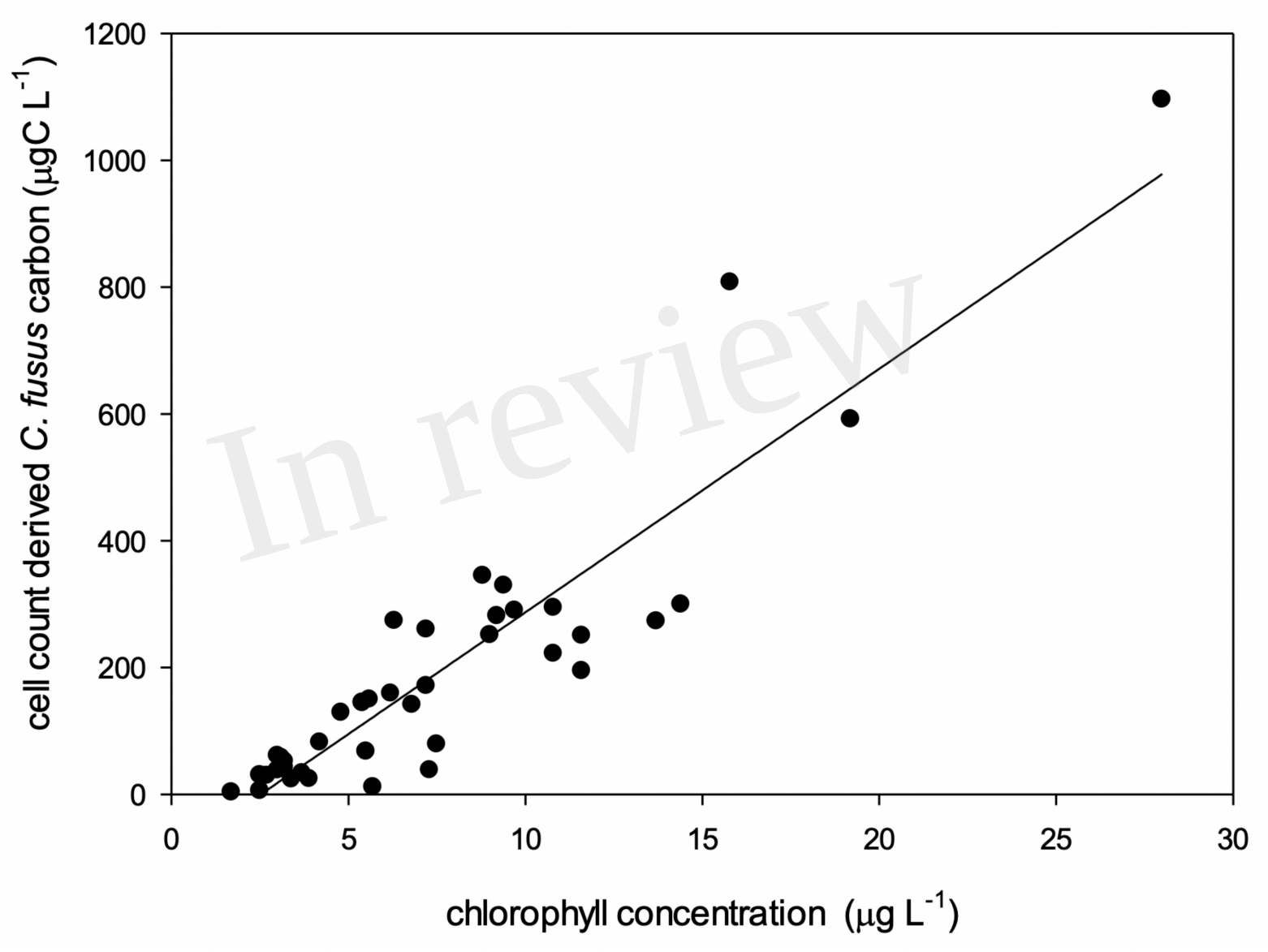
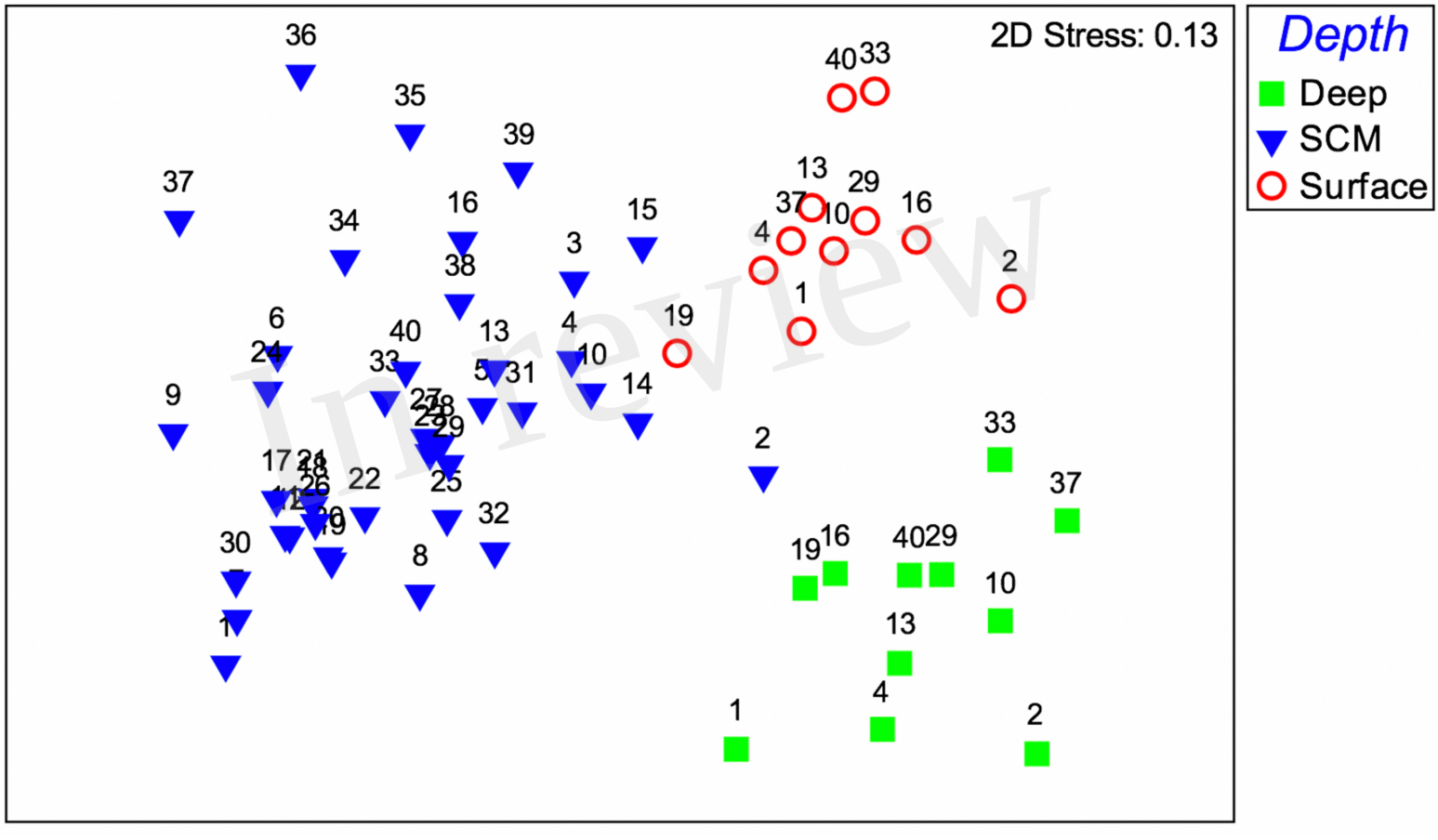
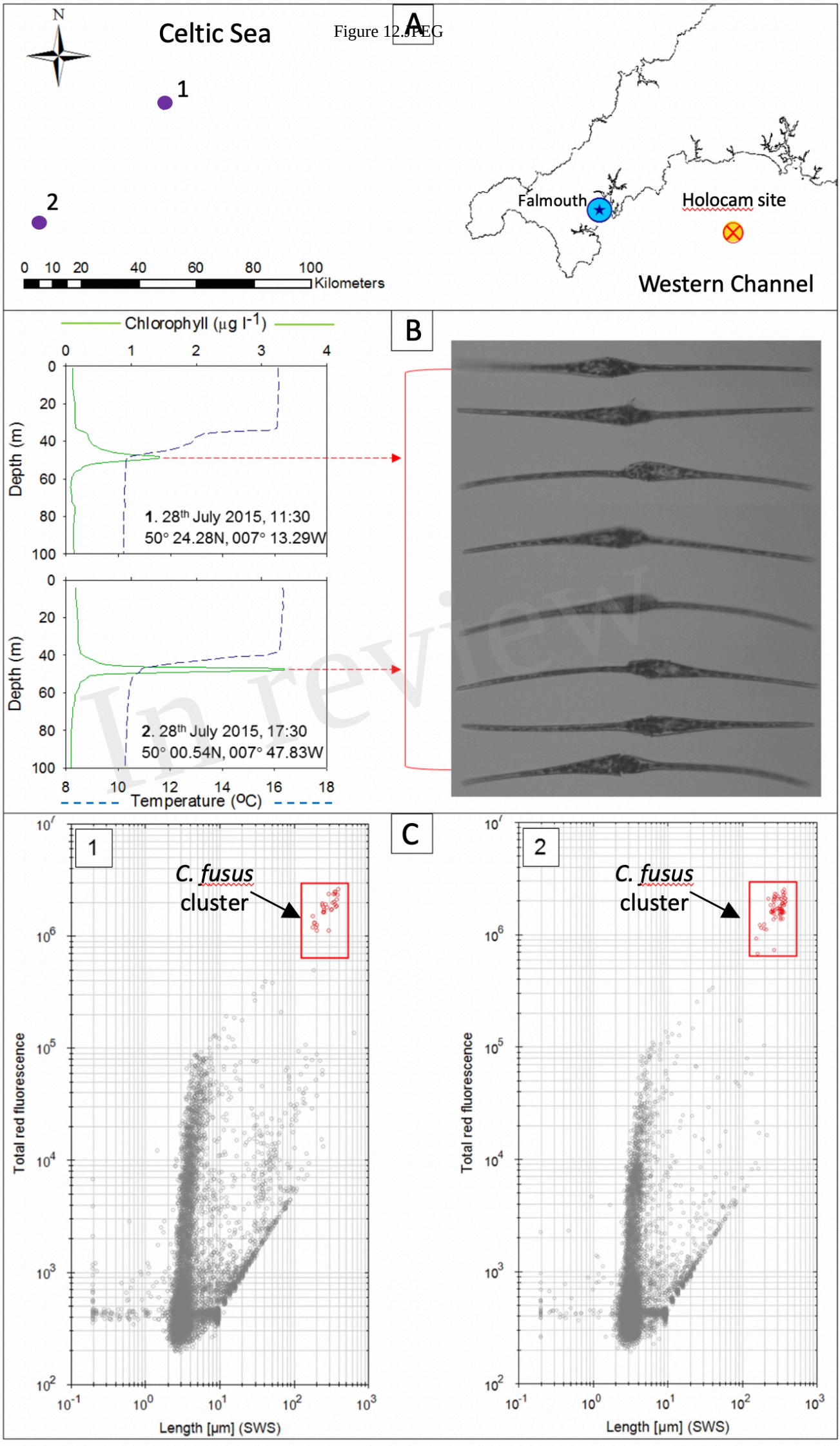


Figure 11.JPEG




A
Celtic Sea
1
2
N
0 10 20 40 60 80 100
Kilometers
Falmouth
Holocam site
Western Channel
B
Chlorophyll (µg l-1)
Depth (m)
1. 28th July 2015, 11:30
50° 24.28N, 007° 13.29W
2. 28th July 2015, 17:30
50° 00.54N, 007° 47.83W
Temperature (°C)
C
1
2
C. fusus
cluster
Total red fluorescence
Length [µm] (SWS)